



SOURCES AND PROPERTIES OF AMAZONIAN AEROSOL PARTICLES

Scot T. Martin,¹ Meinrat O. Andreae,² Paulo Artaxo,³ Darrel Baumgardner,⁴ Qi Chen,¹ Allen H. Goldstein,⁵ Alex Guenther,⁶ Colette L. Heald,⁷ Olga L. Mayol-Bracero,⁸ Peter H. McMurry,⁹ Theotonio Pauliquevis,¹⁰ Ulrich Pöschl,² Kimberly A. Prather,^{11,12} Gregory C. Roberts,¹³ Scott R. Saleska,¹⁴ M. A. Silva Dias,¹⁵ Dominick V. Spracklen,¹⁶ Erik Swietlicki,¹⁷ and Ivonne Trebs²

Received 8 December 2008; revised 7 August 2009; accepted 22 September 2009; published 16 April 2010.

[1] This review provides a comprehensive account of what is known presently about Amazonian aerosol particles and concludes by formulating outlook and priorities for further research. The review is organized to follow the life cycle of Amazonian aerosol particles. It begins with a discussion of the primary and secondary sources relevant to the Amazonian particle burden, followed by a presentation of the particle properties that characterize the mixed populations present over the Amazon Basin at different times and places. These properties include number and mass concentrations

and distributions, chemical composition, hygroscopicity, and cloud nucleation ability. The review presents Amazonian aerosol particles in the context of natural compared to anthropogenic sources as well as variability with season and meteorology. This review is intended to facilitate an understanding of the current state of knowledge on Amazonian aerosol particles specifically and tropical continental aerosol particles in general and thereby to enhance future research in this area.

Citation: Martin, S. T., et al. (2010), Sources and properties of Amazonian aerosol particles, *Rev. Geophys.*, 48, RG2002, doi:10.1029/2008RG000280.

TABLE OF CONTENTS

1. Introduction	2	3.3. Chemical Composition	24
2. Sources	8	3.4. Hygroscopicity	29
2.1. Primary Particles	8	3.5. Cloud Condensation Nuclei	29
2.2. Secondary Gas-to-Particle Conversion	10	4. Outlook and Future Priorities	31
3. Properties	18	4.1. Priorities for Improved Models	34
3.1. Mass Concentration	18	4.2. Priorities for Improved Measurements	34
3.2. Number-Diameter Distribution	21	5. Concluding Remarks	35

¹School of Engineering and Applied Sciences and Department of Earth and Planetary Sciences, Harvard University, Cambridge, Massachusetts, USA.

²Biogeochemistry Department, Max Planck Institute for Chemistry, Mainz, Germany.

³Institute of Physics, University of São Paulo, São Paulo, Brazil.

⁴Centro de Ciencias de la Atmosfera, Universidad Nacional Autonoma de Mexico, Mexico City, Mexico.

⁵Department of Environmental Science, Policy, and Management, University of California, Berkeley, California, USA.

⁶Biosphere-Atmosphere Interactions Group, Atmospheric Chemistry Division, National Center for Atmospheric Research, Boulder, Colorado, USA.

⁷Department of Atmospheric Science, Colorado State University, Fort Collins, Colorado, USA.

⁸Institute for Tropical Ecosystem Studies, University of Puerto Rico, San Juan, Puerto Rico.

⁹Department of Mechanical Engineering, University of Minnesota, Minneapolis, Minnesota, USA.

¹⁰Climate and Environmental Modelling Group, Instituto Nacional de Pesquisas da Amazônia, Manaus, Brazil.

¹¹Department of Chemistry and Biochemistry, University of California, San Diego, La Jolla, California, USA.

¹²Scripps Institution of Oceanography, University of California, San Diego, La Jolla, California, USA.

¹³Center for Atmospheric Sciences, Scripps Institution of Oceanography, University of California, San Diego, La Jolla, California, USA.

¹⁴Department of Ecology and Evolutionary Biology, University of Arizona, Tucson, Arizona, USA.

¹⁵Center for Weather Forecasting and Climate Studies, National Institute for Space Research, University of São Paulo, São Paulo, Brazil.

¹⁶Institute for Climate and Atmospheric Science, School of Earth and Environment, University of Leeds, Leeds, UK.

¹⁷Department of Physics, Lund University, Lund, Sweden.

1. INTRODUCTION

[2] Aerosol particles in the Amazon Basin have been the focus of numerous field campaigns over the past 20 years (Table 1 and Figure 1). These studies were motivated by a wide range of objectives, the most prominent of which are as follows: (1) The Basin was used as a laboratory to gain baseline knowledge concerning pristine continental aerosol particles, against which the effects of human activities globally could be judged [Andreae, 2007]. (2) An understanding was sought of the effects of biomass-burning aerosol particles on human health, such as increased incidences of morbidity, mortality, and asthma [Ignotti et al., 2007]. (3) The effects of aerosol particles on regional climate were investigated, such as changes in rainfall patterns as a consequence of the redistribution of energy and cloud condensation nuclei [Andreae et al., 2004]. (4) The Basin was studied as an integrated ecosystem to understand the feedback and regulation of plant emissions on rainfall and, in turn, of rainfall on plant growth and emissions [Barth et al., 2005; Keller et al., 2009]. These topics have in common a need to know the sources and properties of Amazonian aerosol particles, yet an integrated summary of results from previous field campaigns (Table 1) and associated statements of future research priorities have not been prepared previously in a comprehensive review article. This gap in the literature is the motivation for this review and defines its scope: the review's goals are to focus the ongoing activities of researchers already investigating the sources and properties of Amazonian aerosol particles and, by organizing and presenting material about what is already known and what remains to be learned, to invite new researchers to join in critical ways. Complementary general reviews of atmospheric particles, especially with regard to the organic component that is dominant in the Amazon Basin, are given by Andreae and Crutzen [1997], Jacobson et al. [2000], Kanakidou et al. [2005], and Fuzzi et al. [2006].

[3] Aerosol sources located within the Amazon Basin are dominated, with the exception of some urbanized areas and transportation corridors, by natural and anthropogenic emissions from the biosphere. Sources include both high but intermittent biomass-burning emissions (both natural and anthropogenic) and low but more consistent production of primary and secondary biological aerosol particles and components (Figure 2). Primary particles are produced both deliberately by flora (e.g., the release of pollen and fungal spores) and incidentally (e.g., as leaf and soil debris or as suspended microbes). Substantial production of secondary aerosol occurs by the atmospheric oxidation of trace gases to low-volatility compounds. These products can deposit on preexisting particles or possibly nucleate new particles.

[4] Once in the atmosphere, particles undergo continuous transformations (Figure 3). Processes include (photo)chemical reactions that occur between compounds within the particles as well as interactions that occur between compounds within the particles and those in the gas phase, such as the condensation of low-volatility compounds or reactions with highly reactive gaseous species like the OH radical. Clouds are present at varying abundances almost

everywhere and all the time over the Amazon Basin, and most particles also undergo several cycles of cloud processing during their residence in the Basin. The time scale of cloud cycling of boundary layer air (of the order of hours) is considerably shorter than the residence time of air over the Basin or the deposition lifetime of aerosol particles (of the order of days). Cloud processing can modify particle properties both by chemical reactions in the liquid phase and by interactions between droplets (e.g., collision and coagulation). Particles leave the Amazonian atmosphere by dry deposition to the vegetation surface, by cloud scavenging and precipitation, and by advection out of the region.

[5] In addition to particle sources within the Basin, there are also important and at times dominant long-range natural and anthropogenic sources (Figure 4). The influence of long-range transport is particularly important when in-Basin sources are weak, such as in the wet season, and under these conditions the particle population can be dominated at times both in mass and number by outflow from other areas. The Atlantic Ocean, upwind of the Basin, is a strong source of marine particles that are generated both directly by sea spray as well as indirectly by the conversion of gases, such as the oxidation of dimethyl sulfide (DMS) to form sulfate. Across the Atlantic and farther upwind of the Basin, the Saharan desert is the world's largest source of mineral dust. Sub-Saharan Africa is one of the most important sources of smoke from vegetation fires. Although the inflowing air masses that arrive with the trade winds from the Atlantic shed much of their particle burden in transit from Africa and Eurasia, the impact of transatlantic transport on particles measured in the Basin can, nevertheless, be substantial and at times dominant [Prospero et al., 1981; Andreae et al., 1990a; Artaxo et al., 1990; Swap et al., 1992; Formenti et al., 2001; Chen et al., 2009]. Furthermore, pollution-derived particles from urban and industrialized areas in southern and eastern Brazil and other South American countries can also be transported into the Basin, especially in the dry season.

[6] During the wet season (December–March, i.e., summer of the Southern Hemisphere), atmospheric particles are removed relatively quickly by wet deposition, and anthropogenic sources such as biomass burning are weak throughout the Basin. This combination of circumstances results in natural processes (including contributions from marine and African sources) as the dominant contributors to the ambient particle populations over large expanses of the Basin and during a significant part of the year. The particle concentrations measured during these conditions are among the lowest found on any continent and are similar to those over the remote oceans [Andreae, 2009]. The Basin has been dubbed the “green ocean” because of the similarities in particle concentrations and cloud microphysics between it and remote oceanic regions [Williams et al., 2002]. The Amazon Basin may be the only region on the tropical continents where there remains the possibility to find at times populations of nearly pristine aerosol particles free of direct anthropogenic influences.

TABLE 1. Major International Measurement Campaigns of Amazonian Aerosol Particles^a

Dates	Acronym	Campaign	Location and Description	Ground	Airborne	Science Focus	Reference
Jul–Aug 1985	ABLE-2A	Amazon Boundary Layer Experiment: dry season	dry season in Amazonia	Y (Manaus)	Y (ER-2 and Electra)	trace gas and meteorological measurements	Harriss <i>et al.</i> [1988]
Apr–May 1987	ABLE-2B	Amazon Boundary Layer Experiment: wet season	wet season in Amazonia	Y (Manaus)	Y (ER-2 and Electra)	trace gas and meteorological measurements	Harriss <i>et al.</i> [1990]
Aug–Sep 1995	SCAR-B	Smoke, Clouds, and Radiation–Brazil	dry season in disturbed southern Amazonia (Brasília and Alta Floresta in southern Amazonia)	Y	Y (ER-2 and Convair)	effects of biomass burning on climate (direct effect)	Kaufman <i>et al.</i> [1998]
Mar–Apr 1998	CLAIRE-98	Cooperative LBA Airborne Regional Experiment	wet season in undisturbed central Amazonia (Balbina in central Amazonia)	Y		chemistry and physics of biogenic aerosols	Avissar <i>et al.</i> [2002]
Apr–May 1999 and Sep–Oct 1999	LBA-EUSTACH	European Studies on Trace Gases and Atmospheric Chemistry as a Contribution to the Large-Scale Biosphere–Atmosphere Experiment in Amazonia	wet season in disturbed western Amazonia (FNS and RBJ in Rondônia in southwestern Amazonia, K34 and CI4 in central Amazonia, and Caxiuanã in eastern Amazonia)	Y	Y (Bandeirante)	chemistry and physics of wet season biogenic aerosols and trace gas emissions and reactivity	Andreae <i>et al.</i> [2002]
Jul 2001	CLAIRE-01	Cooperative LBA Airborne Regional Experiment	transition of wet to dry season in undisturbed central Amazonia (Balbina)	Y	Y (Bandeirante)	chemistry and physics of biogenic aerosols and trace gas emissions and reactivity	Fuzzi <i>et al.</i> [2007]
Sep–Nov 2002	LBA-SMOCC	Large-Scale Biosphere–Atmosphere Experiment in Amazonia: Smoke, Aerosols, Clouds, Rainfall, and Climate Field Campaign	dry season biomass burning experiment in pasture and forest sites in disturbed Amazonia (FNS and RBJ Rondônia)	Y	Y (Bandeirante)	measurements of aerosol and trace gas emission factors from biomass burning	Yokelson <i>et al.</i> [2007]
Aug–Sep 2004	TROFEE	Tropical Forest and Fire Emissions Experiment	dry season in disturbed southern Amazonia (Alta Floresta)		Y (Bandeirante)	measurements of aerosol and trace gas emission factors from biomass burning	S. T. Martin <i>et al.</i> (manuscript in preparation, 2010)
Feb–Mar 2008	AMAZE-08	Amazonian Aerosol Characterization Experiment	wet season in undisturbed central Amazonia (TT34 north of Manaus in central Amazonia)	Y	N	chemistry and physics of biogenic aerosols and trace gas emissions and reactivity	

^aBeginning in the 1980s, large-scale measurement campaigns focusing on the atmospheric sciences began in the Amazon Basin, and many of them had at least a component devoted to aerosol research. Over the same time span, these limited-term campaigns have been complemented by some extended measurement series at fixed ground stations. Key: Y, yes; N, no; FNS, Fazenda Nossa Senhora Aparecida; RBJ, Reserva Biológica Jari.

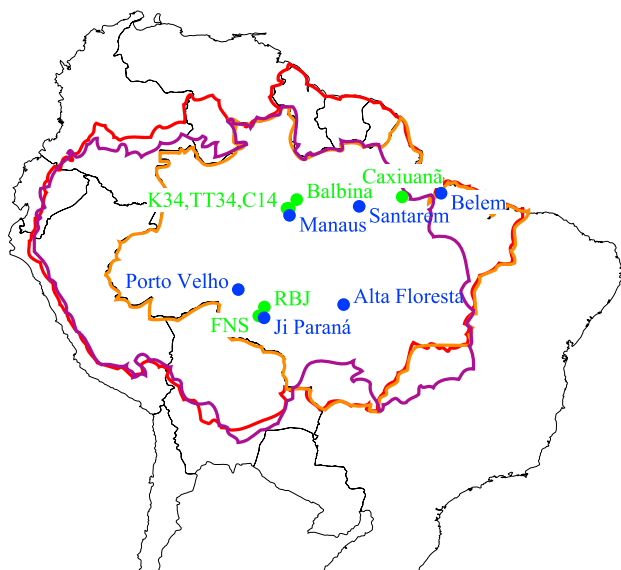


Figure 1. Map showing the geographic boundaries of the Amazon forest (red line), the Amazon-Tocantins river Basins (purple line), and Brazilian Legal Amazon (orange line). The political boundaries of South America are shown as black lines. Cities often mentioned in the literature of Amazonian aerosol particles are indicated. Specific research sites for some of the campaigns listed in Table 1 are also highlighted, including (1) Fazenda Nossa Senhora Aparecida (FNS) and Reserva Biológica Jarú (RBJ) in Rondônia in southwestern Amazonia; (2) K34, TT34, and C14 north of Manaus in central Amazonia; (3) Balbina in central Amazonia; and (4) Caxiuana in eastern Amazonia. The map covers 30°S–15°N and 81°W–35°W.

[7] Compared to these green-ocean conditions of the wet season, there is stark contrast in the dry season for large regions of the Basin (June–September, i.e., winter of the Southern Hemisphere). Vast numbers of deforestation fires burn during the dry season, especially along the peripheries of the forest, and large parts of the Basin become among the most polluted places on Earth [Artaxo et al., 2002; Cardoso et al., 2003]. As one effect, the regional energy balance is changed because the high particle concentrations affect the amount and location of solar radiation absorbed by the planet. Simulations using regional climate models show that the changes in energy delivery significantly influence regional patterns of atmospheric circulation and meteorology [Zhang et al., 2009]. The high particle concentrations change cloud microphysics and rainfall, with a significant influence on the overall water cycle [Andreae et al., 2004; Intergovernmental Panel on Climate Change, 2007; Rosenfeld et al., 2008]. They also influence air quality by degrading visibility and affecting human health [Reinhardt et al., 2001; Schwartz et al., 2002; Watson, 2002; Pope and Dockery, 2006; Barregard et al., 2006].

[8] The meteorology of South America guides the transport of particle-free or particle-laden air masses into the Amazon Basin. It also affects the rates of wet and dry par-

ticle deposition. A brief introduction to the meteorology of South America is therefore provided herein, in the next few paragraphs. A more in-depth presentation is given by Satyamurty et al. [1998], and an introduction to the regional climate of the Amazon Basin (including a history of field campaigns focused on meteorology) is given by Nobre et al. [2004].

[9] The convective activity and the atmospheric circulation of tropical South America are part of a monsoon system [Zhou and Lau, 1998]. The seasonal cycle of circulation and convection is revealed by changes both in the low-level wind field (see Figures 5a and 5b) and in the outgoing longwave radiation (see Figures 5c and 5d) for December, January, and February (DJF) compared to June, July, and August (JJA) [Kalnay et al., 1996]. In both the wet and dry

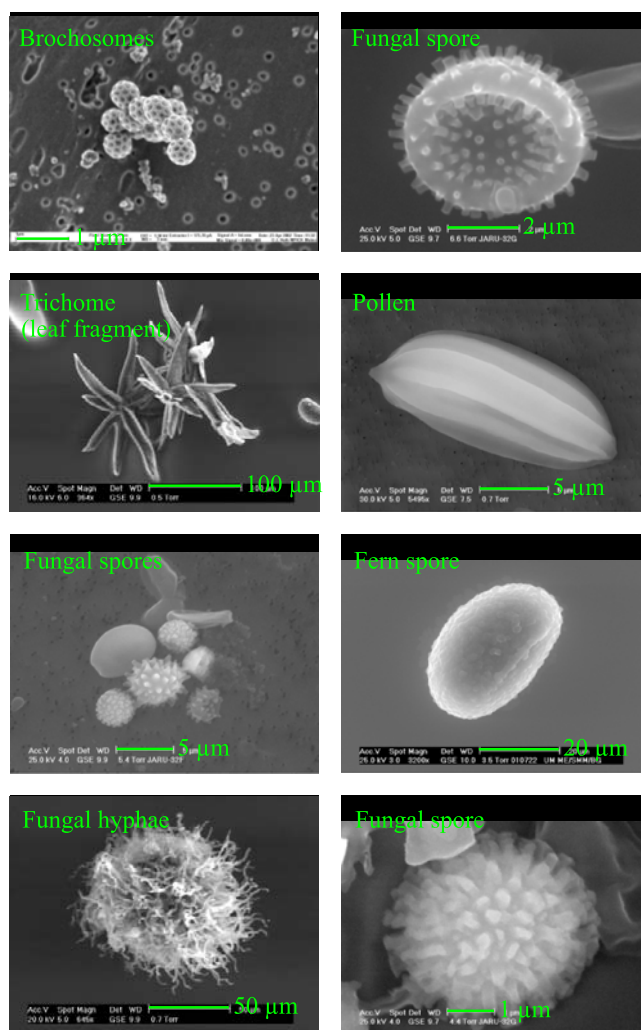


Figure 2. Scanning electron micrographs of primary biological particles collected in the Amazon Basin. From unpublished results of the Max Planck Institute for Chemistry, Mainz, Germany, for particles collected during the project European Studies on Trace Gases and Atmospheric Chemistry as a Contribution to the Large-Scale Biosphere-Atmosphere Experiment in the Amazon Basin (LBA-EUSTACH) in 1999.

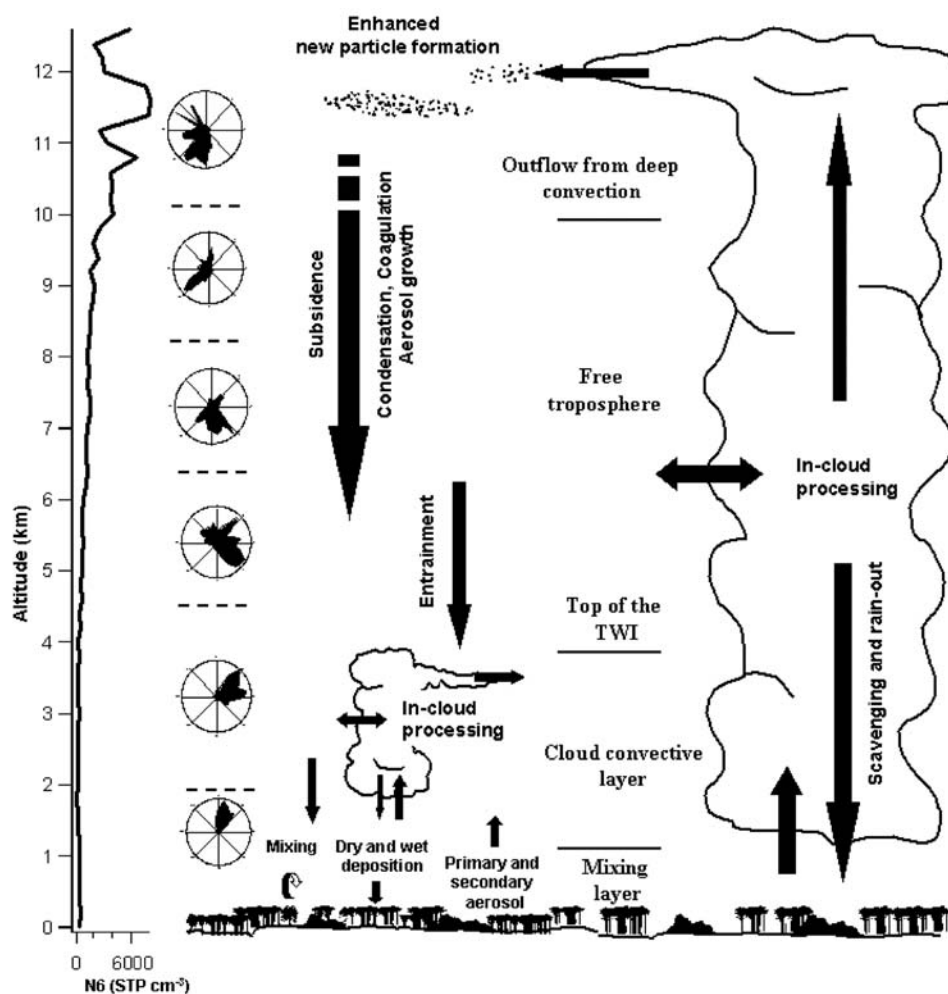


Figure 3. Illustration of vertical mixing processes that affect the particle number-diameter distribution of aerosol particles in the Amazon Basin. The region of the lowest 4 km represents daytime conditions with a fully developed mixing layer and shallow convection in the transition layer. Upward transport is controlled by deep convection and fair-weather cumulus clouds. Subsidence dominates the large-scale downward transport. Figure 3 is based on *Krejci et al.* [2003, Figure 9]; the original intent was to describe observations over Suriname, but the processes depicted are applicable to the wider Amazon Basin.

seasons, there is considerable inflow coming from the east into the Basin. These air masses originate in the Atlantic, and many of them pass through the semiarid region of northeastern Brazil before entering central and southern regions of the Basin [*Satyamurty et al.*, 1998].

[10] Several aspects of Figure 5 are important for understanding the variability of particle concentrations and compositions observed in the Amazon Basin. For northern and central Amazonia, the Intertropical Convergence Zone (ITCZ), which is the confluence between the northeastern and southeastern trade winds extending from West Africa to South America (Figures 5a and 5b), has an important influence. The ITCZ reaches the northern coast of South America just south of the equator in DJF but just north of it in JJA. When the ITCZ is located to the south of the equator in DJF (i.e., a large part of the Basin is under the influence of air from the Northern Hemisphere), low-level winds from

the global northeast reach the northern coast of South America and open up the possibility of advection of African dust and biomass-burning particles [e.g., *Artaxo and Hansson*, 1995; *Formenti et al.*, 2001]. During the wet season, Figure 5c shows that the outgoing longwave radiation has a minimum over central Amazonia, implying cold cloud tops and hence deep clouds with strong convection and high rates of wet deposition. There is also strong vertical transport and redistribution of particles [*Freitas et al.*, 2000; *Andreae et al.*, 2001]. In comparison, during the dry season the convection weakens and shifts to the north-western edge of geographical Amazonia (Figure 5d).

[11] In the southern part of the Amazon Basin, north-westerlies prevail during the wet season and are associated with the South Atlantic Convergence Zone (SACZ), a region of enhanced convective activity extending from central Amazonia to the southeast [*Kodama*, 1992]. The SACZ

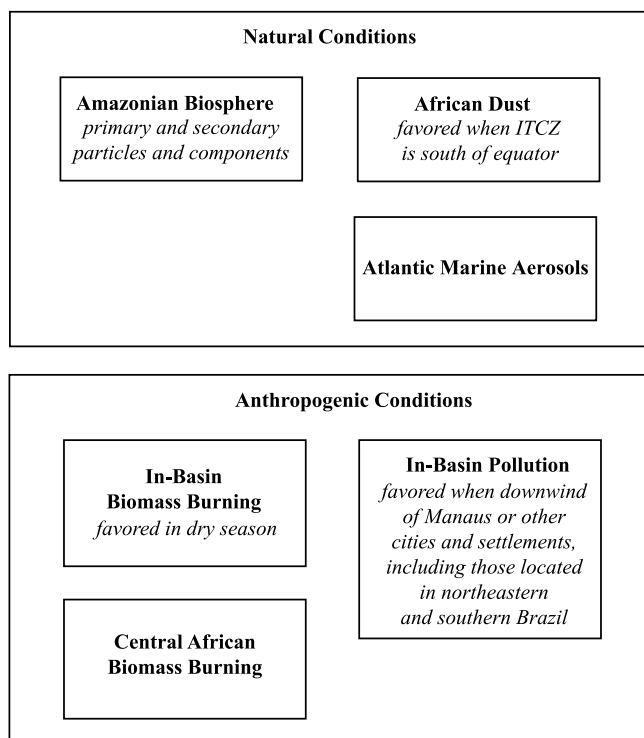


Figure 4. Source classification scheme for Amazonian aerosol particles. Although emissions from the Amazonian biosphere are active at all times and have low variability year round, they are relatively weak, and particles and their components can be dominated at times by influences from outside of the Amazon Basin, such as from Saharan dust, African biomass burning, or Atlantic marine emissions. The dust and marine emissions are a natural contribution because they were present in the year 1750. At other times, anthropogenic influences such as in-Basin biomass burning in the dry season can dominate the type, the number, and the mass concentrations of Amazonian aerosol particles. “Favored” suggests conditions of greater probability, although all influences are possible under most conditions. For example, in the wet season biomass burning can still influence some observations, such as sampling sites downwind of the border area of Brazil and Guyana and Suriname, which engage in biomass burning during the northern Amazonian wet season (see Figure 8).

forms a quasi-stationary front with significant intraseasonal variability. As the remains of midlatitude fronts reach the tropical region [Garreaud and Wallace, 1998], the SACZ is reinforced. In the absence of perturbations coming from the south, the SACZ weakens and may disappear. This variability leads to a change in low-level winds from easterlies in the break periods to westerlies in the periods of well-defined SACZ. The active and break periods are related to intraseasonal oscillations [Nogues-Paegle and Mo, 1997] that affect the whole region from the northern coast [Wang and Fu, 2002] to the more continental area [Jones and Carvalho, 2002]. During the break periods, deep convective systems are commonly isolated, and (given the absence of sinks) particle concentrations are relatively high. By

comparison, during SACZ events convection and rainfall are more widespread, the atmosphere is clean, and particle concentrations are relatively low [Silva Dias et al., 2002; Williams et al., 2002].

[12] With the coming of the dry season, the low-level winds in the southern part of the Basin reverse, changing from northwesterlies to southeasterlies (Figures 5a and 5b). These patterns favor the flow of pollution from urban and industrial Brazil into this region of Amazonia, thereby increasing particle concentrations. Alternatively, cold-front southerlies (i.e., winter in the Southern Hemisphere) can bring clean air into the southern part of Amazonia, in some cases extending as far as the equator [Marengo et al., 1997]. At these times, particle concentrations in southern Amazonia can drop considerably. More typically, however, biomass burning in the Basin leads to sustained high particle concentrations regardless of southerlies or southeasterlies [Artaxo et al., 2002].

[13] The diel evolution of the planetary boundary layer (see Figure 6) also affects particle concentrations measured at the surface. During the afternoon, the boundary layer is well mixed by strong turbulence that is driven by sensible heat flux from the surface. The depth of this convective boundary layer (CBL) is variable depending on land cover and on the meteorology. By late afternoon CBL depths >1000 m are typical in the Amazon Basin, although variability is high. Fisch et al. [2004] observed differences between the wet and dry seasons (e.g., lower than 1500 m in the former and up to 2000 m in the latter) and between forest and pasture landscapes. At sunset, radiative cooling at the surface generates a nocturnal stable layer. The nocturnal boundary layer usually has a depth of a few hundred meters or less. The residual layer above the nocturnal stable layer is typically without turbulence or mixing, although there are exceptions at times when higher-altitude shearing jets are present that induce turnover and hence the cleansing of the residual layer with cleaner higher-altitude air. In the absence of cleansing, the nighttime residual layer conserves the properties (such as particle and gas concentrations) of the previous afternoon until the following morning.

[14] These basic characteristics of the daytime and nighttime boundary layers affect particle concentrations measured at the surface. During the afternoon, emissions from the surface get mixed through the whole volume of the convective boundary layer, diluting their concentrations. Moreover, at the top of the boundary layer, turbulence in fair weather conditions maintains an entrainment of cleaner air from higher altitudes while convective clouds associated with rainy weather pump the aerosol particles to higher levels. These effects favor a dip of particle concentrations in the early afternoon. In comparison, the stable nocturnal boundary layer traps emissions in a volume near the surface, thereby favoring higher particle concentrations at night. This effect is amplified during the dry season because biomass burning usually begins at midday and continues into the evening hours. The residual layer above the nocturnal boundary layer influences the surface concentrations the following morning because the development of the con-

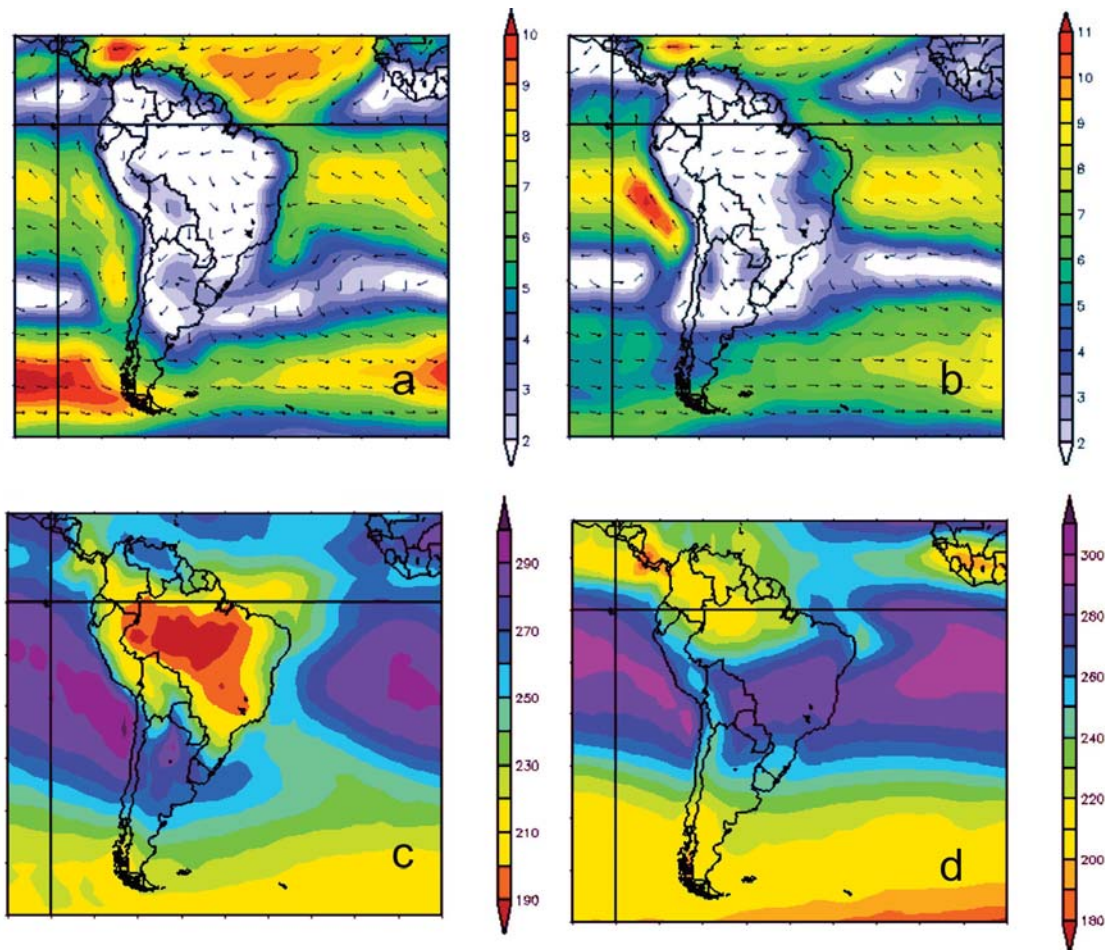


Figure 5. (a) December, January, and February (DJF) wind vector and wind speed (m s^{-1}) at 1000 hPa for South America. (b) Same as Figure 5a but averaged for June, July, and August (JJA). (c) Outgoing longwave radiation for DJF (W m^{-2}). (d) Same as Figure 5c but for JJA. The horizontal black line shows the position of the equator. Data represent the average reanalysis of 1988–2007 from the National Centers for Environmental Prediction.

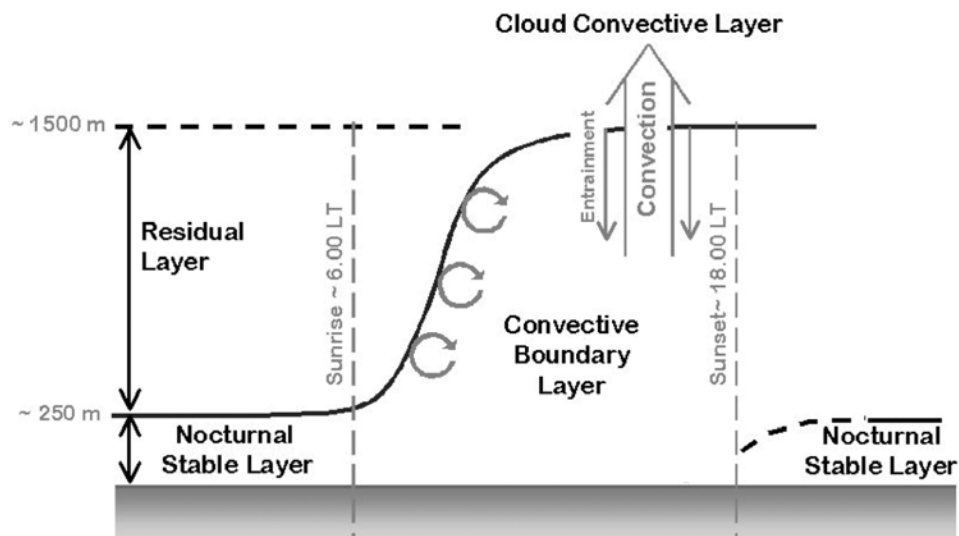


Figure 6. Illustration of the typical diel evolution of the lower troposphere in the Amazon Basin. Adapted from Rissler *et al.* [2006].

vective boundary layer mixes the nocturnal layer into the residual layer. As a result, middle to late morning surface concentrations (in the absence of nighttime cleansing of the residual layer) can be similar to nighttime concentrations measured the afternoon before in the convective boundary layer [Rissler et al., 2006].

[15] Given the seasonal and regional variability of the contributions from different aerosol sources, the changing transport paths of air masses, and the different removal rates in the dry and wet seasons, the inference (which is supported by the observations) is that Amazonian aerosol particles have considerable variations in space and time, and consequently, there is a considerable body of literature to review. Our approach in this review is to follow the life cycle of Amazonian aerosol particles, as outlined in the previous paragraphs. The review begins with a discussion of the primary and secondary sources relevant to the Amazonian particle burden, followed by a presentation of the particle properties that characterize the mixed populations present over the Amazon Basin at different times and places. These properties include number and mass concentrations and distributions, chemical composition, hygroscopicity, and cloud nucleation ability. The review presents Amazonian aerosol particles in the context of natural compared to anthropogenic sources as well as variability with season and meteorology. The review concludes with an outlook and priorities for further research.

2. SOURCES

[16] Amazonian aerosol particles have a wide range of natural and anthropogenic sources, and the integrated effects of emission and processing in the atmosphere lead to complex internal and external mixtures of particles, even within an apparently homogeneous air mass. The complexity can be usefully dissected by conceptualizing a single particle as composed of components. Depending on their origin, components are usually classified as primary or secondary. Primary components are directly emitted from a source into the atmosphere; secondary components are formed in the atmosphere [Fuzzi et al., 2006]. A single particle composed mainly of primary components can be called a primary aerosol particle, and a single particle composed mainly of secondary components can be called a secondary aerosol particle. After some air mass aging, many, if not most, individual particles can be composed of significant quantities of both types of components. In the Amazon Basin, organic components typically constitute ~70%–90% of the particle mass concentration in both the fine and coarse size fractions [Graham et al., 2003a; Fuzzi et al., 2007].

[17] Examples of primary biological aerosol (PBA) particles emitted in the Amazon Basin include pollen, bacteria, fungal and fern spores, viruses, and fragments of plants and animals [Elbert et al., 2007]. Anthropogenic biomass burning is also an important and at times dominant source at some locations, especially during the dry season. In addition to sources within the Amazon Basin, primary par-

ticles are also brought in by long-range transport, such as marine particles from the Atlantic Ocean and desert dust or biomass-burning particles from Africa [Andreae et al., 1990a; Artaxo et al., 1990; Swap et al., 1992; Formenti et al., 2001; Chen et al., 2009]. Regional urban and industrial activities, including traffic and industry in Manaus and other cities and settlements in northeastern and southern Brazil, also have outflow plumes containing combustion-derived particles and dust, and these plumes are significant when a sampling location lies within them.

[18] Examples of components of secondary organic aerosol (SOA) are the low-volatility molecules that result from the reactions of O₃ and OH with biogenic volatile organic compounds (BVOCs), such as isoprene and terpenes. BVOCs are emitted in large quantities to the gas phase by plants. Low-volatility BVOC oxidation products can condense from the gas phase onto preexisting particles or alternatively can contribute to new-particle formation. Liquid-phase reactions inside cloud droplets can also yield low-volatility BVOC oxidation products, serving as another source of organic components in particles for cloud droplets that evaporate [Blando and Turpin, 2000; Lim et al., 2005; Carlton et al., 2006].

2.1. Primary Particles

2.1.1. Primary Biological Particles

[19] Emissions of primary biological particles are often wind-driven, such as suspension of pollen, plant debris, or soil dust [Jaenicke, 2005; Pöschl, 2005]. In addition to wind-driven release, certain biological organisms also actively eject materials into the air for reproductive purposes, such as wet-discharged fungal spores [Elbert et al., 2007]. Coarse-mode PBA particles in the Amazon Basin have sizes ranging from several to tens of micrometers and include fragments of plants and insects, pollen grains, algae, fern spores, and fungal spores [Graham et al., 2003a]. Microscopic analyses of collected particles show that in the absence of African dust and Atlantic marine emissions, morphologically identifiable biological particles dominate both the number-diameter and volume-diameter distributions of the coarse fraction for natural conditions (Figure 7). PBA components like carbohydrates, proteins, and lipids, as well as elemental tracers, have also been detected in the fine fraction [Artaxo and Hansson, 1995; Andreae and Crutzen, 1997; Graham et al., 2003a]. Observations during the Amazonian Aerosol Characterization Experiment 2008 (AMAZE-08) by Sinha et al. [2009] and Chen et al. [2009], however, suggest that PBA components contribute in a minor way to the size class below 1 μm. Nevertheless, the actual number, mass, and size of PBA particles emitted to the fine fraction remain to be fully quantified for Amazonia, and the abundance and composition of PBA particles are highly variable and still poorly characterized, partially because the distinction between biological and other carbonaceous components requires advanced analytical techniques and intensive investigation [Pöschl, 2005; Fuzzi et al., 2006; Despres et al., 2007].

[20] Fungi are an especially important source of coarse-mode PBA particles in the Basin. They actively discharge

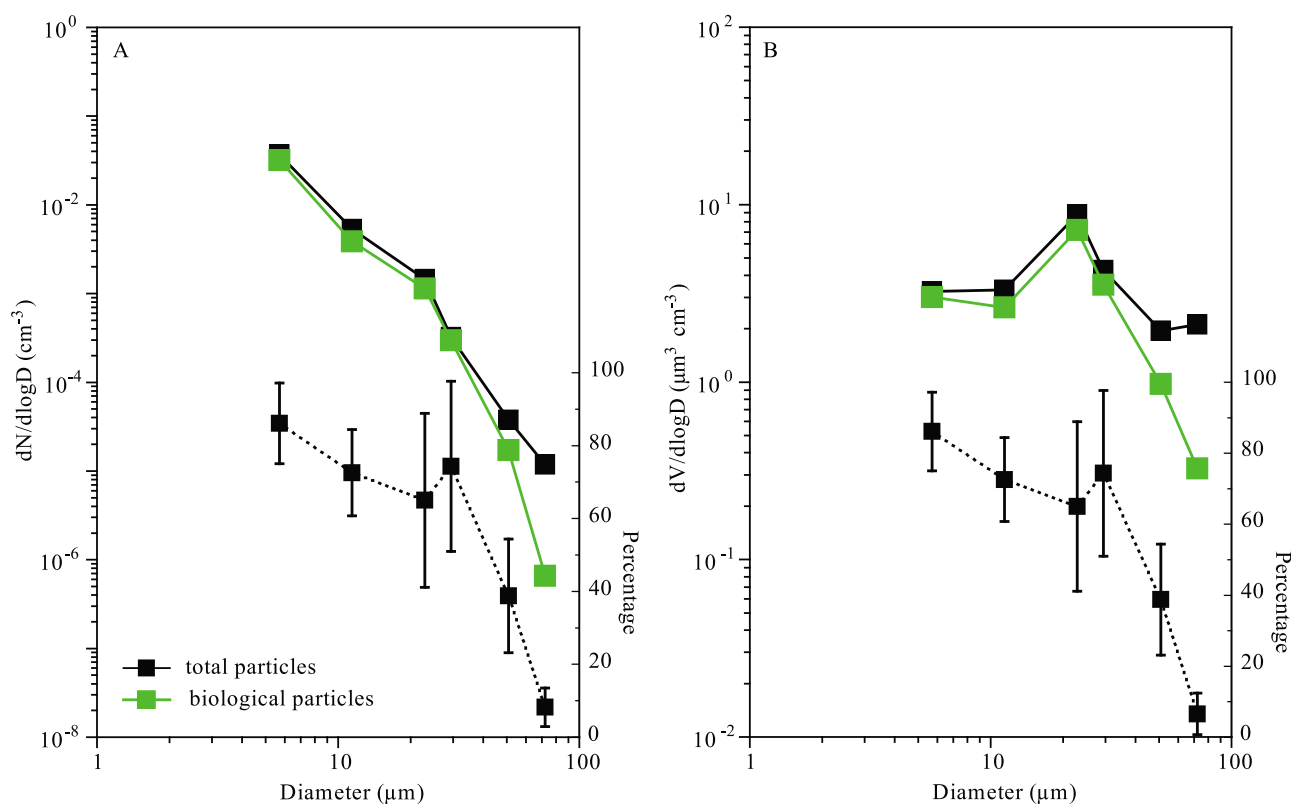


Figure 7. (a) Number-diameter and (b) volume-diameter distributions for giant Amazon aerosol particles ($>4 \mu\text{m}$), as determined by light microscopic analysis of samples collected on glass slides. Shown are particle totals, biological particle totals, and percentage contribution of biological particles to the total. Data are from M. Andreae (personal communication, 2009).

their spores with liquid jets and droplets into the air, and these processes are most active under humid conditions, such as those in the rain forest [Gilbert, 2005]. For particles of 1–10 μm , fungi are estimated to contribute 25% of the particles during the day and 45% at night. In corroboration, measurements using an ultraviolet aerodynamic particle sizer deployed during AMAZE-08 indicate that fungal spores and other viable PBA particles account for up to 80% of coarse-mode particles during time periods of weak influence from sources outside of the Basin (U. Pöschl, manuscript in preparation, 2010). Actively wet spore discharging fungi also emit dissolved inorganic salts like potassium chloride and organic substances like sugar alcohols such as mannitol in aqueous jets, and these jets break up to form droplets that can dry as fine particles [Elbert et al., 2007].

2.1.2. Biomass Burning

[21] Anthropogenic aerosol particles are generated in great number and mass concentration during episodes of biomass burning employed for land clearing and pasture maintenance [Andreae et al., 1988; Artaxo et al., 1998; Andreae et al., 2002; Artaxo et al., 2002; Guyon et al., 2003b; Andreae et al., 2004; Freitas et al., 2005; Guyon et al., 2005; Fuzzi et al., 2007; Yokelson et al., 2007]. The resulting particles are predominantly carbonaceous, consisting largely of organic carbon but also with significant near-elemental soot carbon and minor amounts of inorganic materials. The particles

consist of a variety of internal and external mixtures, including soot carbon fractal aggregates, tarry materials (sometimes present as spherical “tar balls”), grains of inorganic salts (KCl, K₂SO₄, KNO₃, (NH₄)₂SO₄, etc.), and ash and char particles [Allen and Miguel, 1995; Andreae et al., 1998; Ferek et al., 1998; Yamasoe et al., 2000; Posfai et al., 2003, 2004].

[22] Emission factors from primary deforestation fires and pasture maintenance fires in tropical rain forests range from 6 to 25 g kg⁻¹ for total particulate matter (PM) and 7.5 to 15 g kg⁻¹ for PM smaller than 2.5 μm (PM_{2.5}), expressed as mass of emitted primary particles per mass unit of dry fuel. For Amazonia, the estimates for the emission rates of PM_{2.5} and PM₁₀ are 8 and 10 Tg yr⁻¹, respectively [Yokelson et al., 2008]. The pollution plumes can be hundreds of kilometers across and thousands of kilometers long. With an optical thickness regularly exceeding 1.0 in the peak period of biomass burning in the Basin (i.e., September and October), the plumes are clearly visible in satellite images, and biomass burning is the dominant particle source in affected areas. At those times, pollution from biomass smoke typically accounts for >90% of the fine particles and ~50% of the coarse particles. The annual mean of optical thickness over Amazonia and regions nearby is dominated by biomass-burning emissions [Tegen et al., 1997; Tie et al., 2005].

[23] Fire counts observed by satellite over the Amazon Basin in February, May, August, and November 2007 are shown in Figure 8. As expected, the most numerous and intense fires were in the dry season (i.e., August 2007) in the southern part of Amazonia. Fires were also important in northeastern Brazil in November 2007, and the prevailing flow patterns carried the biomass-burning emissions into the central and southern Amazon Basin (Figure 5). In the wet season, biomass burning took place along the northern rim of the Basin, and at times local meteorological variability transported the biomass-burning emissions into the Basin.

[24] In addition to biomass burning within South America, emissions from Africa are imported into the Amazon Basin at all times of the year [Talbot et al., 1990; Andreae et al., 1994]. Fires burn in tropical and subtropical Africa year round, with a maximum early in the year in the Northern Hemisphere and a maximum in the second half of the year in the Southern Hemisphere. At least some of the smoke from these fires is transported across the Atlantic by the trade winds [Andreae et al., 1994]. The particles emitted to the Northern Hemisphere enter northern Amazonia in the wet season and can be important, given the weak baseline production mechanisms of the Amazonian biosphere. In comparison, the contribution from the Southern Hemisphere enters southern Amazonia during the regional burning season of the latter, and the African contribution is therefore typically of less relative importance to Amazonia at that time. In summary, although the highest concentrations of biomass-burning particles are observed in southern Amazonia during the dry season, lower levels of biomass-burning particles can be important intermittently at any time and at most locations of the Basin throughout the year.

2.1.3. African Mineral Dust

[25] Saharan dust is a prominent out-of-Basin particle source. The importance of the transatlantic transport of dust was recognized by Prospero et al. [1981] and has been observed in several subsequent measurement campaigns [Swap et al., 1992; Artaxo et al., 1998; Formenti et al., 2001]. Imported dust occurs at its highest concentrations in those parts of the Basin that are north of the ITCZ. The maximum dust concentrations at the surface are typically reached around March and April, coinciding with the wet season in the central part of the Basin. A significant fraction of the delivered mineral dust is submicron, as explained by the large transport distance from Africa and the preferential loss of coarse-mode particles along the way. The dust is observed at near-surface stations in pulses of high concentrations that last from one to several days, and when present, mineral dust often dominates the total particle mass concentrations [Andreae et al., 1990b; Talbot et al., 1990; Formenti et al., 2001; Worobiec et al., 2007].

2.1.4. Marine Emissions

[26] Crossing the coast of the Amazon Basin with the trade wind flow, large concentrations of marine aerosol particles are progressively removed by wet and dry deposition as air masses travel deeper into the Basin [Andreae and Andreae, 1988; Talbot et al., 1988; Andreae et al., 1990a, 1990b; Talbot et al., 1990; Worobiec et al., 2007].

The contribution by marine particles to the total Amazonian particle mass concentration can remain significant even over the central parts of the Basin. Marine aerosol particles consist in large part of primary sea spray particles, which are composed mainly of coarse-mode inorganic salts mixed with lesser amounts of the primary biological material that was partitioned to the ocean surface [O'Dowd et al., 2004; Andreae and Rosenfeld, 2008]. The sodium and chloride content of the coarse fraction of the particle population in the Amazon Basin is explained almost entirely by marine sources. Marine aerosol particles also have a substantial contribution from secondary processes, such as sulfates produced by the oxidation of dimethyl sulfide and organic material produced by the oxidation of volatile organic compounds (VOCs) [Ceburnis et al., 2008]. Much of the secondary material occurs in the fine mode. In the absence of biomass-burning particles or dust, approximately half of the submicron sulfate fraction is attributable to secondary sulfate produced from marine emissions, especially from DMS oxidation [Andreae et al., 1990a; Worobiec et al., 2007].

[27] Priorities for progress in identifying the sources of primary particles in the Amazon Basin and quantifying their emissions include the following: (1) characterization and quantification of different types of primary biological, biomass burning, mineral dust, and marine aerosol particles, including long-term trends, seasonal cycles, and diel differences, and identification of their mixing states, including the relative contributions of primary and secondary components; (2) discrimination and quantification of the relative fractions of in-Basin and out-of-Basin sources of all particle types; (3) improved characterization and understanding of Amazonian aerosol particles by application of a combination of advanced measurement techniques, such as bulk and single-particle mass spectrometry, X-ray microanalysis, fluorescence spectroscopy, electron microscopy, and DNA analysis; and (4) development of process models describing the emission of primary biological particles from the Amazonian ecosystem and implementation of these process models in regional and global models of atmospheric chemistry, transport, and climate.

2.2. Secondary Gas-to-Particle Conversion

[28] The production mechanisms for secondary particle components involve many trace gases, in particular, biogenic volatile organic compounds, nitrogen oxides (NO_x), ozone (O_3), hydroxyl radical (OH), and sulfur species including DMS and sulfur dioxide (SO_2) [Andreae and Andreae, 1988; Jacob and Wofsy, 1988; Andreae et al., 1990a; Browell et al., 1990; Jacob and Wofsy, 1990; Kesselmeier et al., 2000; Andreae et al., 2002]. DMS and SO_2 are oxidized to form particle sulfate. BVOCs react with O_3 and OH to produce oxidized organic products, a fraction of which have low enough volatility to condense and serve as particle components. BVOCs and NO_x together affect the concentrations of O_3 and OH, thereby influencing the production of BVOC oxidation products. Reactions both in the gas phase and in cloud waters are important. BVOCs and

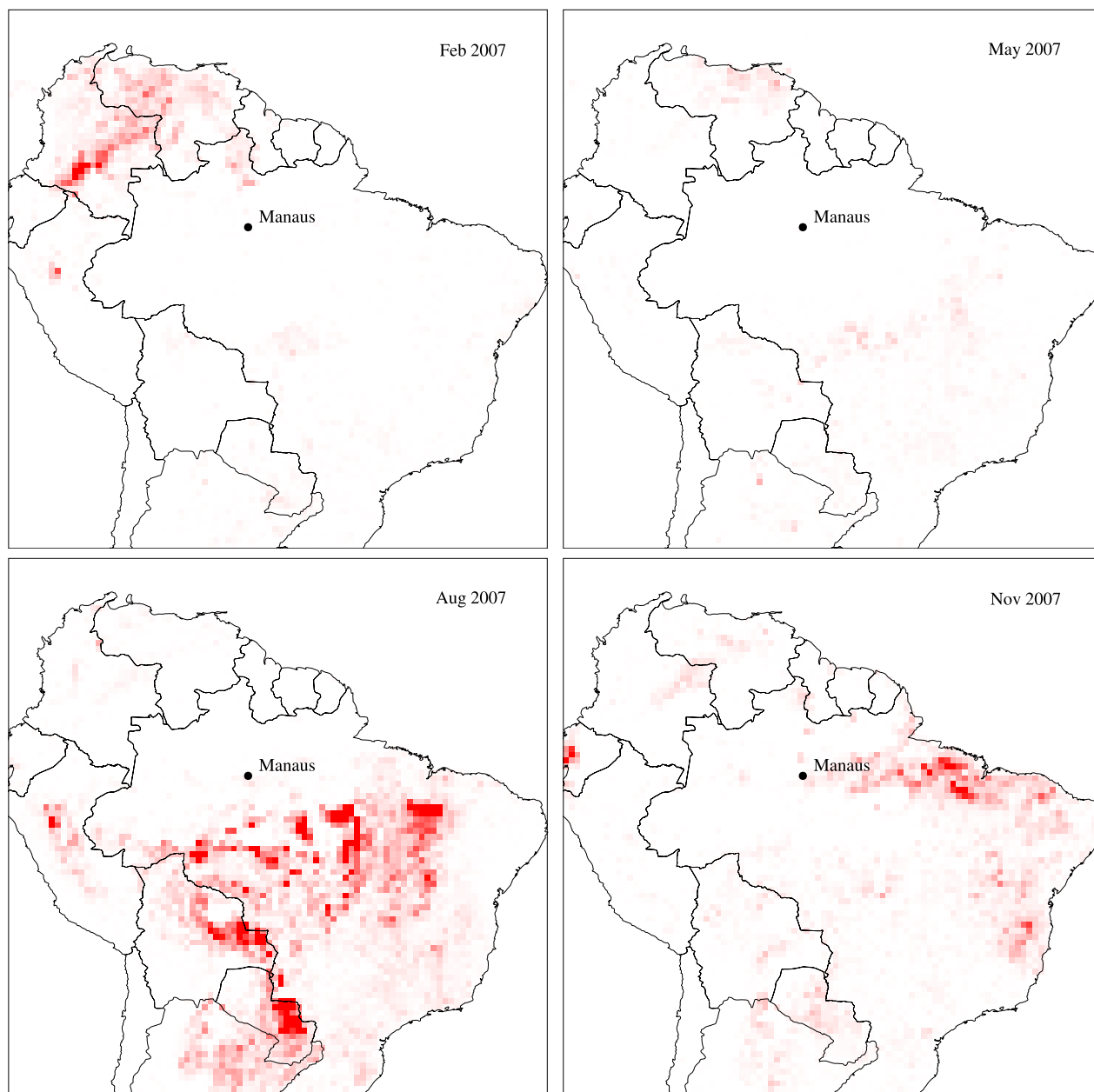


Figure 8. Fire locations in February, May, August, and November 2007 based on the Moderate Resolution Imaging Spectroradiometer Collection 5 Active Fire Product [Giglio *et al.*, 2006]. Red (scaled from 0 to 300) is the number of fire pixels, with corrections for cloud cover. Some boxes have more than 300 fire pixels (e.g., the maximum value for the scenes shown is 1144 fire pixels). Prepared by S. T. Martin and C. L. Heald.

NO_x emitted by biomass burning can lead to regional concentrations of these trace gases and ozone that are similar to those found in industrially polluted regions.

[29] For natural conditions (i.e., as defined by the influences in Figure 4), Figure 9 shows that NO_x and O_3 concentrations in the convective boundary layer over Amazonia are low, of the order of 20–200 ppt and 5–20 ppb, respectively. At night, NO_x concentrations within the canopy space can increase by a factor of ~ 10 because the trace gases emitted from the soils get trapped by the nocturnal inversion (Figure 10). At the same time, O_3 concentrations can drop to nearly 0 ppb in the canopy space, both because of reactions

between NO and O_3 and because of deposition to the vegetation surfaces. The enrichment of trace gases in the nocturnal boundary layer, especially in topographic depressions, can persist into the morning hours, thus providing conditions that may be conducive to intense photochemical processes, possibly including the formation of low-volatility products relevant to particle growth.

2.2.1. Biogenic Volatile Organic Compounds

[30] Biogenic volatile organic compounds are emitted from plants during growth, maintenance, decay, and consumption, and average emission rates account for more than 2% of net primary productivity in the Basin and other regions

[Zimmerman *et al.*, 1988]. The Amazon Basin contains on the order of 10^5 plant species, each having unique signatures of BVOC emissions. Estimates of BVOC emissions from the whole of the Amazon Basin represent a challenging but important task. Prior to new studies conducted in the past decade, Amazonian BVOC emission estimates were based on a small set of measurements conducted by Zimmerman *et al.* [1988]. Kuhn *et al.* [2007] and Karl *et al.* [2007] showed that under some circumstances for specific compounds models of biogenic emissions accurately simulated measured BVOC fluxes in the region. The substantial recent progress in understanding Amazonian BVOC emissions and the major remaining uncertainties are described in detail by Kesselmeier *et al.* [2009].

[31] The specific BVOC compounds emitted and their relative rates of emissions vary widely by plant species and environmental conditions. Major BVOCs emitted include isoprene (C_5H_8), monoterpenes (i.e., compounds composed of two isoprene moieties), sesquiterpenes (i.e., three isoprene moieties), ethane, and oxygenated VOCs. BVOC emissions typically increase exponentially with temperature, doubling every 5–15 K depending on the compound. For some compounds such as isoprene, emissions also increase with available sunlight, and light-dependent monoterpene emissions result in a pronounced diel flux and mixing ratio cycle (Figure 11). The controlling mechanisms for emission differ among compounds. Some are emitted immediately following production, and others are stored in plant tissues. Examples of directly emitted compounds include isoprene and some types of monoterpenes and sesquiterpenes. Examples of stored compounds include other types of monoterpenes and sesquiterpenes, some oxygenated terpenes, and some components of plant oils. The factors influencing the emission of stored compounds to the atmosphere are complex, depending on molecular vapor pressure, animal herbivory, and plant phenology, moisture, or stress.

[32] Although tropical forests are the dominant global source of atmospheric BVOCs and the Amazon Basin is a major contributor [Rasmussen and Khalil, 1988], BVOC emissions have been studied more extensively in temperate regions. The high species diversity in the Amazon Basin is coupled with an ecological complexity and a seasonality, however, that is very different from temperate regions, yielding significantly different emission trends with different forest types. For example, because of the consistently high temperatures over the Amazon Basin, BVOC emissions do not exhibit large seasonality there. Isoprene and monoterpene emissions and concentrations are also strongly correlated in the Amazon (Figure 11), in contrast to their anticorrelated behavior in temperate forests. Isoprene concentrations are highest at midday in temperate forests while monoterpene concentrations are highest at night, corresponding to their emission into a shallow boundary layer. Monoterpene release by plants in those forests is dominated by the emission of stored compounds, and the diel monoterpene emission pattern is therefore significantly different from that of isoprene, which favors release during time periods of intense sunlight. The explanation for the different diel

monoterpene emission pattern in the Amazon rain forest is not yet fully known.

[33] Emissions of BVOCs have been incorporated into global chemical transport models, and the contribution of low-volatility BVOC oxidation products to the mass concentration of organic particles has been predicted [Chung and Seinfeld, 2002; Tsigaridis and Kanakidou, 2003; Hoyle *et al.*, 2007; Heald *et al.*, 2008]. Heald *et al.* [2008] estimated that the conversion of South American BVOCs into secondary particle mass contributes 40% of the annual global production of this particle component. Simulated concentrations over the Amazon Basin varied from 0.6 to $3 \mu\text{g m}^{-3}$ and peaked in the dry season, corresponding to decreased wet deposition during that time period.

[34] A significant underestimate by models of ambient organic particle concentrations, as reported for a number of anthropogenically influenced environments, has been attributed to underestimated conversion yields of oxidized VOCs to secondary particle mass [Volkamer *et al.*, 2007]. The models employ laboratory-based yields of a few percent for isoprene and 10%–15% for most terpenes [Chung and Seinfeld, 2002; Henze and Seinfeld, 2006]. These yields, however, may significantly underestimate what occurs over longer time scales in the atmosphere [Ng *et al.*, 2006]. Recent laboratory studies carried out for BVOC concentrations present in the atmosphere have also shown that yields of secondary particle mass are higher than those obtained by the extrapolation of earlier laboratory results carried out at higher BVOC concentrations [Shilling *et al.*, 2008], and this finding could potentially bring models and observations into closer agreement. A mass balance approach, based on analysis of the fate of BVOCs as either (1) oxidation to CO and CO_2 or (2) deposition with the remainder assumed to form particle mass, leads to much higher production estimates of secondary particle mass [Goldstein and Galbally, 2007]. For the Amazon Basin, the underpredictions by chemical transport models of observed concentrations appear significantly less (e.g., 35% reported by Chen *et al.* [2009]) than those reported for anthropogenically influenced regions of the world.

[35] Further refinement of the mass balance approach requires better measurements of BVOC oxidation products and their rates of wet and dry deposition. Studies of plant signaling, defense, and food and flavor chemistry have led to the detection of thousands of individual BVOCs [Hines, 2006], yet only a few of these have been studied for their emission rates, their atmospheric chemistry, and their contribution to secondary particle mass. Given the limitations in knowledge of emissions rates, oxidation pathways, and yields of particle mass, monoterpene, sesquiterpene, and other BVOC emissions are generally lumped into a few categories for both emission modeling and atmospheric chemistry modeling. As analytical techniques have improved in recent decades, a much broader array of highly reactive and oxygenated BVOCs and their oxidation products has been identified [Holzinger *et al.*, 2005; Lee *et al.*, 2006; Surratt *et al.*, 2008]. The hardest to measure of the directly emitted compounds tend to be the most reactive,

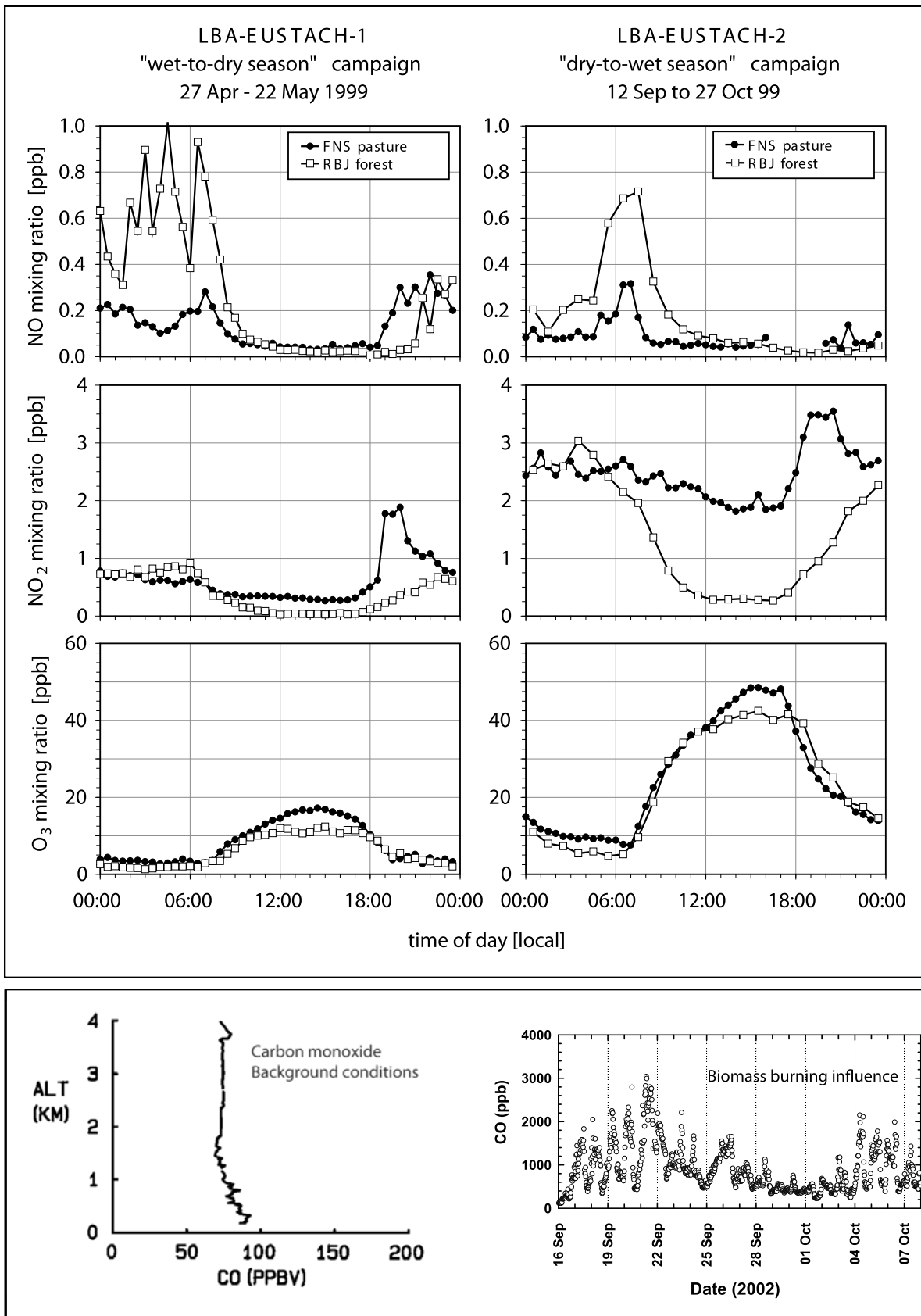


Figure 9

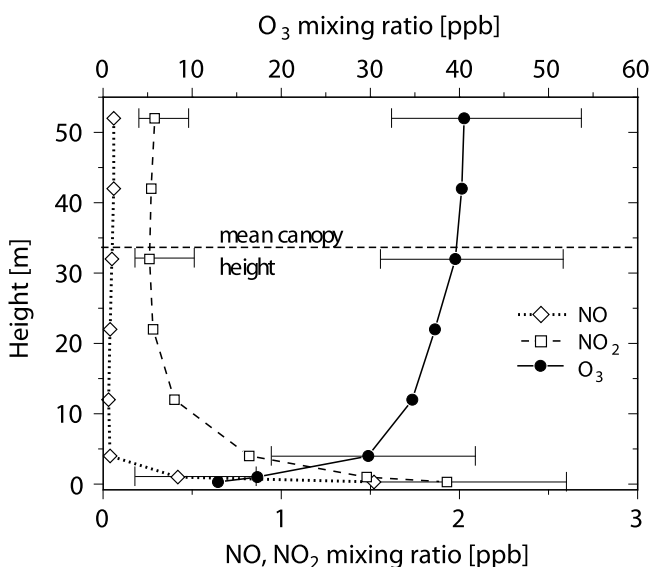


Figure 10. Mean vertical profiles of NO, NO₂, and O₃ in the forest canopy. Profiles represent an average over 43 days in Rondônia at 1400 local time during LBA-EUSTACH-2. Data points of concentrations are medians (0.5-quantile), and corresponding variations are indicated by horizontal bars (0.25-quantile on the left end and 0.75-quantile on the right end). Variation of the NO concentration above 5 m is smaller than symbol size. Adapted from Andreae et al. [2002].

and these same compounds, following multistep oxidation in the atmosphere, often have the highest potential for the production of secondary particle components [Ng et al., 2006]. Sesquiterpenes, which have atmospheric lifetimes of a few minutes or less, are the best known example [Ciccioli et al., 1999].

[36] One approach to estimate BVOC emissions for incorporation in chemical transport models is a bottom-up calculation constrained by leaf, branch, or canopy-scale fluxes [Guenther et al., 1995]. Bottom-up models are based primarily on enclosure measurements that characterize emissions associated with the foliage of an individual plant species. The resulting emission factors are combined with the distributions of plant species to estimate landscape-level emission rates. This approach, however, is difficult to apply in the Amazon Basin because of the high species diversity. Direct BVOC emission measurements are available for <0.2% of the 10⁵ plant species in the Amazon Basin [Harley et al., 2004], and investigations performed to date consist of

only a few measurements per plant species, with analyses that include only a limited subset of all BVOCs. Even so, Harley et al. [2004] demonstrate the utility of this approach for specific Amazonian landscapes. Accurate characterization of the species-dependent emission rates typically requires a large number of measurements over different seasons and locations because of the substantial variability. Representative measurements also require sampling of upper canopy foliage, which is often difficult to access in rain forests. Above-canopy flux measurements, summarized by Kesselmeier et al. [2009] for Amazonia, provide an alternative approach to the parameterization of bottom-up models.

[37] Another approach to estimate BVOC emissions for use in chemical transport models is a top-down calculation driven by satellite observations of the atmospheric distribution of formaldehyde (HCHO). Formaldehyde is a high-yield oxidation product of isoprene, and thus enhancements above the methane background can be used as a proxy for this emission source [Palmer et al., 2003]. The first global top-down study by Shim et al. [2005] provided a global emission estimate that was ~13% higher than the bottom-up estimate of Guenther et al. [1995]. The estimates by Shim et al. [2005], however, were 35% lower for South America than those of Guenther et al. [1995]. Recent top-down investigations have begun to specifically focus on the Amazon Basin. Palmer et al. [2007] examined seasonal variations and found a good correlation between satellite HCHO column measurements and isoprene concentrations measured above the eastern part of the Basin. A related top-down study by Barkley et al. [2008] estimated isoprene emissions that were 35% lower than the bottom-up estimates of Guenther et al. [2006]. This result differed from that of Stavrou et al. [2009], who estimated in a top-down study that isoprene emissions were slightly higher than the bottom-up emission estimate of Guenther et al. [2006]. The bottom-up and top-down models, nevertheless, can be considered to generally agree because each has an uncertainty factor of ~2.

[38] After emission, some fraction of the oxidized BVOCs yields secondary particle mass. Went [1960] first proposed that BVOCs when oxidized create the blue haze observed in the atmosphere above many forested regions. BVOCs are usually oxidized by OH, O₃, or NO₃ in the atmosphere. The OH pathway is particularly important for BVOC oxidation in the tropics given the high light levels and H₂O concentrations. The NO₃ pathway is minor for the usual conditions in the Amazon Basin because there is low O₃ at night and low NO₂ in the day so that the rate of NO₃ formation by

Figure 9. (top) Mean diel NO, NO₂, and O₃ concentrations at the LBA-EUSTACH primary forest site RBJ and at the LBA-EUSTACH pasture site FNS in 1999. Both sites are located in Rondônia (see Figure 1). Measurements were taken 3.5 m above ground at the pasture site and 20 m above the rain forest canopy. Data are presented as 1-h medians over 27 days of the LBA-EUSTACH-1 campaign (mostly natural conditions) and 46 days of the LBA-EUSTACH-2 campaign (strongly influenced by biomass burning). (bottom left) A midday CO profile taken from a flight out of Manaus on 18 July over remote forest in 1985 for natural conditions (Amazon Boundary Layer Experiment: dry season (ABLE-2)). (bottom right) Temporal variation of the CO mixing ratio at 30 min intervals at the FNS surface station during Large-Scale Biosphere-Atmosphere Experiment in Amazonia: Smoke, Aerosols, Clouds, Rainfall, and Climate Field Campaign (LBA-SMOCC) in 2002. Adapted from Sachse et al. [1988], Andreae et al. [2002], and Chand et al. [2006].

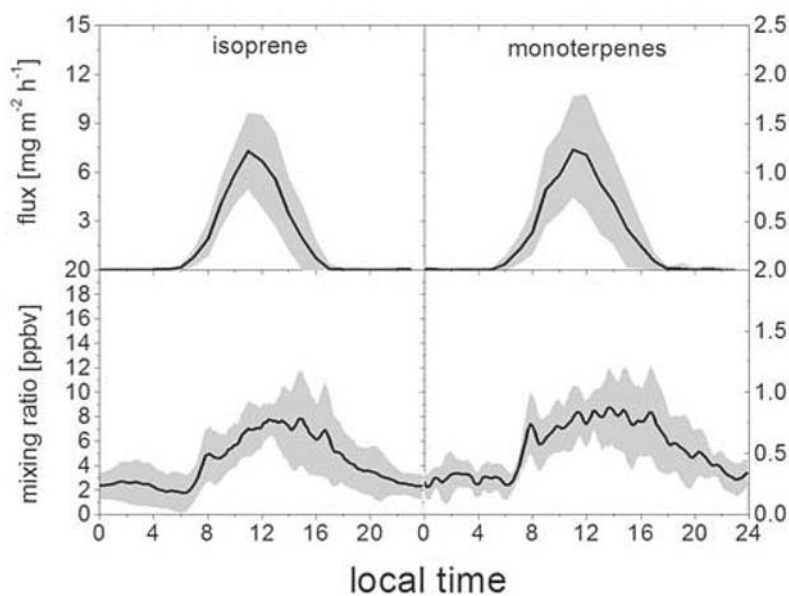


Figure 11. (top) Diel fluxes and (bottom) mixing ratios of isoprene and monoterpenes measured in central Amazonia (tower C14; see Figure 1) between 14 and 29 September 2004. Black lines represent the study mean, and gray area represents the standard deviation. Adapted from *Karl et al.* [2007].

$O_3 + NO_2$ is slow. Overall, knowledge of the composition, the sources, the chemistry, and the role of the secondary organic components of particles in the atmosphere and Earth's climate system is still extremely limited. Even for the well-studied compound isoprene, recent analysis suggests that state-of-the-art atmospheric chemistry models greatly underpredict OH concentrations [*Lelieveld et al.*, 2008], possibly implying important missing chemistry [*Karl et al.*, 2009]. The OH concentrations measured in flights by *Lelieveld et al.* [2008] over the Amazon forest ($5.6 \pm 1.9 \cdot 10^6$ molecules cm^{-3} in the boundary layer and $8.2 \pm 3.0 \cdot 10^6$ molecules cm^{-3} in the free troposphere) were significantly higher than anticipated from model calculations, suggesting an overlooked pathway mediated by organic peroxy radicals for production of OH (Figure 12). Isoprene emissions estimated using inverse modeling can be artificially low for times during which OH is underestimated [*Kuhn et al.*, 2007]. In addition, higher OH estimates change greatly our understanding of photochemistry in the tropics and the rate of transformations of BVOCs into particle mass.

[39] Opportunities for progress in identifying the correct BVOC precursors to secondary particle mass include the following:

[40] 1. For emissions, develop approaches for measuring total nonmethane BVOC and improve approaches for measuring total oxidant reactivity (principally OH and O_3), quantify the major contributions of known compounds to these totals and investigate any residuals, and characterize and understand regional and seasonal variations in canopy-scale emissions using airborne and tower networks of above-canopy flux measurements and satellite observations.

[41] 2. In order to better understand the role of oxidation, develop conceptual approaches to utilize mass balance of

organic material in the atmosphere as a diagnostic tool to test current understanding, to predict compounds that should be in the atmosphere (oxidation products), and to search for them in a systematic way. Explore by experimental or modeling studies the possible formation of low-volatility products relevant to particle production by intense photochemical processes at daybreak, following the enrichment of trace gases in the nocturnal boundary layer.

[42] 3. Develop laboratory, in situ, and remote-sensing techniques to scan the atmosphere for currently unmeasured compounds, to observe sums of compounds by functional classes and compare with measured individual species, to more broadly utilize comprehensive separation technology, and to quantify wet and dry atmospheric deposition of gas-phase and particle-phase organic molecules. Focus on the broad array of semivolatile organic species present in the atmosphere, the majority of which are likely oxidation products of primary BVOC emissions and can potentially condense as secondary organic components of particles.

[43] 4. Formulate models to constrain the chemical influence and fate of products from atmospheric BVOC reactions, to assess the importance of additional organic compounds for atmospheric photochemistry and secondary particle mass, and to represent the full range of BVOCs and their gas-phase and particle-phase products in chemistry and climate simulations.

2.2.2. New-Particle Formation

[44] Near-surface measurements of particle number-diameter distributions and ion number-mobility distributions at many terrestrial sites around the globe, but excluding the Amazon Basin, show that new-particle formation occurs frequently [*Hörrak et al.*, 1994; *Kulmala et al.*, 2004; *Laakso et al.*, 2007; *Iida et al.*, 2008]. Evidence for these

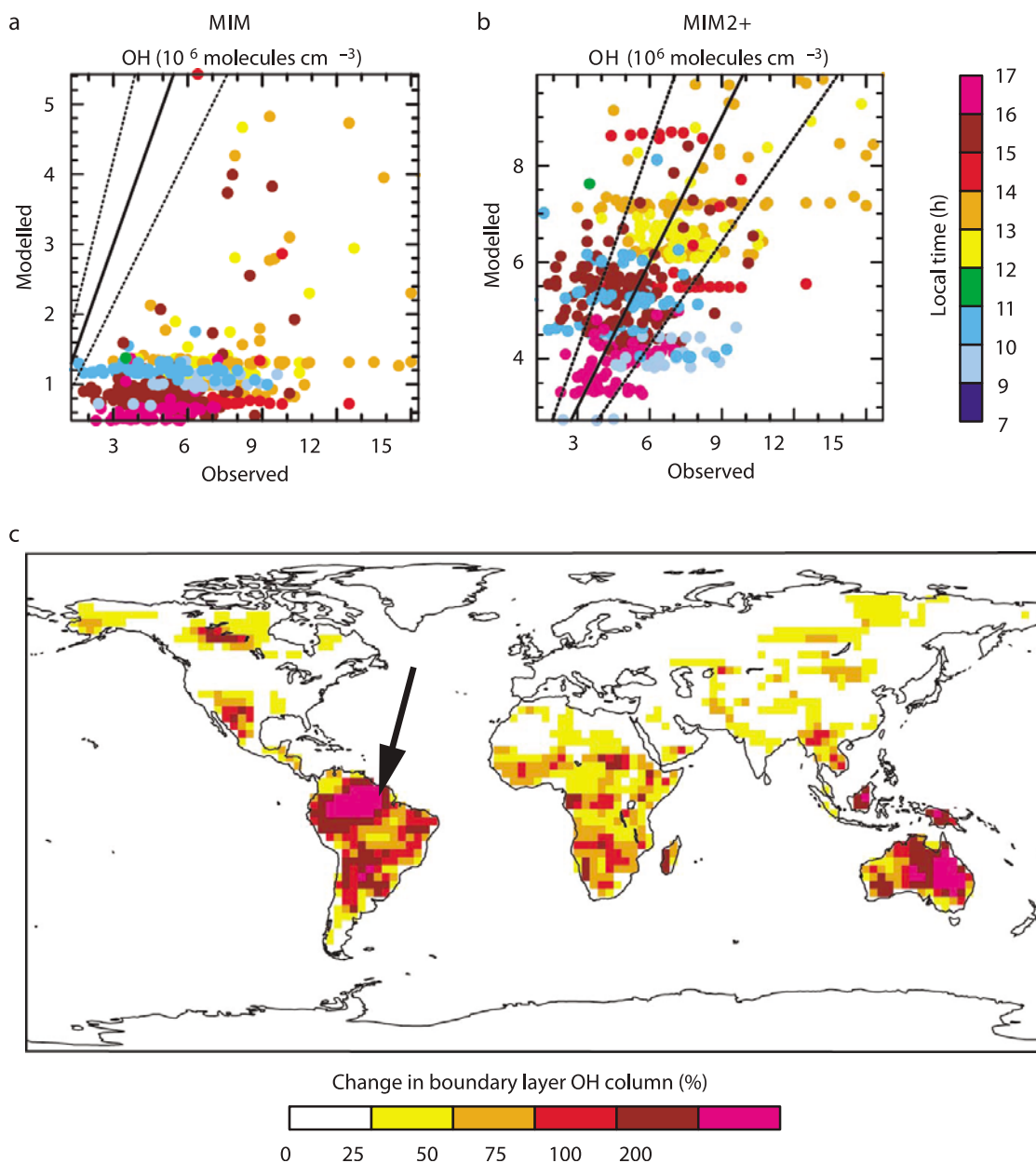


Figure 12. Scatterplots between the amount of OH observed from aircraft and that modeled for the boundary layer over Suriname in October 2005. (a) The standard model. (b) An updated model including the role of organic peroxy radicals. The solid lines indicate ideal agreement, and the dashed lines indicate the $\pm 40\%$ range based on the measurement accuracy. (c) Percentage difference in the annual mean OH, as calculated using the updated model compared to the standard model (the arrow indicates the location of Suriname). The aircraft measurements were performed in October 2005 over the pristine forests of Suriname, Guyana, and Guyane (French Guiana). Adapted by permission from Macmillan Publishers Ltd [Lelieveld *et al.*, 2008], copyright 2008.

events (outside the Amazon) is the appearance of neutral and electrically charged nanoparticles at diameters well below 10 nm and their subsequent growth to larger diameters. Measurements carried out in parallel at sites located distances of several hundred kilometers apart show that the events are often regional [Stanier *et al.*, 2004; Vana *et al.*, 2004; Komppula *et al.*, 2006]. In contrast, measurements in the Amazon Basin provide little evidence for near-surface regional-scale production of new particles [Zhou *et al.*,

2002; Krejci *et al.*, 2003, 2005; Rissler *et al.*, 2004, 2006]. Whereas in other continental locations 3-nm particles are regularly observed at near-surface measurement sites and also seen to grow into the Aitken mode above 30 nm, in the Amazon Basin the smallest particles typically have sizes of 10–20 nm, and continuous growth to larger diameters is rarely observed. Growth rates for Amazonia under pristine conditions have been reported as 5 nm h^{-1} in one location for a limited set of measurements [Zhou *et al.*, 2002], implying

that the observed ultrafine particles nucleated 2–4 h prior to the observations. Further observations are needed, however, to define the possible variability of growth rates in the Amazon Basin.

[45] A constraint on an observable new-particle mode is that freshly nucleated particles must grow more quickly than they are scavenged by coagulation with preexisting larger particles. The likelihood of satisfying this condition increases both with faster growth rates of the nucleated particles and with lower concentrations of preexisting particles [McMurry and Friedlander, 1979; Kerminen and Kulmala, 2002; McMurry et al., 2005; Lehtinen et al., 2007]. The constraint on an observable new-particle mode is typically satisfied for growth rates of 1–10 nm h⁻¹. The condensation of sulfuric acid vapor typically accounts for only a fraction of this growth [Fiedler et al., 2005; Stolzenburg et al., 2005], with most of the balance due to condensation of VOC oxidation products [Smith et al., 2008]. Under the assumption that the growth rate of 5 nm h⁻¹ reported by Zhou et al. [2002] can be broadly extrapolated to the Amazon Basin (i.e., both geographically and seasonally), the discussed constraint should be satisfied, and nucleation events should therefore be observable as new-particle modes. Their absence in observations therefore suggests that near-surface nucleation is not widespread.

[46] Some evidence for the Amazon Basin shows that nucleation occurs at high altitudes and that the entrainment of these particles to the near-surface layer explains surface observations of ultrafine particles. Aircraft measurements above Suriname in northern Amazonia observed enhanced ultrafine number concentrations at 2–4 km in regions of cloud outflow (Figure 13), suggesting nucleation [Krejci et al., 2003, 2005], which is in broad agreement with observations of cloud outflow from other locations worldwide [Perry and Hobbs, 1994; Clarke et al., 1998, 1999]. For comparison, measurements of vertical profiles with a tethered balloon in Melpitz, Germany, showed that nucleation first occurred aloft in the residual layer prior to breakup of the nocturnal inversion and then continued in the mixed layer during and after breakup, all in the absence of clouds [Stratmann et al., 2003]. Over the boreal forests of Finland, a similar measurement program showed that nucleation occurred within the boundary layer but not aloft [Laakso et al., 2007]. In the Amazon Basin, nocturnal events of apparent nucleation, which cannot be explained by outflow from deep convective clouds, have also been observed intermittently for short periods [Rissler et al., 2004, 2006]. Diel patterns of intermittent nucleation were similar for a wide variety of conditions, including periods of intensive biomass burning as well as natural conditions. Nucleation-mode particle concentrations were highest at sunrise and sunset, with average concentrations exceeding 1000 cm⁻³ (Figure 14). Similar nocturnal events were observed in an Australian eucalypt forest [Suni et al., 2008].

[47] The implications of these observations for the Amazon Basin compared to those worldwide are that several different chemical processes may be capable of separately

inducing nucleation and growth and, further, that these different processes may occur in different regions of the atmosphere. Further research, however, may yet succeed in unifying these presently disparate observations by using a mechanistic approach to the problem. A comprehensive analysis of particle formation events recorded at continental locations around the world shows that the nucleation rates J , which quantify the rates at which stable molecular clusters are produced, satisfy the following empirical expression [Riipinen et al., 2007; Kuang et al., 2008]: $J = k[\text{H}_2\text{SO}_4]^p$ for $1 < p < 2$, where $[\text{H}_2\text{SO}_4]$ is the sulfuric acid vapor concentration and k is a kinetic prefactor that varies from location to location. Mechanisms responsible for the variability in k are not yet understood. One hypothesis is that k accounts for the concentrations of species that conucleate with sulfuric acid. Laaksonen et al. [2008] have proposed that BVOC oxidation products may be important conucleating species over forested regions.

[48] Of critical importance for the application of this nucleation equation to the near-surface layer over the Amazon Basin are the weak sulfur sources within the Basin, which lead to sulfur dioxide concentrations of typically 20–30 ppt [Andreae and Andreae, 1988; Andreae et al., 1990a]. This value is more than an order of magnitude lower than the values commonly found under remote conditions over the continents of the Northern Hemisphere. This concentration of SO₂ plausibly implies correspondingly low gas-phase H₂SO₄ concentrations, although no direct observations have ever been made in the Amazon Basin to provide confirmation. Simulated peak daytime near-surface H₂SO₄ concentrations are $< 5 \times 10^5 \text{ cm}^{-3}$ (0.019 ppt) [Spracklen et al., 2005]. By the above nucleation equation, this modeled H₂SO₄ concentration is too low to result in near-surface particle formation because preexisting particles should scavenge any incipient molecular clusters before they grow to new particles [Spracklen et al., 2006]. An alternative mechanism to the H₂SO₄ pathway, namely, ion-mediated nucleation, is also modeled as an unimportant source of nuclei over the Basin [Yu et al., 2008].

[49] In comparison to the absence of predicted new particle formation in the near-surface region of the Basin, models predict that new-particle formation upwind or aloft, in particular, within the upper troposphere followed by growth and entrainment into the near-surface layer, contributes significantly to the Amazonian particle number concentrations, especially during the wet season [Spracklen et al., 2005]. Although sufficient for nucleation at higher altitudes, modeled H₂SO₄ concentrations are insufficient to explain the subsequent rate of particle growth observed in convective outflow over the Basin, suggesting that other gas-phase species such as the oxidation products of BVOCs may have a role [Ekman et al., 2008].

[50] Opportunities for progress to better constrain and quantify mechanisms of new-particle formation over the Amazon include the following:

[51] 1. Perform simultaneous observations of sulfuric acid vapor concentrations, particle nucleation rates, and par-

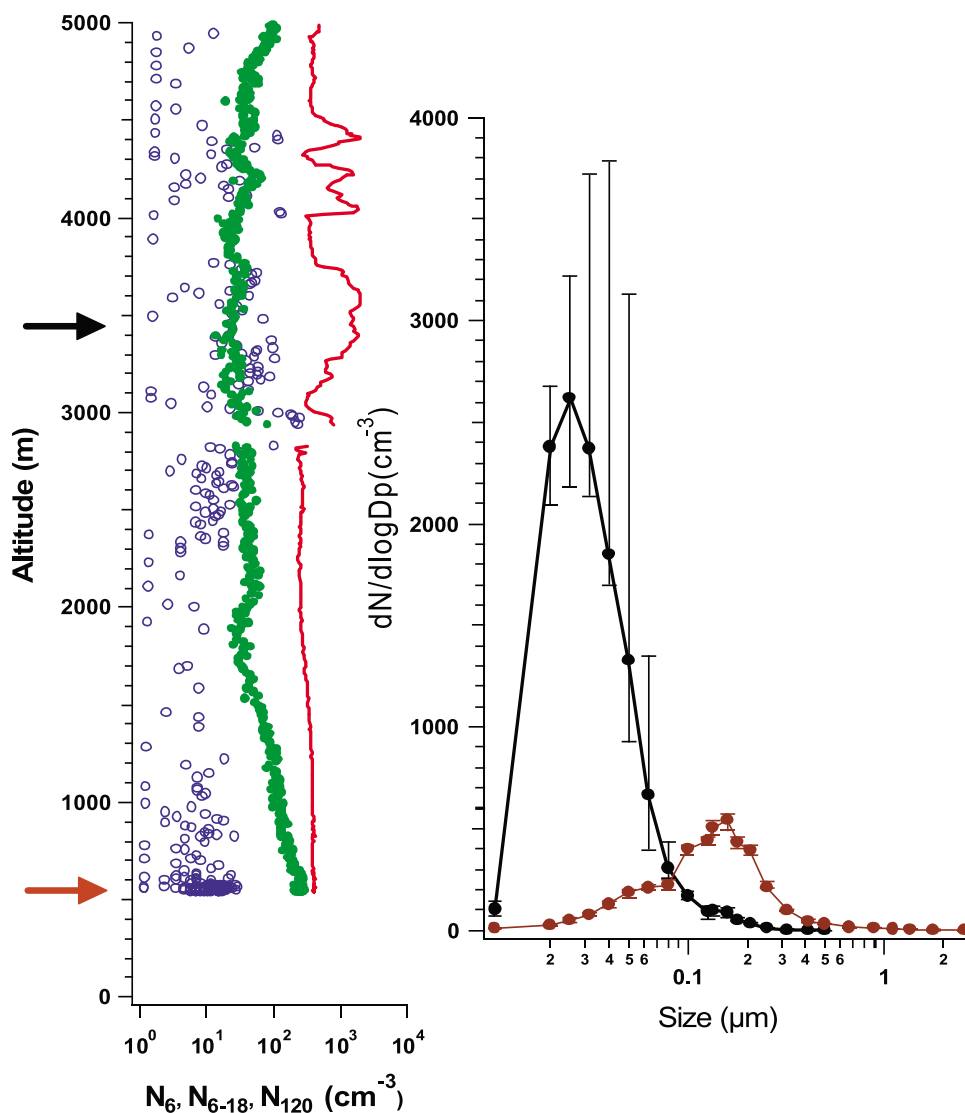


Figure 13. Particle number-diameter distributions observed over southern Suriname. N_6 , N_{18} , and N_{120} represent the number concentration of particles larger than 6, 18, and 120 nm, respectively. The difference N_{6-18} , equal to $N_6 - N_{18}$, quantifies the concentration of nucleation-mode particles. The arrows on the left mark the altitude where the average number-diameter distributions on the right were measured. The error bars on the average distributions represent lower and upper quartiles. Adapted from *Krejci et al.* [2005].

particle number-diameter distributions over the Amazon Basin. Establish whether the apparent absence of surface-level nucleation in the Basin is consistent with our understanding of the atmospheric conditions that lead to new-particle formation in other locations [*McMurry et al.*, 2005].

[52] 2. Measure the growth rates of ultrafine particles over the Basin and evaluate if the contributions by sulfuric acid and known BVOC oxidation products are sufficient to explain the observed rates.

[53] 3. Deploy an air-ion spectrometer [*Hörrak et al.*, 1994; *Mirme et al.*, 2007] in the Basin to give information on very small particles (diameter < 3 nm) and the earliest steps of new-particle formation.

[54] 4. Develop models to assess the contributions of different nucleation mechanisms to aerosol particles in the Basin. Evaluate candidate mechanisms by comparisons of

model predictions made using these mechanisms against past and newly available observations.

3. PROPERTIES

3.1. Mass Concentration

[55] The mass concentrations of particles in the Amazon Basin vary strongly with season and location, modulated to the largest extent by the influence of in-Basin biomass burning with other important influences by the episodic long-range transport of African dust and biomass burning [*Artaxo et al.*, 2002; *Guyon et al.*, 2003b]. Figure 15 shows the time series of PM_{10} concentrations for central Amazonia (Balbina), eastern Amazonia (Santarem), and southern Amazonia (Alta Floresta). These locations are influenced seasonally in varying degrees by in-Basin biomass-burning

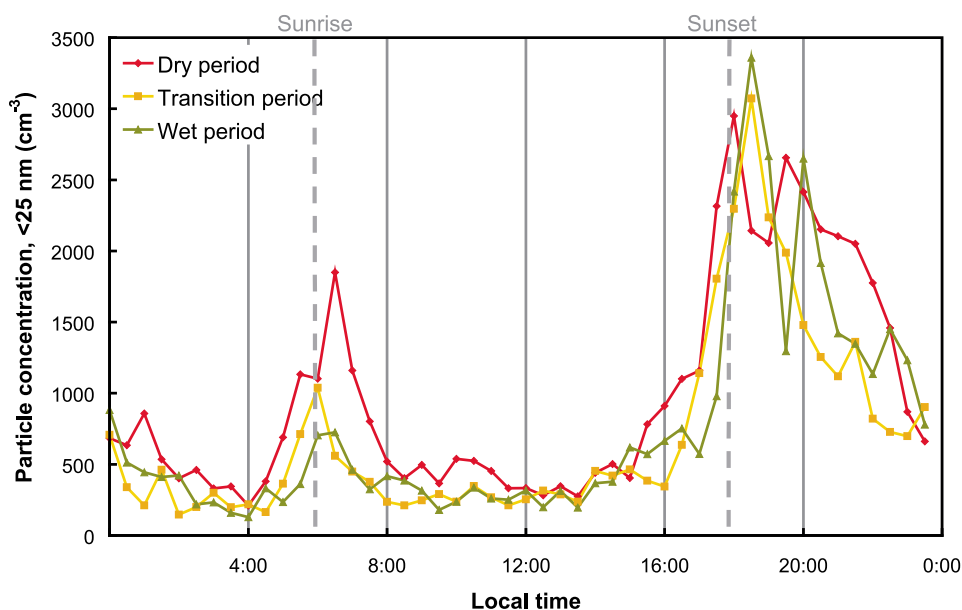


Figure 14. Diel variation in the number concentration of nucleation-mode (namely, <25 nm) particles. Shown are averages for periods of LBA-SMOCC that were weakly, moderately, and strongly influenced by biomass burning. Adapted from *Rissler et al.* [2006].

emissions. In Alta Floresta (August 1992 to February 2005), two different prevailing regimes of mass concentration occur. In the wet season, in the absence of biomass burning, the PM_{10} concentration is $9\text{--}12\ \mu\text{g m}^{-3}$, with a fine fraction of $2\text{--}3\ \mu\text{g m}^{-3}$. In the dry season, the PM_{10} concentration approaches $300\text{--}600\ \mu\text{g m}^{-3}$, producing an optical thickness of more than 4 at 500 nm [*Schafer et al.*, 2008].

[56] For comparison, in central Amazonia where the influence of biomass burning is less, the mass concentration is low even in the dry season (Figure 15, Balbina), with an average PM_{10} concentration of $11\ \mu\text{g m}^{-3}$ (October 1998 to February 2005). The typical concentration of fine particles increases from $2\ \mu\text{g m}^{-3}$ in the wet season to $4\ \mu\text{g m}^{-3}$ in the dry season. The corresponding fine-mode black-carbon-equivalent (BC_e) (see section 3.3.1) mass concentration ranges from 100 to $150\ \text{ng m}^{-3}$ during the wet season and from 600 to $800\ \text{ng m}^{-3}$ during the dry season. Abrupt pulses of relatively high mass concentration can occur in both the fine and coarse fractions (Figure 15, Balbina), and these pulses are explained by African dust outflow that reaches the observation site.

[57] The influence of biomass burning on mass concentration in Santarem in eastern Amazonia is intermediate compared to Balbina and Alta Floresta. From March 2000 to January 2005 in Santarem, the PM_{10} concentration increases from $\sim 10\ \mu\text{g m}^{-3}$ in the wet season to $40\ \mu\text{g m}^{-3}$ in the dry season. The fine fraction, typically as low as $2\ \mu\text{g m}^{-3}$ in the wet season, reaches $20\text{--}30\ \mu\text{g m}^{-3}$ in the dry season. During the wet season, the ratio of the fine-to-coarse fraction is lower for Santarem than Balbina, possibly suggesting an increased relative importance of out-of-Basin coarse-mode particles over eastern Amazonia. Figure 15 also shows that the influence of biomass burning is strongest in the fourth

quarter of each year at Santarem, whereas it is strongest in the third quarter at Alta Floresta. These seasonal patterns match those of vegetation fires in eastern and southern Amazonia, as is apparent in Figure 8.

[58] Figure 16 (top) shows that the particle mass-diameter distribution is dominated for natural conditions by coarse-mode particles, corresponding to primary biological particles possibly coated by secondary material. These data were obtained by gravimetric analysis of the stages of a multi-orifice uniform deposit impactor (MOUDI) during March and April 2008 as part of AMAZE-08 in central Amazonia (S. T. Martin et al., Amazonian Aerosol Characterization Experiment 2008 (AMAZE-08), manuscript in preparation, 2010). The selected data correspond to time periods during which the influence of sources outside of the Amazon Basin was weak [*Chen et al.*, 2009]. *Artaxo and Hansson* [1995] applied principal component analysis to the elemental composition and the mass concentrations recorded on five stages of a cascade impactor for various levels within the canopy and found that the concentrations of potassium and phosphorus, indicative of primary particles, were prominent in the coarse fraction, especially during the night.

[59] The mass-diameter distribution shifts from the coarse to the fine fraction during times of strong influence by in-Basin biomass burning. Figure 16 (bottom) shows MOUDI measurements recorded during Large-Scale Biosphere-Atmosphere Experiment in Amazonia: Smoke, Aerosols, Clouds, Rainfall, and Climate Field Campaign (LBA-SMOCC) in southwestern Amazonia (Rondônia). The total particle mass concentration was $154\ \mu\text{g m}^{-3}$, emphasizing the overwhelming influence of biomass-burning particles compared to any other types at the time of sampling. Figure 16 (bottom) shows that the mass-diameter distribution is

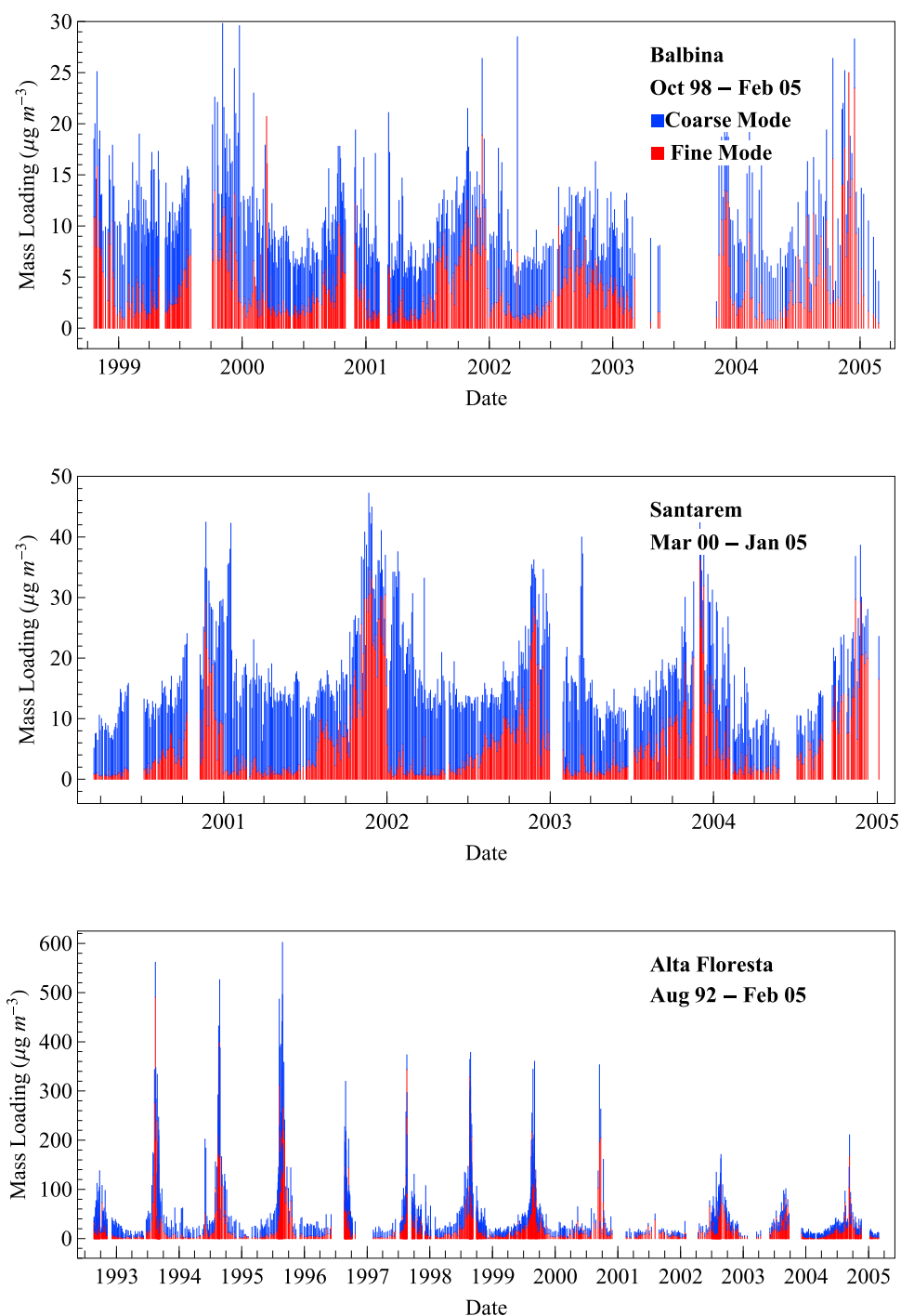


Figure 15. Time series of particle mass concentrations in Balbina (central Amazonia), Santarem (eastern Amazonia), and Alta Floresta (southern Amazonia). Data are shown as stack bar plots of fine (red, $< 2 \mu\text{m}$) and coarse (blue, $2\text{--}10 \mu\text{m}$) fractions. The PM_{10} concentration is the sum of the two. The measurement protocol follows that of the U.S. Environmental Protection Agency for weighing filters, and some water was therefore also included in the mass measurement. The relative mass of water, however, was small because of the low hygroscopic growth factors of Amazonian aerosol particles. Adapted and updated from Artaxo et al. [2002].

heavily loaded in the accumulation mode, with a mass median aerodynamic diameter between 0.33 and $0.56 \mu\text{m}$. Also apparent, however, is that the mass concentrations in the coarse fraction are similar in Figures 16 (top) and 16 (bottom). The implication is that the mass concentration

of coarse-mode natural biogenic particles (possibly coated by secondary materials) is similar in the dry and wet seasons as well as for different locations within the Amazon Basin.

[60] Opportunities for progress to identify and quantify the processes that control mass concentrations and mass-

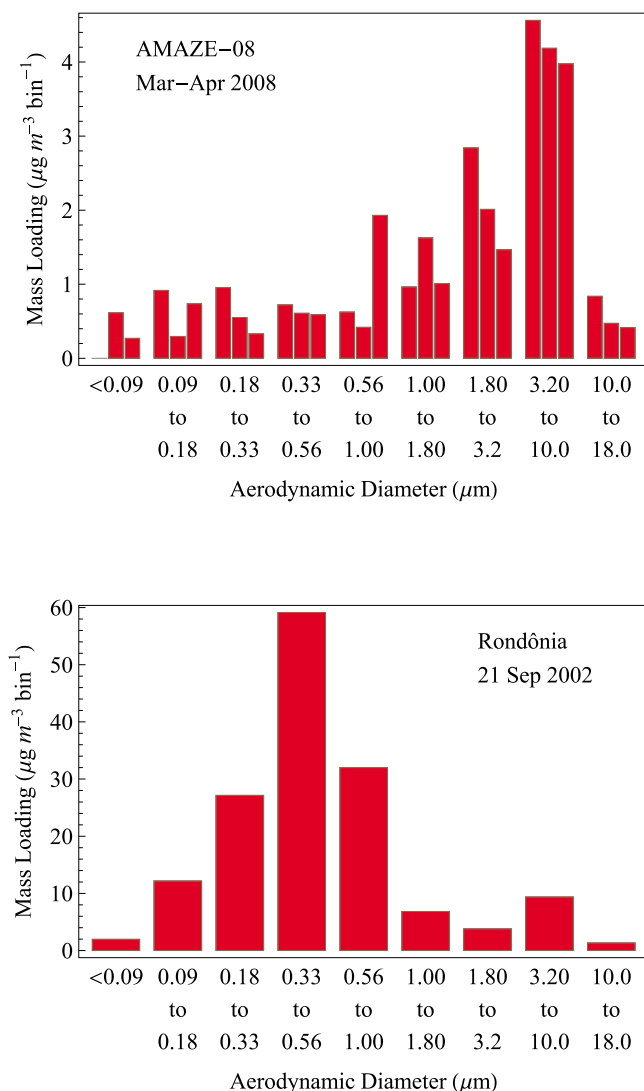


Figure 16. Particle mass-diameter distributions from gravimetric analysis of multiorifice uniform deposit impactor stages. (top) Examples of three distributions collected during Amazonian Aerosol Characterization Experiment 2008 (AMAZE-08) at times when natural conditions prevailed. Each distribution represents 1 week of data (10–16 March 2008, 22–31 March 2008, and 9–17 April 2008). (bottom) Representative distribution collected during LBA-SMOCC during time periods strongly influenced by biomass burning. A bin label such as “3.2 to 10 μm ” represents a mass filter having 50% cut points at 3.2 and 10 μm . Data are from P. Artaxo (personal communication, 2009).

diameter distributions of particles over the Amazon Basin include the following:

[61] 1. Measurements show that long-range transport aerosol particles from Africa and the Atlantic Ocean are present within the Amazon Basin and at times dominate observations, yet the concentrations and distribution of these out-of-Basin particles are not fully understood, especially in relation to the extensive cloud development and opportunities for wet scavenging within the Basin. Long-term measurements at a minimum of three locations across the

Basin along an east-west transect are therefore motivated to provide observations of concentration gradients. Such a data set would provide insights into the sources of particles from outside the Basin as well as sinks within the Basin, thereby placing new constraints on the processes affecting particle lifetime.

[62] 2. Systematic measurements of the size-segregated composition of aerosol particles are scarce in the Amazon Basin. Such measurements are particularly needed in the remote and seasonally pristine areas in the western half of the Basin.

3.2. Number-Diameter Distribution

[63] At times of prevailing natural conditions, particle number concentrations in the mixed layer of the Basin are 200–400 cm^{-3} , lower by 10–100 times than continental and urban concentrations elsewhere in the world having high anthropogenic influence. The baseline, low number concentrations in the Basin are derived from biogenic sources that have low variability and low concentration throughout the year, and strong deviations from this norm that are episodically observed at near-surface stations must be explained by either local pollution or long-range transport from distant sources [Pauliquevis *et al.*, 2007].

[64] The particle number-diameter distributions have been measured in the Basin by techniques such as mobility particle sizing (diameter range of 3–850 nm) [e.g., Zhou *et al.*, 2002], optical particle counting (diameters of 100 nm and larger) [e.g., Krejci *et al.*, 2003], and aerodynamic particle sizing (diameters of 500 nm and larger) [e.g., Rissler *et al.*, 2006]. A typical number-diameter distribution measured by a differential mobility particle sizer (DMPS) operated during a period of natural conditions is shown in Figure 17a (i.e., for the wet season in central Amazonia during the Cooperative LBA Airborne Regional Experiment (CLAIRE-98)) [Zhou *et al.*, 2002]. Aitken and accumulation modes are apparent at 70 and 150 nm, respectively. The histograms of Aitken, accumulation mode, and particle number concentrations for the observations from CLAIRE-98 are shown in Figures 17b–17d, and the modal statistics are summarized in Table 2. The characteristic diameters of the nucleation, Aitken, and accumulation modes are well confined within diameter ranges, and the modes are clearly separated (Table 2). These results of CLAIRE-98 are comparable to those of CLAIRE-01 [Rissler *et al.*, 2004] at the same site and of AMAZE-08 ~100 km away (Figure 18, top) (E. Swietlicki, unpublished data, 2009).

[65] In Figure 17 a Hoppel gap, which is persistent in Amazonian number-size distributions, separates the Aitken and accumulation modes and differentiates to a large extent those particles that have been subjected to in-cloud processing from those that have not. According to Hoppel *et al.* [1986, 1994], cloud droplet activation on Aitken particles is followed by the uptake and reaction of soluble gases to form low-volatility products, and net diameter growth therefore occurs upon evaporation of the cloud droplets. In the Amazon Basin, the soluble gases that become fixed are

expected mostly to be BVOCs, which then react in the cloud waters to form, at least in part, low-volatility BVOC oxidation products (i.e., SOA particle components) [Blando and Turpin, 2000; Lim et al., 2005; Carlton et al., 2006].

Alternative mechanisms for converting Aitken particles into the accumulation mode, such as out-of-cloud coagulation and condensational growth, are too slow in clean atmospheres such as the green-ocean Amazon.

[66] In fair weather, a diel pattern during CLAIRE-98 in the number concentration of the two modes, specifically that the Aitken number concentration decreased while that of the accumulation mode increased as the day progressed, was linked to the diel variation of the lower atmosphere (Figure 6) [Zhou et al., 2002]. Cloud processing above the boundary layer led to the depletion of Aitken particles and to the growth of accumulation-mode particles in that layer, and this cloud convective layer mixed into the surface layer during the day, strongly influencing observations there. In rainy weather, the behavior was interrupted by strong scavenging of particle number, volume, and mass, followed by quick recovery (due to regional mixing) after rainfall ceased.

[67] Measurements in CLAIRE-98 showed that the diameters of both the Aitken and accumulation modes continuously increased from sunrise to sunset, with few exceptions. The growth of the accumulation-mode particles was attributed to cloud processing, with downward mixing of these larger particles throughout the day. The growth of the Aitken particles ($\sim 5 \text{ nm h}^{-1}$) was plausibly explained by the condensation of low-volatility vapors resulting from the photo-oxidation of BVOCs (i.e., SOA production) (see section 2.2.1). Much of this particle growth was proposed to occur in the boundary layer itself. Rainfall temporarily halted the steady diameter growth of the Aitken particles. This observation suggested a down-mixing of somewhat smaller Aitken particles in association with cold downdrafts.

[68] These near-surface observations in central Amazonia during CLAIRE-98, CLAIRE-01, and AMAZE-08 were made on air masses that had spent several days within the Basin and thus were highly processed (e.g., ecosystem emissions, cloud cycles, and so forth). For contrast, measurements were made during flights over Suriname and French Guyana as part of CLAIRE-98 as air masses first entered the Amazon Basin and to some extent can therefore be considered the initialization conditions for processing within the Amazon Basin [Krejci et al., 2003]. The average number-diameter distribution from the lowest flight level within the well-mixed boundary layer (0.2–1.2 km) is shown in Figure 18 (bottom). The distribution was depleted in Aitken particles compared to the near-surface measurements described for central Amazonia (e.g., CLAIRE-98) (see Figures 18 (top) and 18 (bottom)), offering evidence for the formation of Aitken particles within the Basin.

[69] The particle number-diameter distribution changes greatly at locations in the Basin that are influenced by in-

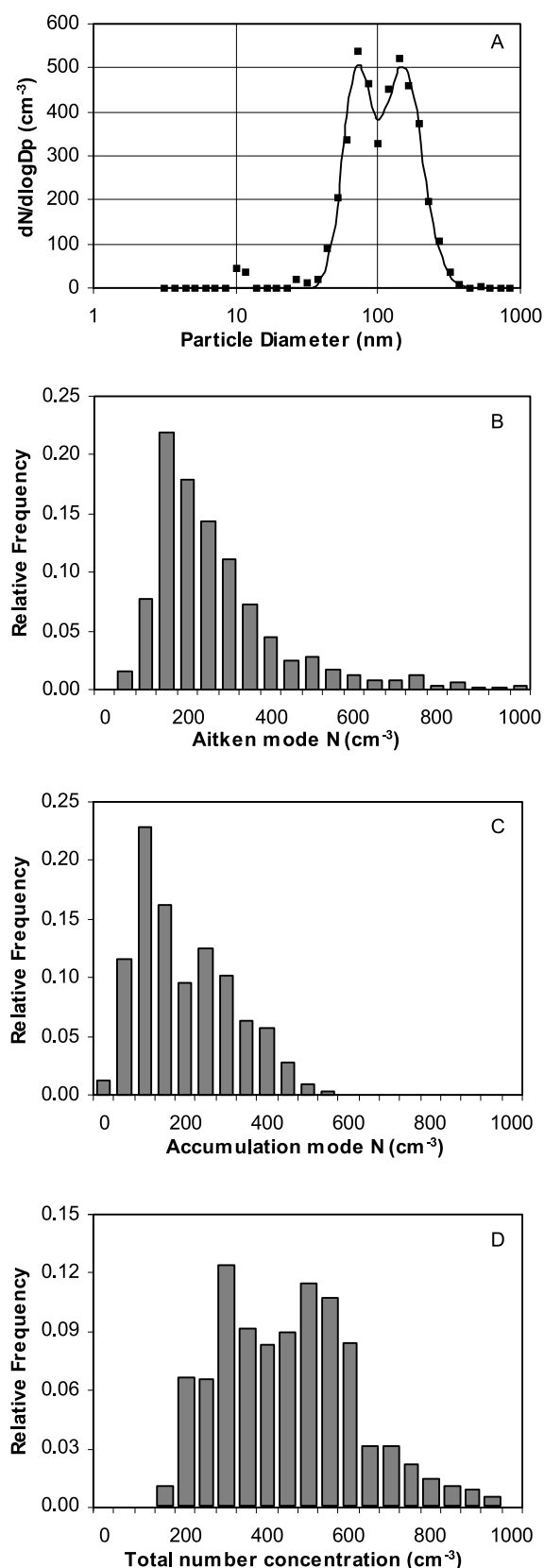


Figure 17. (a) A typical particle number-diameter distribution observed for natural conditions in central Amazonia during the Cooperative LBA Airborne Regional Experiment (CLAIRE-98). Whole-campaign histograms of (b) Aitken, (c) accumulation-mode, and (d) total particle number concentrations. Natural conditions mostly prevailed. Adapted from Zhou et al. [2002].

TABLE 2. Statistics of Particle Number-Diameter Distributions Measured During CLAIRE-98

Mode	Frequency of Occurrence (%)	Number Concentration ($\# \text{ cm}^{-3}$)			Geometric Mean Diameter (nm)	Geometric Standard Deviation
		Mean \pm Standard Deviation	Geometric Mean	Median		
Ultrafine	18	92 \pm 99	55	48	24 \pm 10	1.31 \pm 0.15
Aitken	100	239 \pm 154	200	200	68 \pm 12	1.40 \pm 0.14
Accumulation	100	177 \pm 115	137	146	151 \pm 22	1.40 \pm 0.10

Basin biomass burning [Reid *et al.*, 2005]. The mode of the distribution is typically between 100 and 160 nm. The higher values are found in more aged pollution plumes as a result of coagulation with other particles and the condensation from the gas phase of low-volatility species. Given these high concentrations, when they are present, aerosol particles resulting from in-Basin biomass burning dominate the overall features of the Amazonian aerosol, tending to minimize the impact of other processes such as particles imported from outside of the Basin or the processes of the natural Amazonian biosphere.

[70] DMPS measurements conducted at a ground site during LBA-SMOCC showed a single number median diameter of 135 nm for fresh smoke and average particle number concentrations of $10,500 \text{ cm}^{-3}$ for the diameter range of 30–850 nm [Rissler *et al.*, 2006]. The number concentration of particles in the nucleation mode (i.e., from 3 to 30 nm diameter) was also relatively high, averaging $800\text{--}1000 \text{ cm}^{-3}$, although their presence was intermittent. Airborne measurements of particle number-diameter distributions were also performed [Guyon *et al.*, 2005], and the geometric mean diameters were $110 \pm 15 \text{ nm}$ in 69 plumes within the boundary layer and $139 \pm 17 \text{ nm}$ for 50 smoke plumes detrained above the boundary layer, mostly from nonprecipitating clouds. Biomass-burning particles that enter higher altitudes and escape wet deposition can be exported out of the Basin, affecting particle number and mass concentrations in distant regions of the Southern Hemisphere.

[71] Opportunities for progress in identifying and quantifying the processes that control the number concentration and the number-diameter distribution of particles in the Amazon Basin include the following:

[72] 1. The consistent appearance of the Aitken and accumulation modes in confined diameter windows points to the importance of in-cloud processing, but the source and sink rates of Aitken and accumulation-mode particles, as well as the vertical structure and mixing of particles (including quantifying of cloud convective mixing and associated downdrafts as mechanisms for entraining particles into the boundary layer), must yet be quantified.

[73] 2. Long-term measurements (i.e., years) of size distributions are needed. Needed instruments include an air ion spectrometer, twin scanning mobility particle sizers, and ultraviolet and normal aerodynamic particle sizers, preferably both just over the canopy as well as higher on a tall tower to observe vertical gradients. Measurements should be made of size-resolved particle number fluxes by eddy-covariance techniques, preferably at several altitudes in a high tower. Tethered balloons should be equipped with condensation

particle counters having various smaller cutoff diameters or diffusion batteries to locate altitudes having increased new-particle formation. Ground-based long-term lidar measurements should be employed for vertical profiling. This set of instruments should be applied to closure studies between (1) number and mass, (2) the number-diameter distribution and light scattering, and (3) the number-diameter distribution (including hygroscopic properties) and the concentration of cloud condensation nuclei.

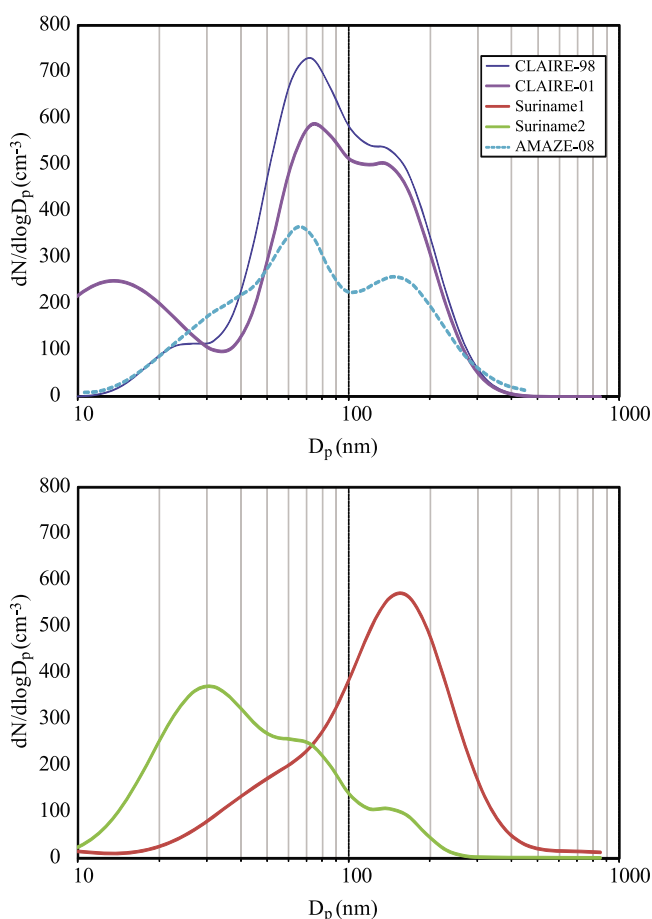


Figure 18. Average particle number-diameter distributions observed in the Amazon Basin for periods during which natural conditions prevailed. (top) CLAIRE-98 (Balbina, 18 days of data) [Zhou *et al.*, 2002], CLAIRE-01 (Balbina, 2 days of data) [Rissler *et al.*, 2004], and AMAZE-08 (approximately equidistant to Manaus and Balbina, 22 February to 12 March 2008) (E. Swietlicki, personal communication, 2009). (bottom) Airborne measurements over Suriname during CLAIRE-98 [Krejci *et al.*, 2003]. Altitude ranges were from (1) 0.2 to 1.2 km and (2) 1.2 to 2.4 km.

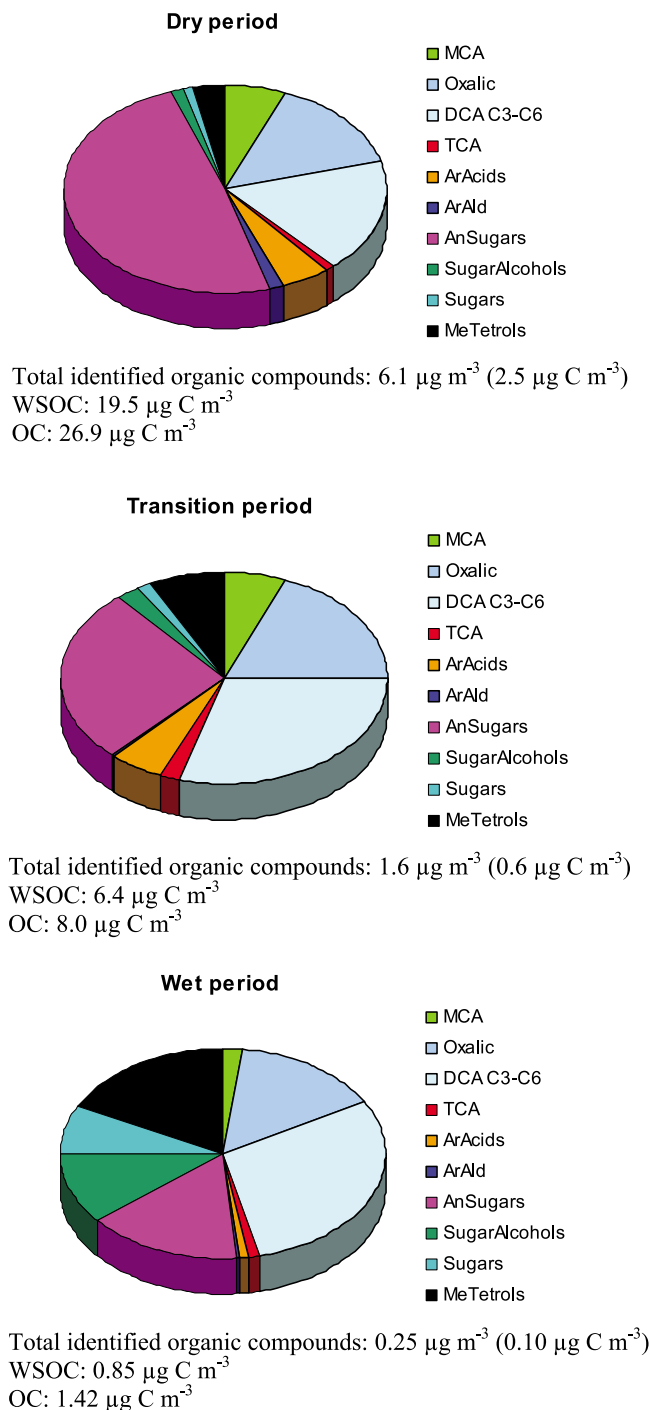


Figure 19. Mean composition of the (identified) water-soluble organic carbon. Fine-mode filter samples collected in the different periods of the LBA-SMOCC campaign were examined by chromatography. Key is as follows: monocarboxylic acids (MCA), oxalic acid, C₃–C₆ dicarboxylic acids (DCA C3–C6), tricarboxylic acids (TCA), aromatic acids (ArAcids), aromatic aldehydes (ArAld), anhydrosugars (AnSugars), sugar alcohols, sugars, 2-methyltetrols (MeTetrols), water-soluble organic carbon (WSOC), and organic carbon (OC). Adapted from Decesari et al. [2006].

3.3. Chemical Composition

[74] Aerosol particles in the Amazon Basin are composed mainly (i.e., 80%–90%) of organic carbon (OC), although African mineral dust and Atlantic sea salt can dominate the mass concentration for short periods [Talbot et al., 1990].

3.3.1. Organic Component

[75] At times of prevailing natural conditions, typical OC concentrations are $1 \mu\text{g C m}^{-3}$ or less in the fine fraction and $1\text{--}3 \mu\text{g C m}^{-3}$ in the coarse fraction [Formenti et al., 2001; Graham et al., 2003a; Guyon et al., 2003b]. A ratio of 1.6–1.7 for OM:OC (i.e., the ratio of organic mass to organic carbon) is estimated from the high-resolution mass spectra of Chen et al. [2009] and in agreement with value of Fuzzi et al. [2007]. Molecular characterization by chromatography has been carried out for the water-soluble organic fraction [Graham et al., 2003b; Claeys et al., 2004]. Results are shown in Figure 19 for particles collected on fine-fraction filters [Decesari et al., 2006]. The identified composition ($0.10 \mu\text{g C m}^{-3}$) is <20% of the water-soluble organic carbon (WSOC) ($0.85 \mu\text{g C m}^{-3}$), indicating the presence of a wide range of other unidentified organic compounds. Dicarboxylic and hydroxyacids are persistently higher during the day compared to the night, consistent with the condensation from the gas phase of BVOC oxidation products [Graham et al., 2003a]. Likewise, the detection of methyltetrols is a signature of isoprene photo-oxidation [Claeys et al., 2004]. BVOC oxidation products may also explain the elevation of fine-fraction OC concentrations by 1.4 ± 0.2 times during the day compared to the night (i.e., greater plant emissions and greater sunlight during the day), although enhanced convective downward mixing of particles from aloft can also play an important role [Graham et al., 2003a, 2003b; Claeys et al., 2004]. The presence of low levels of anhydrosugars (such as levoglucosan, mannosan, and galactosan) even during time periods for which natural conditions prevail demonstrates an influence of biomass burning [Graham et al., 2003a], possibly indicative of the long-range transport of emissions from African fires.

[76] For AMAZE-08, Chen et al. [2009] reported the results of real-time mass spectrometric measurements of submicron particles for a time period of natural conditions. During periods of weak out-of-Basin influence, patterns and identifier peaks in the mass spectra closely resembled those of secondary particle components formed by the oxidation of BVOCs in environmental chambers. Most of the mass concentration of submicron organic particles for the period of AMAZE-08 was attributed to the condensation of BVOC oxidation products as secondary particle components, with a smaller amount to highly oxidized materials representative of humic-like substances (HULIS) that arrived by long-range transport (e.g., aged biomass-burning emissions from Africa).

[77] In regard to the coarse fraction, Graham et al. [2003a] quantified sugars, sugar alcohols, and fatty acids, providing strong evidence for the release of primary biological particles into the forest atmosphere. Trehalose, mannitol, arabitol, and the fatty acids were more prevalent at night, coinciding with

a nocturnal biological activity that increased the release rates of yeasts and other small fungal spores. Glucose, fructose, and sucrose were persistently higher during the day, coinciding with a daytime increase in large fungal spores, fern spores, pollen grains, and, to a lesser extent, plant fragments, as driven by lowered relative humidity and enhanced wind speeds and convective activity during the day. Although mass emissions were reduced at night, coarse-fraction OC concentrations were, nevertheless, elevated at night compared to day by a mean factor of 1.9 ± 0.4 , which was explained by trapping of emitted particles in the nocturnal boundary layer.

[78] The organic compounds constituting the natural particles can be light absorbing. Although the BC_e concentrations are typically $<0.1 \mu\text{g m}^{-3}$ and represent under 5% of the total carbon concentration [Formenti et al., 2001; Graham et al., 2003b], they are nevertheless higher than those of elemental carbon, implying that organic components are contributing to the absorption of light [Guyon et al., 2003a, 2003c]. Because these compounds have a steep increase of light absorption with decreasing wavelength, resulting in a brown color of the filter samples, they have been termed “brown carbon” [Andreae and Gelencser, 2006]. The light-absorbing material is mainly in the coarse fraction and can be explained mostly by chromophores present in primary biological particles and certain BVOC oxidation products.

[79] In addition to natural particles, the composition of biomass-burning particles in the Basin has also been extensively studied [Penner et al., 1991; Andreae, 1993; Falkovich et al., 2005; Decesari et al., 2006; Fuzzi et al., 2007]. The biomass-burning particles, found mostly in the fine size fraction (see section 3.2), are most predominant in southern Amazonia and downwind of it (Figure 5), but even remote areas in northern and central Amazonia are subject to the large-scale transport of biomass-burning emissions [Pauliquevis et al., 2007]. The particles consist of 85%–90% organic carbon [Talbot et al., 1990; Yamasoe et al., 2000; L. L. Soto-García et al., Evaluation of different methods for the determination of BC and OC during biomass burning in the Brazilian Amazon, manuscript in preparation, 2010], of which more than half is water soluble [Graham et al., 2002; Mayol-Bracero et al., 2002; Decesari et al., 2006]. The balance of 10%–15% has typically been operationally defined as black carbon (i.e., apparent elemental carbon). The OC and BC_e concentrations have a diel variability arising both from variations in the thickness of the atmospheric boundary layer and the frequency of fires (Soto-García et al., manuscript in preparation, 2010). Biomass-burning particles also are an important source of water-soluble organic nitrogen, including urea and several amino acids. For example, Mace et al. [2003] reported concentrations of $0.9 \mu\text{g N m}^{-3}$ in the dry season compared to $0.05 \mu\text{g N m}^{-3}$ in the wet season (i.e., a 20-fold difference).

[80] Figure 20 shows the composition of biomass-burning particles in Amazonia, as obtained from a synthesis of data from multiple complementary techniques (e.g., gas chromatography mass spectrometry (GC-MS) and high-pressure liquid chromatography (HPLC)) [Mayol-Bracero et al.,

2002]. The organic compounds were a complex mixture of differing molecular structures, physical properties, and reactivities. Molecular speciation using GC-MS accounted for ~10% of the WSOC. The identified species were mostly pyrolysis products of cellulose, hemicellulose, and lignin [Graham et al., 2002; Zdrahal et al., 2002; Claeys et al., 2004; Trebs et al., 2005; Decesari et al., 2006]. Levoglucosan, a primary vegetation combustion product, was the single most abundant compound identified [Schkolnik et al., 2005; Fuzzi et al., 2007]. It was enriched in samples collected at night compared to those from the day, reflecting the shift from flaming fires in the day to smoldering fires at night [Fuzzi et al., 2007]. The balance of ~90% that eluded analysis by molecular chromatography is expected to be chemical compounds that are predominantly of high molecular weight [Mayol-Bracero et al., 2002; Hoffer et al., 2006]. The HPLC results showed that neutral molecules, monocarboxylic and dicarboxylic acids, and polycarboxylic acids represented ~70% of the WSOC. Decesari et al. [2006] proposed model compounds to reproduce quantitatively the average chemical structure of the WSOC, with the intention that these model compounds can be used as best guess surrogates in microphysical models.

[81] Fuzzi et al. [2006] provide excellent recommendations in a general context on the research needs for improved chemical characterization of organic aerosol particles. For the specific context of the Amazon, high-priority opportunities for increased chemical characterization include the following:

[82] 1. Use sampling techniques such as denuders and real-time measurements that reduce positive and negative sampling artifacts. As necessary, develop new techniques for these purposes. Control sampling with wind direction to facilitate the interpretation of the results. Increase use of size-segregated sampling.

[83] 2. Develop analytical methods to improve chemically resolved mass balance at both the molecular and common property levels. Develop innovative new methods for the analysis of high molecular weight compounds. Determine the composition of light-absorbing material. Develop analytical techniques for airborne measurements having a similar capability as ground-level measurements.

[84] 3. Identify and characterize the molecules and the molecular families constituting the water-insoluble organic fraction. Determine the relative contributions of humic-like substances and BVOC oxidation products to OC mass concentrations during the wet and dry seasons. Characterize and improve the understanding of water-soluble organic nitrogen in biomass-burning particles.

3.3.2. Inorganic Component

[85] Table 3 summarizes the composition of the inorganic component of Amazonian aerosol particles. The inorganic component typically constitutes 10%–20% of the total mass in the fine fraction and less in the coarse fraction, with the balance largely from organic components and at times African dust and African and South American biomass burning.

[86] The two most commonly applied techniques for the study of the inorganic composition have been off-line ion

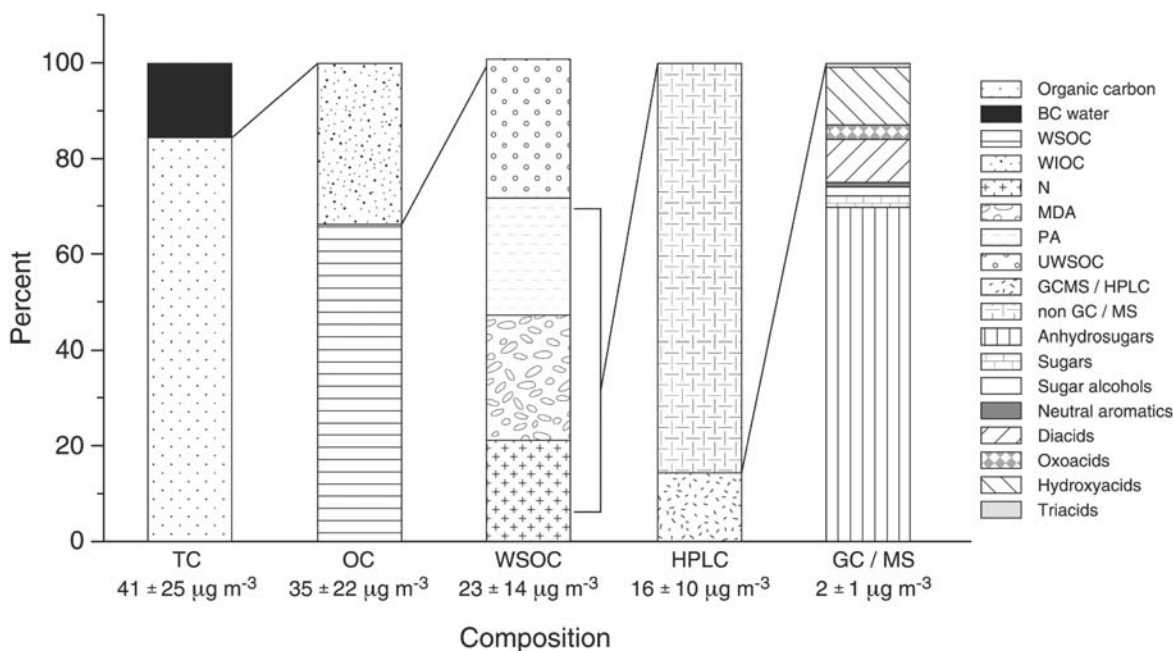


Figure 20. Total carbon apportionment for biomass-burning particles collected during the dry season (LBA-EUSTACH). Total carbon is divided into black carbon and organic carbon, organic carbon is partitioned into water-insoluble and water-soluble fractions, the portion of the water-soluble fraction that is elutable and identifiable by high-pressure liquid chromatography (HPLC) is indicated, the fraction of that eluate that is identifiable by gas chromatography mass spectrometry (GC-MS) is indicated, and the partitioning of that fraction into the chemical species is shown. This final fraction is also represented in Figure 19. “BC water” is based on thermal analysis for black carbon after washing the sample with water. Adapted from *Mayol-Bracero et al.* [2002].

chromatography and proton-induced X-ray emission. Samples were collected for at least half a day up to several days on a single filter to obtain sufficient quantities for analysis. The LBA-SMOCC campaign (Table 1) introduced the use of selective semicontinuous measurements of water-soluble inorganic compounds (i.e., SO_4^{2-} , NO_3^- , Cl^- , and NH_4^+) via a steam-jet aerosol collector and their gaseous precursors (i.e., SO_2 , HNO_3 , HCl , and NH_3) by use of a rotating wet annular denuder [Trebs et al., 2004, 2005]. In AMAZE-08, an Aerodyne aerosol mass spectrometer was employed to quantify nonrefractory sulfate, nitrate, and ammonium every 5 min (Figure 21) [Chen et al., 2009].

[87] Sulfate is the major water-soluble inorganic anion and is primarily distributed in the fine fraction. Sulfate is formed by the reactions of DMS, H_2S , and CS_2 emitted by plants and micro-organisms within the Amazon Basin, and the in-Basin source contributes $\sim 0.05 \mu\text{g m}^{-3}$ to the sulfate mass concentration [Andreae et al., 1990a]. Even for natural conditions, however, sulfate concentrations averaged over several weeks are 3–5 times greater than this in-Basin contribution (Table 3) [Artaxo et al., 1990; Formenti et al., 2001; Artaxo et al., 2002; Fuzzi et al., 2007]. Part of the explanation is that marine DMS is persistently imported from the Atlantic Ocean, representing an integral part of the natural processes of the Basin. Episodic importation of sulfate included as part of African dust and biomass-burning

particles also occurs, and these episodes can increase sulfate concentrations in both the fine and coarse fractions by factors of 2–3 or more [Talbot et al., 1990; Formenti et al., 2001]. In the dry season, the average sulfate concentration increases by a factor of 2–3 (Table 3), even for nominally clean conditions. The increase is attributed in large part to the reduction in wet deposition as well as to the presence of biomass-burning particles diluted throughout the Basin, rather than a change in biogenic gaseous emissions or shifts in the patterns of imported sulfate precursors [Artaxo et al., 1988; Talbot et al., 1988; Artaxo et al., 2002; Graham et al., 2003a; Fuzzi et al., 2007]. At sampling locations and times repeatedly and heavily affected by biomass burning (e.g., in plumes), the average sulfate concentration can increase by a factor of 10 or more.

[88] Nitrate is found predominately in the coarse fraction of natural Amazonian aerosol particles and, compared to sulfate, is a minor component by mass (Table 3) [Talbot et al., 1988, 1990; Graham et al., 2003a; Fuzzi et al., 2007]. The ambient HNO_3 concentrations are too low and the temperatures typical of the Basin are too high to favor elevated concentrations of particle-phase nitrate. Significant nitrate enrichment to the fine fraction occurs, however, for locations affected by in-Basin biomass burning, which can be attributed to increased NO_x emissions followed by oxidation and subsequent condensation of HNO_3 [Trebs et al.,

TABLE 3. Study Average Concentrations Measured for Fine- and Coarse-Mode Inorganic Species During the Wet and Dry Seasons of the Amazon Basin^a

In-Basin Biomass Burning	Out-of-Basin Dust, Marine, African Biomass Burning		SO ₄ ²⁻	NO ₃ ⁻	NH ₄ ⁺	K	P	Ca	Zn	Cl	Na	Al	Si	Ti	Mn	Fe	References ^b
Fine weakly	strongly	weakly	170–260	60–140	31–340	18–34	2.8–7.4	5.0–12	0.25–0.78	0.75–65	26–49	23–62	38–160	1.5–5.0	0.22–0.86	12–37	1–7
		weakly	690	–	–	200	13	120	1.6	180	270	690	1400	46	5.5	340	2
		moderately	510–1000	21	140	80–220	1.9–14	8.0–40	1.2–2.3	1.0–37	51–190	33–130	27–300	2.1–11	0.53–3.6	24–82	8–10
		weakly	38–140	25–36	5.0–13	39–110	8.0–38	11–39	0.60–1.6	8.9–160	54–91	44–92	50–220	4.0–8.6	0.40–1.9	25–64	1–4, 6, 11, and 12
weakly	strongly	moderately	430	–	–	260	37	280	1.9	1100	660	1100	2100	72	10	540	2
		moderately	200–420	–	–	140–270	7.6–87	44–230	2.2–7.4	14–1100	82–2100	100–450	230–970	17–46	2.4–59	190–790	8–10
Fine weakly	strongly	weakly	320–780	14–68	5.0–180	35–160	4.0–6.6	4.0–9.2	0.93–2.6	6.0–8.2	46	8.0–16	25	0.90–1.5	0.19–0.80	6.5–11	13–15
		weakly	1600–3300	1000–1400	860–1600	510–1100	27–33	25–51	4.2–8.3	20–62	–	89–270	120–360	6.0–33	1.9–5.0	28–160	3–6
		strongly	1200–2900	–	–	440–1300	4.7–32	1.5–53	5.3–8.6	12–110	96	100–640	230–540	11–28	1.5–3.8	190–270	8 and 10
Coarse weakly	strongly	weakly	51–200	110–210	27–40	33–100	8.9–32	11–18	0.48–1.8	51–81	–	19–35	33–66	1.4–3.5	0.32–0.60	11–23	13–15
		strongly	180–470	48–520	97–170	88–240	47–68	14–52	2.9–4.3	10–28	–	79–500	90–880	8.0–65	3.9–16	48–460	3, 4, 6, and 12
strongly	strongly	480–3600	–	–	390–1600	17–140	0.86–1300	4.7–11	24–220	–	–	1200–3100	2100–3100	110–140	14–68	1100–1800	8, 10, and 16

^aConcentrations are in ng m⁻³. Time periods of the studies are categorized as weakly, moderately, or strongly affected by biomass burning. Time periods are also categorized for the relative influence of sources from outside the Basin. Entries are rounded to two significant figures for presentation. Fine fraction is defined as aerodynamic diameter under 1 μm for studies 6 and 14; under 2 μm for studies 1–4, 8–10, 13, and 15; and under 2.5 μm for studies 5 and 7. Coarse fraction is defined as aerodynamic diameters from 2 to 10 μm for studies 2–4, 8–10, and 15; of 2–15 μm for studies 1 and 13; under 10 μm for study 12; and >1 μm for study 14. Total suspended particulate matter is used as an approximation of coarse-mode mass for studies 11 and 16. For studies 5 and 7, medians rather than means are listed. The sulfate column includes that measured by ion chromatography as well as that inferred on an equivalent basis from proton-induced X-ray emission measurements of elemental sulfur (i.e., elemental sulfur attributed entirely to sulfate in the analysis).

^bReferences are as follows: 1, *Artaxo et al.* [1990]; 2, *Formenti et al.* [2001]; 3, *Artaxo et al.* [2002]; 4, *Gayon et al.* [2003c]; 5, *Trebs et al.* [2007]; 6, *Fuzzi et al.* [2008]; 7, *Trebs et al.* [2008]; 8, *Echalar et al.* [1998]; 9, *Gerab et al.* [1998]; 10, *Maenhaut et al.* [2002]; 11, *Talbot et al.* [1990]; 12, *Mace et al.* [2003]; 13, *Artaxo et al.* [1988]; 14, *Talbot et al.* [1988]; 15, *Graham et al.* [2003a]; 16, *Artaxo et al.* [1998].

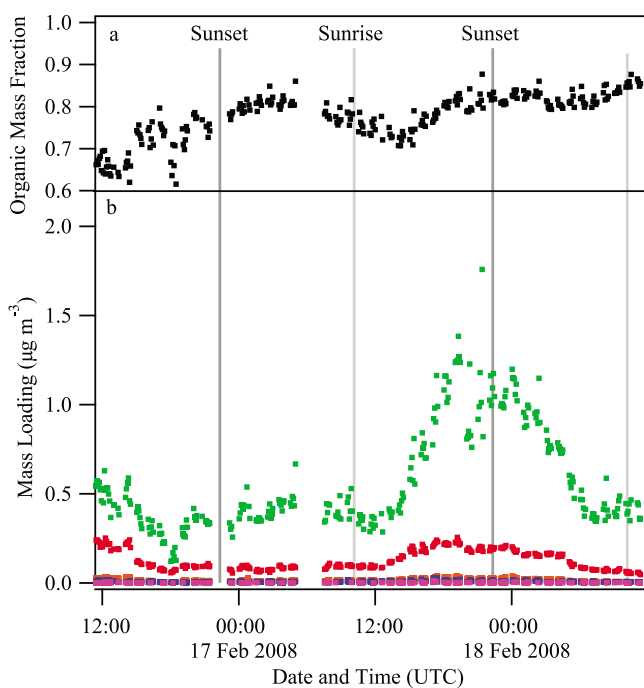


Figure 21. Two-day time series of chemically apportioned, submicron particle mass concentrations measured on 17 and 18 February 2008 in central Amazonia as part of AMAZE-08. Measurements were made using an Aerodyne high-resolution time-of-flight aerosol mass spectrometer during a time period for which natural conditions prevailed. (a) Time series of organic mass fraction. (b) Time series of submicron mass concentrations of organic material (green), sulfate (red), ammonium (orange), nitrate (blue), and chloride (magenta). The measurement window of the AMS is $\sim 60\text{--}600$ nm in vacuum aerodynamic diameter. Data are from S. T. Martin (personal communication, 2009).

2004]. At these times, the nitrate and sulfate concentrations can be nearly equal [Trebs et al., 2005].

[89] Ammonium concentrations in the fine and coarse fractions are typically more than sufficient during both the wet and dry seasons to neutralize the nitrate and sulfate in each size fraction, particularly at night [Talbot et al., 1988; Graham et al., 2003a; Mace et al., 2003; Fuzzi et al., 2007]. Trebs et al. [2005] concluded that organic anions, such as formate, acetate, and oxalate, balance the ammonium in the particles, implying the presence of species such as ammonium oxalate. At locations having a strong influence from biomass burning, ammonium concentrations, along with those of nitrate, are elevated in the fine fraction during the night [Trebs et al., 2004], which can be explained by higher concentrations of HNO_3 and NH_3 in the vicinity of biomass burning and favorable gas-to-particle partitioning for higher relative humidity and cooler temperatures.

[90] The mass concentrations of crustal elements (e.g., Al, Si, Fe, Ti, and Mn) are low in the absence of episodic inputs of African mineral dust (see Table 3) [Talbot et al., 1990; Formenti et al., 2001; Graham et al., 2003a; Trebs et al., 2005]. During episodes, their concentrations are equally

distributed in the fine and coarse fractions, representing a fractionation that is significantly different from the dominance of the coarse mode near source regions. After long transport over the Atlantic Ocean, the coarse mode of these elements is lost preferentially over the fine mode, resulting in the observations seen for the Amazon, including an important tail into the accumulation mode. Wet-season concentrations are several times higher than those during the dry season, which is explained by a shift of the ITCZ that favors the transport to the Amazon Basin of air from the Saharan region (see section 1).

[91] The mass concentrations of P, K, and Zn in the coarse fraction are derived primarily from PBA particles [Artaxo and Hansson, 1995; Guyon et al., 2003c], as evidenced by their greater concentrations underneath the forest canopy compared to above it and by their greater concentrations at night than day (see section 3.3.1). Although the long-range transport of phosphorus from African dust is important for the fertilization of the Basin on long time scales [Mahowald et al., 2005], the incremental concentrations are low and difficult to detect compared to the in-Basin cycling of phosphorus by PBA particle emission and deposition [Formenti et al., 2001]. For natural conditions, K occurs almost exclusively in the coarse fraction, but a fine mode appears at times of biomass burning influence, increasing the mass concentration by a factor of 10 or more in fresh plumes downwind of biomass burning [Echalar et al., 1998; Yamasoe et al., 2000; Artaxo et al., 2002; Maenhaut et al., 2002; Trebs et al., 2005]. Fire is therefore important for the recycling of K throughout the Basin, and its presence in the fine fraction is often used as a marker for the influence of biomass burning.

[92] High priorities for the improved chemical characterization of the inorganic component of Amazonian aerosol particles include the following:

[93] 1. In terms of instrumentation, positive and negative sampling artifacts can be introduced, especially for semi-volatile compounds like ammonium nitrate, by swings in temperature and relative humidity during the long sampling periods required by many techniques when employed at the low particle mass concentrations characteristic of the Amazon Basin [Solomon and Sioutas, 2008; Trebs et al., 2008]. The development of instrumentation that can overcome these obstacles in tropical environments is needed.

[94] 2. Investigations are needed (1) that target the influence of water-soluble organic compounds on the gas particle partitioning of inorganic species, (2) that test for the possible presence of organosulfates and understand how these compounds may be tracers for the sources of particle components, and (3) that focus on connections between inorganic compounds and their physical properties (e.g., their activity as cloud condensation nuclei).

[95] 3. Long-term measurements are needed for a better understanding of seasonal and annual variability, especially the effects of the long-range transport of African dust and biomass-burning emissions. Measurements with improved geographic coverage are needed from the eastern edge of the

Basin into central parts to quantify gradients in African and Atlantic aerosol particles and thereby to understand better their influence.

3.4. Hygroscopicity

[96] The hygroscopic properties of submicron Amazonian aerosol particles have been studied both for the wet season during periods of weak out-of-Basin influence [Zhou et al., 2002] and for the dry season at times of strong in-Basin biomass burning [Rissler et al., 2004, 2006]. Irrespective of season and the air mass type, the hygroscopic diameter growth factor measured at 90% relative humidity (RH) by use of a tandem differential mobility analyzer is typically 1.05–1.35, with few exceptions. This finding differentiates Amazonia from rural sites on other continents, for which “highly hygroscopic” particles having growth factors of 1.7 that approach those of inorganic salts are observed for at least a fraction of the particles [Swietlicki et al., 2008]. The consistently high organic fraction of submicron particles (see section 3.3) can explain the absence of highly hygroscopic particles in the Amazon Basin. As the exception, “highly hygroscopic” particles have, however, been observed in the Amazon Basin at least once [Zhou et al., 2002], plausibly corresponding to the presence of marine particles imported with an Atlantic air mass [Formenti et al., 2001].

[97] For natural conditions, “moderately hygroscopic” particles dominate the submicron particle population in the Amazon Basin [Zhou et al., 2002; Rissler et al., 2004]. Zhou et al. [2002] find that the hygroscopic growth factor increases from 1.17 at 35 nm to 1.32 at 264 nm. Ammonium bisulfate dry-volume fractions ranging from 0.17 at 35 nm to 0.27 at 265 nm, with the balance of the dry-volume fraction corresponding to an insoluble core, can equivalently represent the observed hygroscopic growth. This equivalent representation must not, however, be interpreted as implying that the water-soluble components of the real particles are composed solely of ammonium and bisulfate ions. Water-soluble organic compounds constitute a large fraction of the particle components (see Figures 19 and 20), although many of the substances, such as the larger dicarboxylic acids, fulvic acids [Svenningsson et al., 2006], or HULIS [Ziese et al., 2008], have low water uptake. There is also the important possibility of a difference between the WSOC fraction measured for relatively dilute aqueous solution (i.e., as represented in Figures 19 and 20) and that relevant to the lower water activity of 90% RH at which measurements of hygroscopic growth have been made.

[98] For locations strongly influenced by fresh in-Basin biomass burning, an external mixture of “moderately hygroscopic” (as described above) and “barely hygroscopic” particles is observed [Rissler et al., 2006; Vestin et al., 2007]. “Barely hygroscopic” particles have growth factors from 1.06 at 20 nm to 1.12 at 440 nm, corresponding to inorganic-equivalent dry-volume fractions of ~ 0.07 . The “barely hygroscopic” particles can dominate the number balance of the external mixture by a factor of 5–10, with a larger fraction at smaller particle sizes. Open-air biomass

burning produces particles largely composed of organic components that have a limited propensity for water uptake.

[99] In an advance compared to the inorganic-equivalent representation for hygroscopicity, Mircea et al. [2005] provided a more comprehensive treatment of chemical hygroscopic closure that included a treatment of the organic component. Growth factors observed during LBA-SMOCC (i.e., moderately to strongly influenced by biomass burning) for particles 420 nm and smaller were compared with the predictions of a water-uptake model that incorporated the size-segregated chemical composition. Inorganic components accounted for $\sim 10\%$ of the size-segregated mass concentrations, and the balance was carbonaceous [Fuzzi et al., 2007]. Water-soluble organic compounds constituted 50%–60% of the mass concentrations [Decesari et al., 2006]. The water-uptake model used for the hygroscopic closure simplified the organic composition by choosing nine model compounds derived from functional group analyses and other analytical techniques (see further description in section 3.3). The model, combined with the size-segregated chemical composition, accurately predicted the measured growth factors for an assumption of limited solubility of the organic compounds at 90% RH. Closure could not be obtained for other candidate assumptions, including complete solubility or complete insolubility.

[100] Priorities for progress to better constrain the hygroscopic behavior of Amazonian aerosol particles include the following:

[101] 1. Amazonian biomass-burning particles were extensively characterized during the LBA-SMOCC experiment in 2002. Less is known about natural particles in pristine rain forest environments, essentially based on 18 days of data from Zhou et al. [2002] and two days of data from Rissler et al. [2004], and further measurements are therefore highly motivated with a focus on understanding temporal and geographic variability.

[102] 2. The links between size-segregated chemical composition and hygroscopic behavior, for instance, by implementing new instrumentation such as aerosol mass spectrometers, as well as between hygroscopic growth and radiative properties, for instance, by conducting experiments using humidity-controlled nephelometers, should be pursued.

[103] 3. The presence of PBA particles should be quantified by utilizing the ability of a tandem differential mobility analyzer to determine the hygroscopic growth of individual particles. The intermittent appearance of externally mixed “barely hygroscopic” particles during the wet season of Zhou et al. [2002] (i.e., a frequency of occurrence of 5%–9% in the Aitken mode and 11%–14% in the accumulation mode) might indicate an influence from PBA particle sources, although anthropogenic pollution was not entirely ruled out as having an influence during that study.

3.5. Cloud Condensation Nuclei

[104] The concentrations of cloud condensation nuclei (CCN), in the absence of an influence by in-Basin biomass burning, are 200–300 cm^{-3} for 1% supersaturation [Roberts

et al., 2001, 2002; *Andreae et al.*, 2004; *Gunthe et al.*, 2009]. These CCN concentrations are comparable to the total particle concentration, and they are lower than typically observed for other rural sites worldwide, a finding which is indicative of strong anthropogenic influence for most other continental locations worldwide [*Andreae*, 2007, 2009]. The CCN properties for natural conditions can be described in large part by an effective CCN hygroscopicity parameter κ of 0.15 [*Gunthe et al.*, 2009]. Such particles in the Aitken and accumulation modes are sufficiently hygroscopic that they activate at supersaturations of 0.1–1% [*Zhou et al.*, 2002; *Svenningsson et al.*, 2006]. For comparison, κ is larger by a factor of two for other continental locations and by a factor of four for typical marine particles [*Andreae and Rosenfeld*, 2008]. The lower κ in Amazonia is consistent with the properties of SOA material reported in laboratory studies [*King et al.*, 2007, 2009; *Prenni et al.*, 2007] and with the report of *Chen et al.* [2009] for AMAZE-08 that high proportions of secondary organic matter constitute the components of submicron particles. Supermicron PBA particles present at relatively low number concentrations can also be important under some circumstances by serving as “giant” CCN, which activate at supersaturations below 0.1% because of their large diameters and enhance the collision-coalescence stage of precipitation formation, especially under polluted conditions [*Yin et al.*, 2000]. PBA particles are also an important source of ice nuclei in the Basin [*Prenni et al.*, 2009].

[105] In stark contrast to the low CCN concentrations observed for natural conditions, regions affected by in-Basin biomass burning can have CCN concentrations of $10,000\text{ cm}^{-3}$ or more (Figure 22) [*Roberts et al.*, 2003; *Rissler et al.*, 2004; *Vestin et al.*, 2007]. The contribution to CCN number concentration arises not only from the increase in particle number concentration but also from increases in mode diameter and in water-soluble fraction, both of which further favor CCN activation [*Mayol-Bracero et al.*, 2002; *Decesari et al.*, 2006; *Fuzzi et al.*, 2007]. For the observations of LBA-SMOCC, *Mircea et al.* [2005] show that CCN closure in the supersaturated regime (i.e., >100% RH) is best achieved by assuming complete water solubility of organic species at the high water activities of CCN activation.

[106] The difference in CCN concentrations between natural and anthropogenically influenced conditions has significant consequences on the microphysical properties of clouds, particularly the average droplet diameter, the maximum in-cloud supersaturation, and the precipitation dynamics. Microphysical properties are most susceptible to increasing CCN concentrations for low base concentrations, such as those of natural Amazonia. The input of additional particles from biomass burning greatly alters the pathways of cloud development. When smoke plumes spread over large areas, shallow clouds are inhibited, causing a reduction in cloud cover [*Koren et al.*, 2004, 2005]. *Feingold et al.* [2005] and *Jiang and Feingold* [2006] also suggest that the extinction of radiation by elevated particle concentrations in the middle troposphere reduces the surface heat flux, thereby stabilizing the boundary layer and further reducing

cloud cover. Furthermore, enhanced CCN concentrations that result from biomass burning reduce the cloud droplet diameter below the collision-coalescence threshold, an effect which reduces warm-cloud precipitation [*Andreae et al.*, 2004]. Another potential effect of increased CCN concentrations, which is enhanced cloud albedo as a result of both the smaller droplet diameter and more numerous droplets [*Twomey*, 1977], is small or negligible in the Amazon Basin because the clouds are usually already optically thick [*Platnick and Twomey*, 1994; *Roberts et al.*, 2003].

[107] The range of microphysical regimes observed in the Basin, including blue-ocean clouds, green-ocean clouds, smoky clouds, and pyroclouds, is illustrated by the mass-diameter distributions of liquid water content (Figure 23) [*Andreae et al.*, 2004]. There is a narrowing of the distributions and a slowing of their rate of broadening with increasing height for the progressively more particle-rich regimes from Figures 23a to 23d. For low CCN concentrations (Figures 23a and 23b), the droplet distributions over the ocean and the Amazon Basin grow and broaden in a similar manner as the parcel rises. In contrast, for the very high CCN concentrations of pyroclouds (Figure 23d) that form in the invigorated updrafts of the smoke plume over an active fire, the droplet distribution stops growing once the air parcel rises above a critical altitude (e.g., above 2800 m in Figure 23). The stunted growth is explained by reduced in-cloud supersaturation that inhibits droplet growth and has the consequence of suppressing precipitation. Pyroclouds embedded in a smoky atmosphere also reduce ground heating by blocking sunlight, and they therefore reduce convective vigor and precipitation. Smoky clouds (Figure 23c) represent an intermediate case compared to green-ocean clouds and pyroclouds. In smoky clouds, the altitude for the onset of precipitation, which corresponds to a modal diameter of the liquid water content that is greater than an approximate threshold of $24\ \mu\text{m}$, shifts higher compared to natural conditions. The consequence is that rain either does not occur or occurs from higher altitudes, including more precipitation initiated through the ice phase. Consequently, there is a vertical redistribution of released heat and, in the case of ice, invigorated convection and lightning [*Rosenfeld et al.*, 2008].

[108] The need to understand how the CCN activity of organic particles evolves and the associated implications for direct and indirect radiative forcing have been highlighted in general reviews [*Kanakidou et al.*, 2005; *Fuzzi et al.*, 2006; *McFiggans et al.*, 2006; *Andreae and Rosenfeld*, 2008]. Specific priorities for better understanding and predicting the CCN activity of Amazonian aerosol particles and their effects on climate include the following:

[109] 1. In terms of measurements, data sets of CCN activity in the Basin are restricted in both time and space and can be considered sparse in comparison to the seasonal and spatial heterogeneities of the Amazon Basin. Most measurements in available data sets were carried out during intensive campaigns lasting a short number of weeks at a single location [*Roberts et al.*, 2002; *Rissler et al.*, 2004; *Vestin et al.*, 2007; *Gunthe et al.*, 2009]. Long-term and distributed measurements are therefore highly motivated so

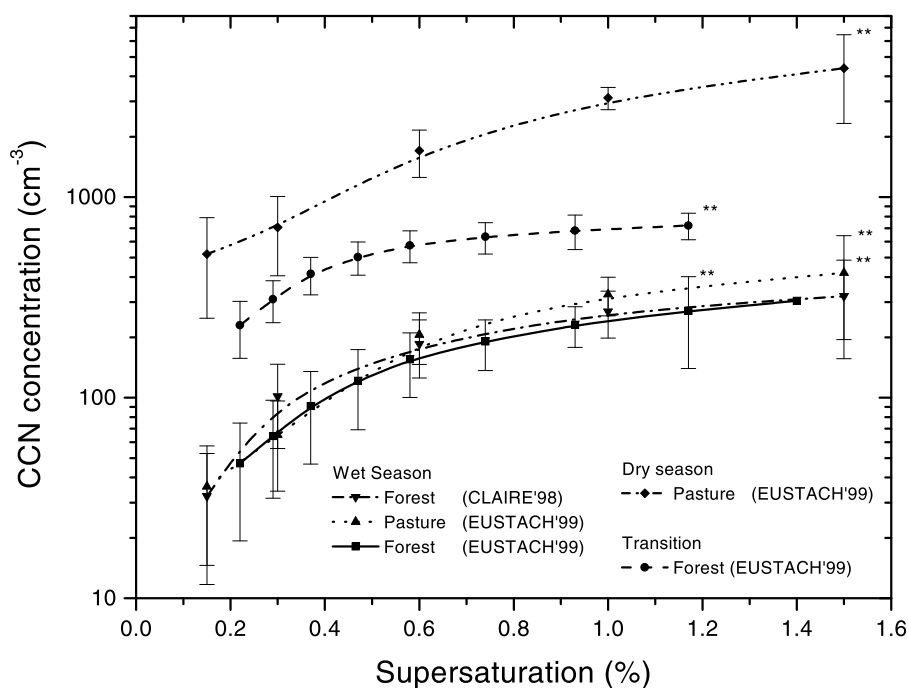


Figure 22. Spectra of cloud condensation nuclei (CCN) activation in the Amazon Basin for a range of natural to very polluted conditions. The highest concentrations occur for locations heavily influenced by in-Basin biomass burning. Adapted from *Roberts et al.* [2003].

that seasonal and annual cycles can be related to other particle parameters, such as chemistry and diameter, as well as to different meteorological conditions.

[110] 2. Regarding the organic components of CCN activity, particle CCN activity depends on both the physical and chemical properties of the particles, especially particle diameter and the water-soluble fraction [*McFiggans et al.*, 2006]. The understanding of the mechanistic processes of CCN activation, in particular, variability in space and time of the organic chemical speciation and most importantly in the water-soluble fraction of Amazonian aerosol particles (see section 3.3.1), must be improved.

[111] 3. Models should be developed that connect the sources of particles and their components with CCN activity, including the further evolution as cloud droplets. This approach should be added to cloud-resolving models, such as the Regional Atmosphere Modeling System (RAMS) or the Weather and Research Forecasting (WRF) model, and large-eddy simulations should be carried out. Interconnections among vegetation, particle chemistry, cloud microphysics, and dynamics could thereby be investigated.

4. OUTLOOK AND FUTURE PRIORITIES

[112] The Amazonian rain forest, having a present area of greater than 5×10^6 km² and corresponding to 85% of its original extent, is the largest on Earth. The future of Amazonian aerosol particles is directly connected to the fate of this forest. Before the 1950s, less than 1% of the area had been deforested. At present, approximately 0.02×10^6 km² is deforested annually, mostly accomplished by biomass

burning. The fate of the forest depends largely on two interacting factors, including (1) the rates and the types of human development in the region and (2) the influences of ongoing global climate change [*Laurance and Williamson*, 2001; *Cardoso et al.*, 2003; *Lewis*, 2006; *Soares-Filho et al.*, 2006; *Malhi et al.*, 2008; *Nepstad et al.*, 2008]. At present, deforestation is largely illegal and driven by business opportunities for cattle ranchers and soy farmers. In the near future, several governments plan to construct a network of criss-crossing roads through the forest (Figure 24). Road access, especially when paved, is the first step of illegal deforestation [*Laurance et al.*, 2001; *Nepstad et al.*, 2008], and forest fragmented by human activities is also more susceptible to drying and subsequent burning, resulting in positive feedbacks between deforestation and wildfire [*Laurance and Williamson*, 2001; *Cardoso et al.*, 2003]. Deforestation by business-as-usual policies is expected to drive cumulative loss of the original extent of the forest from 15% in 2003 to 40% by 2050 (Figure 25) [*Soares-Filho et al.*, 2006]. Even with good governance, however, ongoing global climate change might nevertheless induce large-scale drought [e.g., *Marengo et al.*, 2008]. According to some coupled climate-carbon models, droughts might lead to the loss of most forest in the Amazon Basin and to the conversion to savanna grassland or desert, regardless of activities within political Brazil [*Betts et al.*, 2004; *Saleska et al.*, 2007; *Cox et al.*, 2008]. Some models also indicate that a “tipping point” might occur for a fractional forest cover that falls below a critical value, meaning that deforestation might serve as a feedback that decreases rainfall and thus accelerates deforestation and further reductions in rainfall (i.e., at present,

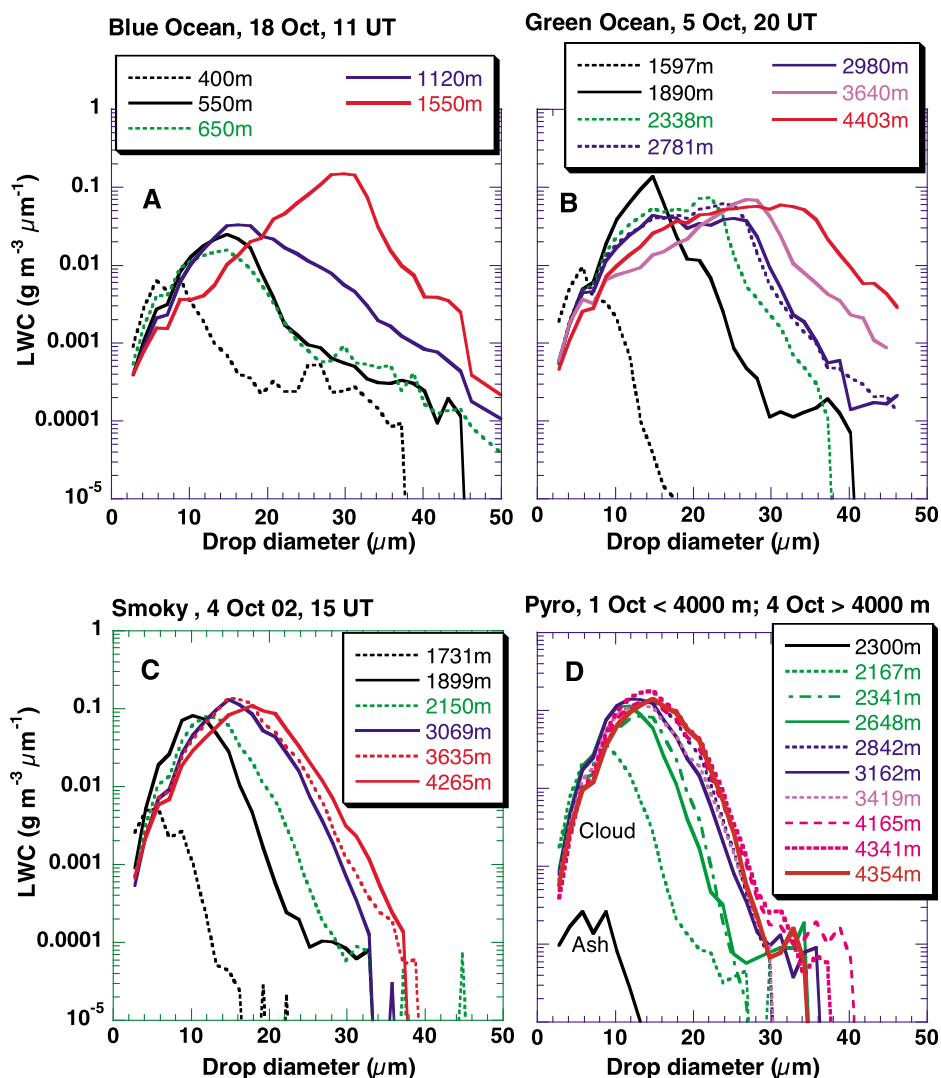


Figure 23. The evolution of the mass-diameter distribution of cloud drops for increasing height in growing convective clouds in regimes of (a) blue ocean, off the northeast Brazilian coast (4°S, 38°W), (b) green ocean (i.e., natural conditions), at the western tip of the Amazon (6°S, 73°W), (c) smoky clouds in Rondônia (10°S, 62°W), and (d) pyroclouds. The lowest-altitude distribution in each plot represents conditions at cloud base, except in Figure 23d for which a distribution for large ash particles outside of the cloud is also shown. Adapted from *Andreae et al.* [2004] with permission from the American Association for the Advancement of Science.

total evapotranspiration approaches 50% of total precipitation), ultimately transforming the region from rain forest to savanna [*Silva Dias et al.*, 2002; *Oyama and Nobre*, 2003; *Nobre et al.*, 2004]. A possible fate of large-scale forest loss has paradigm-shifting implications for the concentrations and the properties of particles as well as for meteorology, cloud type, and rainfall intensity.

[113] According to projections by *Streets* [2007] of the future of the Amazon Basin, emissions of both sulfur dioxide and organic carbon can be expected to grow by 2050 because of high population and economic growth and limited environmental protection. *Heald et al.* [2008] project that a 20% mass increase in the production of particle components resulting from BVOC oxidation can be expected for a business-as-usual scenario, mostly because of changes in climate and emissions over South America by

2100. Nevertheless, this estimate includes uncertain changes in BVOC emissions and land use patterns. Deforestation, for example, might reduce BVOC emissions and therefore reduce some of the increases projected by *Heald et al.* [2008]. Another possibility is that future anthropogenic pollution, in addition to the direct release of primary particles, might enhance oxidant levels and thereby might significantly alter prevailing BVOC oxidation pathways, including the fraction of low-volatility products formed (i.e., those that go on to compose particle components). Changes in land cover might also influence the abundance and the properties of PBA particles and components. The CCN concentrations in the Basin might increase year round, driven both by higher anthropogenic emissions of primary particles and by greater concentrations of sulfuric acid that might increase the rate of new-particle formation. These many possible changes in the

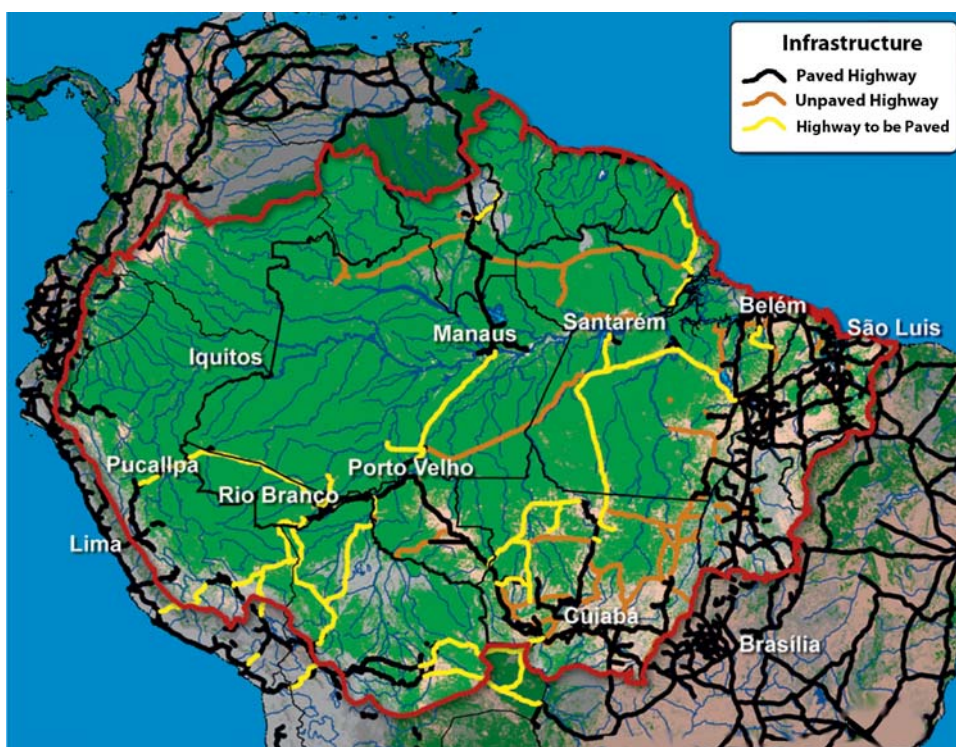


Figure 24. Amazon highways. The government of Brazil plans that the core region of the Amazon will be accessible by all-weather highways (yellow) in the future. Although these investments in transportation promise to lower the production costs of ranching and farming, they potentially threaten to stimulate deforestation. From *Nepstad* [2007] (Woods Hole Research Center).

climatology of Amazonian aerosol particles might, in turn, significantly influence aerosol-cloud-climate interactions in the Basin, providing a feedback to rainfall, vegetation, and climate change [Barth et al., 2005; Andreae and Rosenfeld, 2008]. Davidson and Artaxo [2004] review the in-Basin links between biology and atmospheric properties.

[114] Potentially of particular importance presently and in the future is the export of particles from the Amazon Basin to other parts of the world following deep convection [Andreae et al., 2001], especially during the dry season

when biomass-burning particle concentrations can be very high. Because the ITCZ cuts across the Basin and the general circulation involves the meridional transport of air toward it from both north and south, a substantial fraction of the Amazonian particles experiences deep convection. Whatever does not get scavenged in that process can be injected into the middle and upper troposphere, and some particles can also enter the stratosphere via the Brewer-Dobson circulation. These particles can be transported over long distances between continents. Transported micro-organisms provide

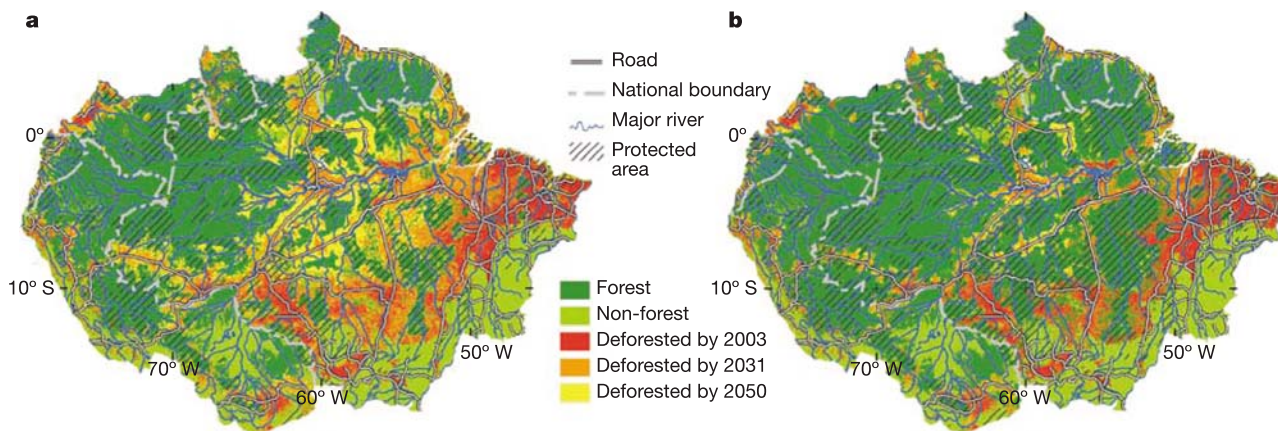


Figure 25. Simulations of forest cover for the year 2050. (a) Scenario of business as usual. (b) Scenario of good governance. Adapted by permission from Macmillan Publishers Ltd [Soares-Filho et al., 2006], copyright 2006.

a clear demonstration of these processes [Griffin *et al.*, 2006; Griffin, 2008]. The Amazon Basin may thus represent a significant global source, both of primary particles and of gaseous precursors to secondary components, and these contributions may strongly influence the pristine conditions that otherwise prevail in the upper troposphere and the stratosphere.

4.1. Priorities for Improved Models

[115] Few modeling studies have focused on aerosol particles in the Amazon Basin. In particular, very little effort has been invested in understanding particle sources for natural conditions when concentrations are low. Regional models have been used, however, to characterize the importance of biomass burning to particle number and mass concentrations as well as to climate [Freitas *et al.*, 2005; Liu, 2005; Martins and Pereira, 2006]. An intercomparison of global models showed that model skill over Amazonia for the annual average optical thickness misleadingly appeared to be good because of compensation by an underestimate of optical thickness for regions influenced by biomass burning and an overestimate for periods during which natural conditions prevailed [Kinne *et al.*, 2003]. The overestimate for natural conditions was puzzling because the global models did not include emissions of PBA particles and had very rudimentary descriptions, if any, of the production of particle-phase BVOC oxidation products. The bias for natural conditions was therefore attributed to one or more of the following: an overestimate of out-of-Basin particles into the region, an underestimate of their in-Basin deposition rates, or a poor characterization of the optical properties of Amazonian particles. Kanakidou *et al.* [2005] summarized the uncertainties and challenges related to global climate model simulations of organic aerosol particles. There is a critical need to validate these model estimates with observations over the Amazon Basin.

[116] High-priority research needs for improving the regional modeling of Amazonian aerosol particles include the following: (1) development of model schemes for emissions of PBA particles in the Amazon Basin (these schemes are entirely absent in state-of-the-art chemical transport models); (2) investigation and implementation of models at the scale necessary to capture how vegetative heterogeneity within the rain forest canopy affects BVOC and PBA particle emissions; (3) inclusion in models at the level of compounds or families of all BVOC emissions contributing to the secondary components of particles (Amazonian vegetation might be a more or less efficient emitter of specific compounds compared to other locations for which standard emissions in models have been calibrated); (4) incorporation in models of new BVOC chemistry, such as the reactions of isoprene in the chemical regimes prevailing in the Amazon Basin (e.g., pristine low NO_x) and more generally of organic peroxy radicals; and (5) attention in models, validated by measurements, of how efficiently particles over the Amazon Basin are removed by precipitation and how this sink term is affected by the processing and alteration of particles during their residence in the atmosphere.

4.2. Priorities for Improved Measurements

[117] There have been many technological advances in the past 10 years for the characterization of aerosol particles globally, yet many of the new instruments have yet to be deployed in the Amazon Basin. The logistical difficulties there have constrained measurements temporally, spatially, and technically to levels insufficient for fully accurate descriptions of Amazonian aerosol particles and the processes affecting them. These difficulties notwithstanding, new instruments, defined in the context of Amazonian aerosol particles both as truly new instruments in the broad scientific community and as more familiar instruments that have never been deployed before for studies in the Basin, should be used to obtain more precise and accurate measurements of key properties of Amazonian particles. The instruments should be deployed to provide complementary information on complex properties, especially related to particle chemical composition. Chemical information can be employed, in conjunction with models, to understand mechanisms of particle formation and subsequent aging processes. In addition, more creative applications should be made of real-time displays and airborne remote sensors for better placement of aircraft, particularly when sampling clouds.

[118] The particle properties that are the most uncertain and thus limit our ability to assess their environmental effects include (1) the molecular composition of the organic component as quantitative tracers of sources and age, (2) the hygroscopicity and mixing states as affected by atmospheric processing, (3) the activity as cloud and ice nuclei, and (4) the optical activity (i.e., extinction, absorption, single-scattering albedo, and asymmetry factor). State-of-the-art measurement systems have the potential to significantly reduce the uncertainties surrounding these properties. For example, improvements in measuring the chemistry of aerosol particles with instruments like mass spectrometers [Prather *et al.*, 1994; Jayne *et al.*, 2000] have provided a detailed look into the chemistry of particles, yet an aerosol mass spectrometer went to Amazonia for the first time only recently as part of AMAZE-08 [Chen *et al.*, 2009]. Similarly, single-column and multicolumn continuous flow CCN counters, including instruments with size-selective inlets, have opened new possibilities for measurements of particle hygroscopicity [Roberts and Nenes, 2005], and these state-of-the-art instruments also were deployed in Amazonia for the first time in 2008 as part of AMAZE-08. There are many other instruments that are currently under development or that have only just been made operational. For example, the single-particle soot photometer is the first instrument to measure the mass in single, light-absorbing carbon particles, and its use in studying biomass-burning particles could improve the connections between the particle chemistry and the particle radiation field [Baumgardner *et al.*, 2004; Schwarz *et al.*, 2006; Moteki *et al.*, 2007]. When an aircraft is available, airborne lidar and radar with real-time displays could be used to locate optimum areas for cloud penetration to understand and quantify the effects of cloud processing on particles.

[119] The use of measurements to understand the evolution of Amazonian aerosol particles, their interactions with clouds and radiation, and their impacts on climate must ultimately be facilitated by global climate models coupled with chemical transport models. These models, however, must be based on particle properties and processes defined by a combination of laboratory and ambient measurements, and significant uncertainties exist in the treatments presently employed for Amazonian aerosol particles (see section 4.1), traceable in part to an insufficient set of measurements. As an example of how the uncertainty in measurements propagates into models, the optical thickness predicted by several different global models varies by more than 50% in the Basin [Kinne *et al.*, 2003, 2006]. These intermodel differences are primarily attributable to uncertainties in the parameterization of particle composition and residence time [Textor *et al.*, 2006]. Reductions in these uncertainties requires knowledge, to be gained through well-designed measurement programs, that focuses on closure studies, where the closure to be achieved is between the predicted and measured particle properties, particularly their hygroscopic, chemical, microphysical, and optical properties. Another important closure, though less precisely defined, is that of particle residence time, meaning an evaluation of sources and sinks.

[120] As one example, obtaining closure for CCN activation requires measuring the chemical composition of particles as a function of size and accurately predicting the number-diameter distribution of CCN as a function of supersaturation that is measured with in situ instrumentation. Gunthe *et al.* [2009] provide one example for studies in the Amazon Basin. The results of this type of closure should be incorporated into modules of cloud microphysics to improve how the aerosol indirect effect is forecast. As a second example, radiation closure requires ground-based and satellite measurements of optical thickness, at multiple wavelengths, to compare with the optical thickness derived from measurements of the chemical and optical properties of particles over a range of altitudes and geographic locations.

[121] The ideal field program to implement the above closure studies would span several dry and wet seasons in the Basin, would require in situ and remote sensing measurements from multiple ground-based and airborne platforms, and would be complemented by satellite observations from which particle and trace gas properties are derived. Autonomous, unmanned aerial vehicles have recently been used to make measurements of particles over the Indian Ocean [Corrigan *et al.*, 2008] and could be extremely useful in the Basin. Establishment of a tall-tower atmospheric observatory could greatly facilitate long-term, high-quality measurements of particle properties, gaseous tracers, and meteorological variables. A tall tower could additionally provide the opportunity to measure vertical profiles through the atmosphere over the forest.

5. CONCLUDING REMARKS

[122] The goal of this review was to provide a synthesis of aerosol research in the Amazon Basin, most of which has

been published during the last two decades. By integrating the information that has been published by a considerable number of authors in a variety of journals over a long span of time, we sought to provide a comprehensive picture about what is known about Amazonian aerosol particles and to bring together the various aspects that are now scattered throughout the literature. We intend that this effort will facilitate an understanding of the current state of knowledge on Amazonian aerosol particles specifically and tropical continental aerosol particles in general and will thereby enhance future research in this area. This review barely touched on the transformations of particles by interactions between particle- and gas-phase species (i.e., condensation, evaporation, and reactions), reactions within the particles, and cloud processing of particles. These processes, though undoubtedly important, have hardly been researched in the Amazon Basin. We therefore identify future studies designed toward these ends as being urgently needed.

[123] **ACKNOWLEDGMENTS.** The authors of this article were part of the International Workshop: Aerosol Particles in the Amazon—Changes and Their Consequences From Past and Future Human Activities, which took place 18–22 February 2008 at ARIAÚ Amazon Towers, north of Manaus, Brazil. Special thanks go to C. H. Martin. Support was received from the U.S. National Science Foundation (NSF) (OISE-0651836), the Brazil LBA Millennium Institute, the Harvard Brazil Studies Program of the David Rockefeller Center of Latin American Studies, Atmospheric Composition Change: The European Network of Excellence (ACCENT), the Max Planck Society, and the Brazilian Large-Scale Biosphere-Atmosphere (LBA) Experiment. S.T.M. also acknowledges support from a Humboldt Research Fellowship: Summer Research Fellowship for U.S. Scientists and Scholars and a Visiting Researcher Award, State of São Paulo Research Foundation (FAPESP), Brazil. This article is dedicated to the memory of Conceição Moreira Silva.

[124] The Editor responsible for this paper was Gerald North. He thanks two anonymous reviewers.

REFERENCES

- Allen, A. G., and A. H. Miguel (1995), Biomass burning in the Amazon—Characterization of the ionic component of aerosols generated from flaming and smoldering rain-forest and savanna, *Environ. Sci. Technol.*, *29*, 486–493, doi:10.1021/es00002a026.
- Andreae, M. O. (1993), The influence of tropical biomass burning on climate and the atmospheric environment, in *Biogeochemistry of Global Change: Radiatively Active Trace Gases*, edited by R. S. Oremland, pp. 113–150, Chapman and Hall, New York.
- Andreae, M. O. (2007), Aerosols before pollution, *Science*, *315*, 50–51, doi:10.1126/science.1136529.
- Andreae, M. O. (2009), Correlation between cloud condensation nuclei concentration and aerosol optical thickness in remote and polluted regions, *Atmos. Chem. Phys.*, *9*, 543–556.
- Andreae, M. O., and T. W. Andreae (1988), The cycle of biogenic sulfur-compounds over the Amazon Basin: 1. Dry season, *J. Geophys. Res.*, *93*, 1487–1497, doi:10.1029/JD093iD02p01487.
- Andreae, M. O., and P. J. Crutzen (1997), Atmospheric aerosols: Biogeochemical sources and role in atmospheric chemistry, *Science*, *276*, 1052–1058, doi:10.1126/science.276.5315.1052.
- Andreae, M. O., and A. Gelencser (2006), Black carbon or brown carbon? The nature of light-absorbing carbonaceous aerosols, *Atmos. Chem. Phys.*, *6*, 3131–3148.

- Andreae, M. O., and D. Rosenfeld (2008), Aerosol-cloud-precipitation interactions. Part 1. The nature and sources of cloud-active aerosols, *Earth Sci. Rev.*, *89*, 13–41, doi:10.1016/j.earscirev.2008.03.001.
- Andreae, M. O., et al. (1988), Biomass-burning emissions and associated haze layers over Amazonia, *J. Geophys. Res.*, *93*, 1509–1527, doi:10.1029/JD093iD02p01509.
- Andreae, M. O., H. Berresheim, H. Bingemer, D. J. Jacob, B. L. Lewis, S. M. Li, and R. W. Talbot (1990a), The atmospheric sulfur cycle over the Amazon Basin: 2. Wet season, *J. Geophys. Res.*, *95*, 16,813–16,824, doi:10.1029/JD095iD10p16813.
- Andreae, M. O., R. W. Talbot, H. Berresheim, and K. M. Beecher (1990b), Precipitation chemistry in central Amazonia, *J. Geophys. Res.*, *95*, 16,987–16,999, doi:10.1029/JD095iD10p16987.
- Andreae, M. O., B. E. Anderson, D. R. Blake, J. D. Bradshaw, J. E. Collins, G. L. Gregory, G. W. Sachse, and M. C. Shipham (1994), Influence of plumes from biomass burning on atmospheric chemistry over the equatorial Atlantic during CITE-3, *J. Geophys. Res.*, *99*, 12,793–12,808.
- Andreae, M. O., et al. (1998), Airborne studies of aerosol emissions from savanna fires in southern Africa: 2. Aerosol chemical composition, *J. Geophys. Res.*, *103*, 32,119–32,128, doi:10.1029/98JD02280.
- Andreae, M. O., et al. (2001), Transport of biomass burning smoke to the upper troposphere by deep convection in the equatorial region, *Geophys. Res. Lett.*, *28*, 951–954, doi:10.1029/2000GL012391.
- Andreae, M. O., et al. (2002), Biogeochemical cycling of carbon, water, energy, trace gases, and aerosols in Amazonia: The LBA-EUSTACH experiments, *J. Geophys. Res.*, *107*(D20), 8066, doi:10.1029/2001JD000524.
- Andreae, M. O., D. Rosenfeld, P. Artaxo, A. A. Costa, G. P. Frank, K. M. Longo, and M. A. F. Silva-Dias (2004), Smoking rain clouds over the Amazon, *Science*, *303*, 1337–1342, doi:10.1126/science.1092779.
- Artaxo, P., and H. C. Hansson (1995), Size distribution of biogenic aerosol-particles from the Amazon Basin, *Atmos. Environ.*, *29*, 393–402, doi:10.1016/1352-2310(94)00178-N.
- Artaxo, P., H. Storms, F. Bruynseels, R. Van Grieken, and W. Maenhaut (1988), Composition and sources of aerosols from the Amazon Basin, *J. Geophys. Res.*, *93*, 1605–1615, doi:10.1029/JD093iD02p01605.
- Artaxo, P., W. Maenhaut, H. Storms, and R. Van Grieken (1990), Aerosol characteristics and sources for the Amazon Basin during the wet season, *J. Geophys. Res.*, *95*, 16,971–16,985, doi:10.1029/JD095iD10p16971.
- Artaxo, P., E. T. Fernandes, J. V. Martins, M. A. Yamasoe, P. V. Hobbs, W. Maenhaut, K. M. Longo, and A. Castanho (1998), Large-scale aerosol source apportionment in Amazonia, *J. Geophys. Res.*, *103*, 31,837–31,847, doi:10.1029/98JD02346.
- Artaxo, P., J. V. Martins, M. A. Yamasoe, A. S. Procopio, T. M. Pauliquevis, M. O. Andreae, P. Guyon, L. V. Gatti, and A. M. C. Leal (2002), Physical and chemical properties of aerosols in the wet and dry seasons in Rondônia, Amazonia, *J. Geophys. Res.*, *107*(D20), 8081, doi:10.1029/2001JD000666.
- Avissar, R., P. L. S. Dias, M. Dias, and C. Nobre (2002), The Large-Scale Biosphere-Atmosphere Experiment in Amazonia (LBA): Insights and future research needs, *J. Geophys. Res.*, *107*(D20), 8086, doi:10.1029/2002JD002704.
- Barkley, M. P., P. I. Palmer, U. Kuhn, J. Kesselmeier, K. Chance, T. P. Kurosui, R. V. Martin, D. Helmig, and A. Guenther (2008), Net ecosystem fluxes of isoprene over tropical South America inferred from GOME observations of HCHO columns, *J. Geophys. Res.*, *113*, D20304, doi:10.1029/2008JD009863.
- Barregard, L., G. Sallsten, P. Gustafson, L. Andersson, L. Johansson, S. Basu, and L. Stigendal (2006), Experimental exposure to wood-smoke particles in healthy humans: Effects on markers of inflammation, coagulation, and lipid peroxidation, *Inhal. Toxicol.*, *18*, 845–853, doi:10.1080/08958370600685798.
- Barth, M., et al. (2005), Coupling between land ecosystems and the atmospheric hydrologic cycle through biogenic aerosol pathways, *Bull. Am. Meteorol. Soc.*, *86*, 1738–1742, doi:10.1175/BAMS-86-12-1738.
- Baumgardner, D., G. Kok, and G. Raga (2004), Warming of the Arctic lower stratosphere by light absorbing particles, *Geophys. Res. Lett.*, *31*, L06117, doi:10.1029/2003GL018883.
- Betts, R. A., P. M. Cox, M. Collins, P. P. Harris, C. Huntingford, and C. D. Jones (2004), The role of ecosystem-atmosphere interactions in simulated Amazonian precipitation decrease and forest dieback under global climate warming, *Theor. Appl. Climatol.*, *78*, 157–175.
- Blando, J. D., and B. J. Turpin (2000), Secondary organic aerosol formation in cloud and fog droplets: A literature evaluation of plausibility, *Atmos. Environ.*, *34*, 1623–1632, doi:10.1016/S1352-2310(99)00392-1.
- Browell, E. V., G. L. Gregory, R. C. Harriss, and V. Kirchner (1990), Ozone and aerosol distributions over the Amazon Basin during the wet season, *J. Geophys. Res.*, *95*, 16,887–16,901, doi:10.1029/JD095iD10p16887.
- Cardoso, M. F., G. C. Hurtt, B. Moore, C. A. Nobre, and E. M. Prins (2003), Projecting future fire activity in Amazonia, *Global Change Biol.*, *9*, 656–669, doi:10.1046/j.1365-2486.2003.00607.x.
- Carlton, A. G., B. J. Turpin, H.-J. Lim, K. E. Altieri, and S. P. Seitzinger (2006), Link between isoprene and secondary organic aerosol (SOA): Pyruvic acid oxidation yields low volatility organic acids in clouds, *Geophys. Res. Lett.*, *33*, L06822, doi:10.1029/2005GL025374.
- Ceburnis, D., C. D. O'Dowd, G. S. Jennings, M. C. Facchini, L. Emblico, S. Decesari, S. Fuzzi, and J. Sakalys (2008), Marine aerosol chemistry gradients: Elucidating primary and secondary processes and fluxes, *Geophys. Res. Lett.*, *35*, L07804, doi:10.1029/2008GL033462.
- Chand, D., P. Guyon, P. Artaxo, O. Schmid, G. P. Frank, L. V. Rizzo, O. L. Mayol-Bracero, L. V. Gatti, and M. O. Andreae (2006), Optical and physical properties of aerosols in the boundary layer and free troposphere over the Amazon Basin during the biomass burning season, *Atmos. Chem. Phys.*, *6*, 2911–2925.
- Chen, Q., et al. (2009), Mass spectral characterization of submicron biogenic organic particles in the Amazon Basin, *Geophys. Res. Lett.*, *36*, L20806, doi:10.1029/2009GL039880.
- Chung, S. H., and J. H. Seinfeld (2002), Global distribution and climate forcing of carbonaceous aerosols, *J. Geophys. Res.*, *107*(D19), 4407, doi:10.1029/2001JD001397.
- Ciccioli, P., et al. (1999), Emission of reactive terpene compounds from orange orchards and their removal by within-canopy processes, *J. Geophys. Res.*, *104*, 8077–8094, doi:10.1029/1998JD100026.
- Claeys, M., et al. (2004), Formation of secondary organic aerosols through photooxidation of isoprene, *Science*, *303*, 1173–1176, doi:10.1126/science.1092805.
- Clarke, A. D., J. L. Varner, F. Eisele, R. L. Mauldin, D. Tanner, and M. Litchy (1998), Particle production in the remote marine atmosphere: Cloud outflow and subsidence during ACE 1, *J. Geophys. Res.*, *103*, 16,397–16,409, doi:10.1029/97JD02987.
- Clarke, A. D., V. N. Kapustin, F. L. Eisele, R. J. Weber, and P. H. McMurry (1999), Particle production near marine clouds: Sulfuric acid and predictions from classical binary nucleation, *Geophys. Res. Lett.*, *26*, 2425–2428, doi:10.1029/1999GL000438.
- Corrigan, C. E., G. C. Roberts, M. V. Ramana, D. Kim, and V. Ramanathan (2008), Capturing vertical profiles of aerosols and black carbon over the Indian Ocean using autonomous unmanned aerial vehicles, *Atmos. Chem. Phys.*, *8*, 737–747.
- Cox, P. M., P. P. Harris, C. Huntingford, R. A. Betts, M. Collins, C. D. Jones, T. E. Jupp, J. A. Marengo, and C. A. Nobre (2008), Increasing risk of Amazonian drought due to decreasing aerosol pollution, *Nature*, *453*, 212–217, doi:10.1038/nature06960.

- Davidson, E. A., and P. Artaxo (2004), Globally significant changes in biological processes of the Amazon Basin: Results of the Large-scale Biosphere-Atmosphere Experiment, *Global Change Biol.*, *10*, 519–529, doi:10.1111/j.1529-8817.2003.00779.x.
- Decesari, S., et al. (2006), Characterization of the organic composition of aerosols from Rondonia, Brazil, during the LBA-SMOCC 2002 experiment and its representation through model compounds, *Atmos. Chem. Phys.*, *6*, 375–402.
- Despres, V. R., J. F. Nowoisky, M. Klose, R. Conrad, M. O. Andreae, and U. Pöschl (2007), Characterization of primary biogenic aerosol particles in urban, rural, and high-alpine air by DNA sequence and restriction fragment analysis of ribosomal RNA genes, *Biogeosciences*, *4*, 1127–1141.
- Echalar, F., P. Artaxo, J. V. Martins, M. Yamasoe, F. Gerab, W. Maenhaut, and B. Holben (1998), Long-term monitoring of atmospheric aerosols in the Amazon Basin: Source identification and apportionment, *J. Geophys. Res.*, *103*, 31,849–31,864, doi:10.1029/98JD01749.
- Ekman, A. M. L., R. Krejci, A. Engstrom, J. Strom, M. de Reus, J. Williams, and M. O. Andreae (2008), Do organics contribute to small particle formation in the Amazonian upper troposphere?, *Geophys. Res. Lett.*, *35*, L17810, doi:10.1029/2008GL034970.
- Elbert, W., P. E. Taylor, M. O. Andreae, and U. Pöschl (2007), Contribution of fungi to primary biogenic aerosols in the atmosphere: Wet and dry discharged spores, carbohydrates, and inorganic ions, *Atmos. Chem. Phys.*, *7*, 4569–4588.
- Falkovich, A. H., E. R. Graber, G. Schkolnik, Y. Rudich, W. Maenhaut, and P. Artaxo (2005), Low molecular weight organic acids in aerosol particles from Rondonia, Brazil, during the biomass-burning, transition and wet periods, *Atmos. Chem. Phys.*, *5*, 781–797.
- Feingold, G., H. L. Jiang, and J. Y. Harrington (2005), On smoke suppression of clouds in Amazonia, *Geophys. Res. Lett.*, *32*, L02804, doi:10.1029/2004GL021369.
- Ferek, R. J., J. S. Reid, P. V. Hobbs, D. R. Blake, and C. Liousse (1998), Emission factors of hydrocarbons, halocarbons, trace gases and particles from biomass burning in Brazil, *J. Geophys. Res.*, *103*, 32,107–32,118, doi:10.1029/98JD00692.
- Fiedler, V., et al. (2005), The contribution of sulphuric acid to atmospheric particle formation and growth: A comparison between boundary layers in Northern and Central Europe, *Atmos. Chem. Phys.*, *5*, 1773–1785.
- Fisch, G., J. Tota, L. A. T. Machado, M. Dias, R. F. D. Lyra, C. A. Nobre, A. J. Dolman, and J. H. C. Gash (2004), The convective boundary layer over pasture and forest in Amazonia, *Theor. Appl. Climatol.*, *78*, 47–59.
- Formenti, P., M. O. Andreae, L. Lange, G. Roberts, J. Cafmeyer, I. Rajta, W. Maenhaut, B. N. Holben, P. Artaxo, and J. Lelieveld (2001), Saharan dust in Brazil and Suriname during the Large-Scale Biosphere-Atmosphere Experiment in Amazonia (LBA)—Cooperative LBA Regional Experiment (CLAIRE) in March 1998, *J. Geophys. Res.*, *106*, 14,919–14,934, doi:10.1029/2000JD900827.
- Freitas, S. R., M. A. F. S. Dias, and P. L. S. Dias (2000), Modeling the convective transport of trace gases by deep and moist convection, *Hybrid Methods Eng.*, *2*, 317–330.
- Freitas, S. R., K. M. Longo, M. Dias, P. L. S. Dias, R. Chatfield, E. Prins, P. Artaxo, G. A. Grell, and F. S. Recuero (2005), Monitoring the transport of biomass-burning emissions in South America, *Environ. Fluid Mech.*, *5*, 135–167, doi:10.1007/s10652-005-0243-7.
- Fuzzi, S., et al. (2006), Critical assessment of the current state of scientific knowledge, terminology, and research needs concerning the role of organic aerosols in the atmosphere, climate, and global change, *Atmos. Chem. Phys.*, *6*, 2017–2038.
- Fuzzi, S., et al. (2007), Overview of the inorganic and organic composition of size-segregated aerosol in Rondonia, Brazil, from the biomass-burning period to the onset of the wet season, *J. Geophys. Res.*, *112*, D01201, doi:10.1029/2005JD006741.
- Garreaud, R. D., and J. M. Wallace (1998), Summertime incursions of midlatitude air into subtropical and tropical South America, *Mon. Weather Rev.*, *126*, 2713–2733, doi:10.1175/1520-0493(1998)126<2713:SIOMAI>2.0.CO;2.
- Gerab, F., P. Artaxo, R. Gillett, and G. Ayers (1998), PIXE, PIGE and ion chromatography of aerosol particles from northeast Amazon Basin, *Nucl. Instrum. Methods Phys. Res.*, *137*, 955–960.
- Giglio, L., I. Csizsar, and C. O. Justice (2006), Global distribution and seasonality of active fires as observed with the Terra and Aqua MODIS sensors, *J. Geophys. Res.*, *111*, G02016, doi:10.1029/2005JG000142.
- Gilbert, G. S. (2005), Nocturnal fungi: Airborne spores in the canopy and understory of a tropical rain forest, *Biotropica*, *37*, 462–464, doi:10.1111/j.1744-7429.2005.00061.x.
- Goldstein, A. H., and I. E. Galbally (2007), Known and unexplored organic constituents in the Earth's atmosphere, *Environ. Sci. Technol.*, *41*, 1514–1521, doi:10.1021/es072476p.
- Graham, B., O. L. Mayol-Bracero, P. Guyon, G. C. Roberts, S. Decesari, M. C. Facchini, P. Artaxo, W. Maenhaut, P. Koll, and M. O. Andreae (2002), Water-soluble organic compounds in biomass burning aerosols over Amazonia: 1. Characterization by NMR and GC-MS, *J. Geophys. Res.*, *107*(D20), 8047, doi:10.1029/2001JD000336.
- Graham, B., et al. (2003a), Composition and diurnal variability of the natural Amazonian aerosol, *J. Geophys. Res.*, *108*(D24), 4765, doi:10.1029/2003JD004049.
- Graham, B., P. Guyon, P. E. Taylor, P. Artaxo, W. Maenhaut, M. M. Glovsky, R. C. Flagan, and M. O. Andreae (2003b), Organic compounds present in the natural Amazonian aerosol: Characterization by gas chromatography–mass spectrometry, *J. Geophys. Res.*, *108*(D24), 4766, doi:10.1029/2003JD003990.
- Griffin, D. W. (2008), Non-spore forming eubacteria isolated at an altitude of 20,000 m in Earth's atmosphere: Extended incubation periods needed for culture-based assays, *Aerobiologia*, *24*, 19–25, doi:10.1007/s10453-007-9078-7.
- Griffin, D. W., D. L. Westphal, and M. A. Gray (2006), Airborne microorganisms in the African desert dust corridor over the mid-Atlantic ridge, Ocean Drilling Program, Leg 209, *Aerobiologia*, *22*, 211–226, doi:10.1007/s10453-006-9033-z.
- Guenther, A., et al. (1995), A global model of natural volatile organic compound emissions, *J. Geophys. Res.*, *100*, 8873–8892, doi:10.1029/94JD02950.
- Guenther, A., T. Karl, P. Harley, C. Wiedinmyer, P. I. Palmer, and C. Geron (2006), Estimates of global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from Nature), *Atmos. Chem. Phys.*, *6*, 3181–3210.
- Gunthe, S. S., et al. (2009), Cloud condensation nuclei in pristine tropical rainforest air of Amazonia: Size-resolved measurements and modeling of atmospheric aerosol composition and CCN activity, *Atmos. Chem. Phys.*, *9*, 7551–7575.
- Guyon, P., O. Boucher, B. Graham, J. Beck, O. L. Mayol-Bracero, G. C. Roberts, W. Maenhaut, P. Artaxo, and M. O. Andreae (2003a), Refractive index of aerosol particles over the Amazon tropical forest during LBA-EUSTACH 1999, *J. Aerosol Sci.*, *34*, 883–907, doi:10.1016/S0021-8502(03)00052-1.
- Guyon, P., B. Graham, J. Beck, O. Boucher, E. Gerasopoulos, O. L. Mayol-Bracero, G. C. Roberts, P. Artaxo, and M. O. Andreae (2003b), Physical properties and concentration of aerosol particles over the Amazon tropical forest during background and biomass burning conditions, *Atmos. Chem. Phys.*, *3*, 951–967.
- Guyon, P., B. Graham, G. C. Roberts, O. L. Mayol-Bracero, W. Maenhaut, P. Artaxo, and M. O. Andreae (2003c), In-canopy gradients, composition, sources, and optical properties of aerosol over the Amazon forest, *J. Geophys. Res.*, *108*(D18), 4591, doi:10.1029/2003JD003465.

- Guyon, P., et al. (2005), Airborne measurements of trace gas and aerosol particle emissions from biomass burning in Amazonia, *Atmos. Chem. Phys.*, *5*, 2989–3002.
- Harley, P., et al. (2004), Variation in potential for isoprene emissions among Neotropical forest sites, *Global Change Biol.*, *10*, 630–650, doi:10.1111/j.1529-8817.2003.00760.x.
- Harriss, R. C., et al. (1988), The Amazon Boundary Layer Experiment (ABLE-2A): Dry season 1985, *J. Geophys. Res.*, *93*, 1351–1360, doi:10.1029/JD093iD02p01351.
- Harriss, R. C., et al. (1990), The Amazon Boundary Layer Experiment: Wet season 1987, *J. Geophys. Res.*, *95*, 16,721–16,736, doi:10.1029/JD095iD10p16721.
- Heald, C. L., et al. (2008), Predicted change in global secondary organic aerosol concentrations in response to future climate, emissions, and land use change, *J. Geophys. Res.*, *113*, D05211, doi:10.1029/2007JD009092.
- Henze, D. K., and J. H. Seinfeld (2006), Global secondary organic aerosol from isoprene oxidation, *Geophys. Res. Lett.*, *33*, L09812, doi:10.1029/2006GL025976.
- Hines, P. J. (2006), The invisible bouquet—Introduction, *Science*, *311*, 803–803, doi:10.1126/science.311.5762.803.
- Hoffer, A., A. Gelencser, M. Blazso, P. Guyon, P. Artaxo, and M. O. Andreae (2006), Diel and seasonal variations in the chemical composition of biomass burning aerosol, *Atmos. Chem. Phys.*, *6*, 3505–3515.
- Holzinger, R., A. Lee, K. T. Paw, and A. H. Goldstein (2005), Observations of oxidation products above a forest imply biogenic emissions of very reactive compounds, *Atmos. Chem. Phys.*, *5*, 67–75.
- Hoppel, W. A., G. M. Frick, and R. E. Larson (1986), Effects of non-precipitating clouds on the aerosol size distribution in the marine boundary layer, *Geophys. Res. Lett.*, *13*, 125–128, doi:10.1029/GL013i002p00125.
- Hoppel, W. A., G. M. Frick, J. W. Fitzgerald, and R. E. Larson (1994), Marine boundary layer measurements of new particle formation and the effects of nonprecipitating clouds on the aerosol size distribution, *J. Geophys. Res.*, *99*, 14,443–14,459, doi:10.1029/94JD00797.
- Hörrak, U., H. Iher, A. Luts, J. Salm, and H. Tammet (1994), Mobility spectrum of air ions at Tahkuse Observatory, *J. Geophys. Res.*, *99*, 10,697–10,700, doi:10.1029/93JD02291.
- Hoyle, C. R., T. Berntsen, G. Myhre, and I. S. A. Isaksen (2007), Secondary organic aerosol in the global aerosol—Chemical transport model Oslo CTM2, *Atmos. Chem. Phys.*, *7*, 5675–5694.
- Ignotti, E., J. Valente, S. Hacon, K. Longo, and S. Freitas (2007), Proportional mortality due to respiratory diseases (RD) and exposure to PM_{2.5} in Amazon region, *Epidemiology*, *18*, S94–S94.
- Iida, K., M. R. Stolzenburg, P. H. McMurry, and J. N. Smith (2008), Estimating nanoparticle growth rates from size-dependent charged fractions: Analysis of new particle formation events in Mexico City, *J. Geophys. Res.*, *113*, D05207, doi:10.1029/2007JD009260.
- Intergovernmental Panel on Climate Change (2007), *Climate Change 2007: The Physical Science Basis—Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by S. Solomon et al., 996 pp., Cambridge Univ. Press, Cambridge, U. K.
- Jacob, D. J., and S. C. Wofsy (1988), Photochemistry of biogenic emissions over the Amazon forest, *J. Geophys. Res.*, *93*, 1477–1486, doi:10.1029/JD093iD02p01477.
- Jacob, D. J., and S. C. Wofsy (1990), Budgets of reactive nitrogen, hydrocarbons, and ozone over the Amazon forest during the wet season, *J. Geophys. Res.*, *95*, 16,737–16,754, doi:10.1029/JD095iD10p16737.
- Jacobson, M. C., H.-C. Hansson, K. J. Noone, and R. J. Charlson (2000), Organic atmospheric aerosols: Review and state of the science, *Rev. Geophys.*, *38*, 267–294, doi:10.1029/1998RG000045.
- Jaenicke, R. (2005), Abundance of cellular material and proteins in the atmosphere, *Science*, *308*, 73, doi:10.1126/science.1106335.
- Jayne, J. T., D. C. Leard, X. F. Zhang, P. Davidovits, K. A. Smith, C. E. Kolb, and D. R. Worsnop (2000), Development of an aerosol mass spectrometer for size and composition analysis of sub-micron particles, *Aerosol Sci. Technol.*, *33*, 49–70, doi:10.1080/027868200410840.
- Jiang, H. L., and G. Feingold (2006), Effect of aerosol on warm convective clouds: Aerosol-cloud-surface flux feedbacks in a new coupled large eddy model, *J. Geophys. Res.*, *111*, D01202, doi:10.1029/2005JD006138.
- Jones, C., and L. M. V. Carvalho (2002), Active and break phases in the South American monsoon system, *J. Clim.*, *15*, 905–914, doi:10.1175/1520-0442(2002)015<0905:AABPIT>2.0.CO;2.
- Kalnay, E., et al. (1996), The NCEP/NCAR 40-year reanalysis project, *Bull. Am. Meteorol. Soc.*, *77*, 437–471, doi:10.1175/1520-0477(1996)077<0437:TNYRP>2.0.CO;2.
- Kanakidou, M., et al. (2005), Organic aerosol and global climate modelling: A review, *Atmos. Chem. Phys.*, *5*, 1053–1123.
- Karl, T., A. Guenther, R. J. Yokelson, J. Greenberg, M. Potosnak, D. R. Blake, and P. Artaxo (2007), The tropical forest and fire emissions experiment: Emission, chemistry, and transport of biogenic volatile organic compounds in the lower atmosphere over Amazonia, *J. Geophys. Res.*, *112*, D18302, doi:10.1029/2007JD008539.
- Karl, T., A. Guenther, A. Turnipseed, P. Artaxo, and S. Martin (2009), Rapid formation of isoprene photo-oxidation products observed in Amazonia, *Atmos. Chem. Phys.*, *9*, 7753–7767.
- Kaufman, Y. J., et al. (1998), Smoke, Clouds, and Radiation—Brazil (SCAR-B) experiment, *J. Geophys. Res.*, *103*, 31,783–31,808, doi:10.1029/98JD02281.
- Keller, M., M. Bustamante, J. Gash, and P. Dias (Eds.) (2009), *Amazonia and Global Change, Geophys. Monogr. Ser.*, vol. 186, AGU, Washington, D. C.
- Kerminen, V. M., and M. Kulmala (2002), Analytical formulae connecting the “real” and the “apparent” nucleation rate and the nuclei number concentration for atmospheric nucleation events, *J. Aerosol Sci.*, *33*, 609–622, doi:10.1016/S0021-8502(01)00194-X.
- Kesselmeier, J., et al. (2000), Atmospheric volatile organic compounds (VOC) at a remote tropical forest site in central Amazonia, *Atmos. Environ.*, *34*, 4063–4072, doi:10.1016/S1352-2310(00)00186-2.
- Kesselmeier, J., A. Guenther, T. Hoffmann, M. T. Piedade, and J. Warnke (2009), Natural volatile organic compound emissions from plants and their roles in oxidant balance and particle formation, in *Amazonia and Global Change, Geophys. Monogr. Ser.*, vol. 186, edited by M. Keller et al., pp. 183–206, AGU, Washington, D. C.
- King, S. M., T. Rosenoern, J. E. Shilling, Q. Chen, and S. T. Martin (2007), Cloud condensation nucleus activity of secondary organic aerosol particles mixed with sulfate, *Geophys. Res. Lett.*, *34*, L24806, doi:10.1029/2007GL030390.
- King, S. M., T. Rosenoern, J. E. Shilling, Q. Chen, and S. T. Martin (2009), Increased cloud activation potential of secondary organic aerosol for atmospheric mass loadings, *Atmos. Chem. Phys.*, *9*, 2959–2972.
- Kinne, S., et al. (2003), Monthly averages of aerosol properties: A global comparison among models, satellite data, and AERONET ground data, *J. Geophys. Res.*, *108*(D20), 4634, doi:10.1029/2001JD001253.
- Kinne, S., et al. (2006), An AeroCom initial assessment—Optical properties in aerosol component modules of global models, *Atmos. Chem. Phys.*, *6*, 1815–1834.
- Kodama, Y. (1992), Large-scale common features of subtropical precipitation zones (the Baiu frontal zone, the SPCZ, and the SACZ). 1. Characteristics of subtropical frontal zones, *J. Meteorol. Soc. Jpn.*, *70*, 813–836.

- Komppula, M., S. L. Sihto, H. Korhonen, H. Lihavainen, V. M. Kerminen, M. Kulmala, and Y. Viisanen (2006), New particle formation in air mass transported between two measurement sites in northern Finland, *Atmos. Chem. Phys.*, *6*, 2811–2824.
- Koren, I., Y. J. Kaufman, L. A. Remer, and J. V. Martins (2004), Measurement of the effect of Amazon smoke on inhibition of cloud formation, *Science*, *303*, 1342–1345, doi:10.1126/science.1089424.
- Koren, I., Y. J. Kaufman, D. Rosenfeld, L. A. Remer, and Y. Rudich (2005), Aerosol invigoration and restructuring of Atlantic convective clouds, *Geophys. Res. Lett.*, *32*, L14828, doi:10.1029/2005GL023187.
- Krejci, R., J. Strom, M. de Reus, P. Hoor, J. Williams, H. Fischer, and H. C. Hansson (2003), Evolution of aerosol properties over the rain forest in Surinam, South America, observed from aircraft during the LBA-CLAIRE 98 experiment, *J. Geophys. Res.*, *108* (D18), 4561, doi:10.1029/2001JD001375.
- Krejci, R., J. Strom, M. de Reus, J. Williams, H. Fischer, M. O. Andreae, and H. C. Hansson (2005), Spatial and temporal distribution of atmospheric aerosols in the lowermost troposphere over the Amazonian tropical rainforest, *Atmos. Chem. Phys.*, *5*, 1527–1543.
- Kuang, C., P. H. McMurry, A. V. McCormick, and F. L. Eisele (2008), Dependence of nucleation rates on sulfuric acid vapor concentrations in diverse atmospheric locations, *J. Geophys. Res.*, *113*, D10209, doi:10.1029/2007JD009253.
- Kuhn, U., et al. (2007), Isoprene and monoterpene fluxes from Central Amazonian rainforest inferred from tower-based and airborne measurements, and implications on the atmospheric chemistry and the local carbon budget, *Atmos. Chem. Phys.*, *7*, 2855–2879.
- Kulmala, M., H. Vehkamäki, T. Petaja, M. Dal Maso, A. Lauri, V. M. Kerminen, W. Birmili, and P. H. McMurry (2004), Formation and growth rates of ultrafine atmospheric particles: A review of observations, *J. Aerosol Sci.*, *35*, 143–176, doi:10.1016/j.jaerosci.2003.10.003.
- Laakso, L., et al. (2007), Hot-air balloon as a platform for boundary layer profile measurements during particle formation, *Boreal Environ. Res.*, *12*, 279–294.
- Laaksonen, A., et al. (2008), The role of VOC oxidation products in continental new particle formation, *Atmos. Chem. Phys.*, *8*, 2657–2665.
- Laurance, W. F., and G. B. Williamson (2001), Positive feedbacks among forest fragmentation, drought, and climate change in the Amazon, *Conserv. Biol.*, *15*, 1529–1535, doi:10.1046/j.1523-1739.2001.01093.x.
- Laurance, W. F., M. A. Cochrane, S. Bergen, P. M. Fearnside, P. Delamonica, C. Barber, S. D'Angelo, and T. Fernandes (2001), Environment—The future of the Brazilian Amazon, *Science*, *291*, 438–439, doi:10.1126/science.291.5503.438.
- Lee, A., A. H. Goldstein, J. H. Kroll, N. L. Ng, V. Varutbangkul, R. C. Flagan, and J. H. Seinfeld (2006), Gas-phase products and secondary aerosol yields from the photooxidation of 16 different terpenes, *J. Geophys. Res.*, *111*, D17305, doi:10.1029/2006JD007050.
- Lehtinen, K. E. J., M. Dal Maso, M. Kulmala, and V. M. Kerminen (2007), Estimating nucleation rates from apparent particle formation rates and vice versa: Revised formulation of the Kerminen-Kulmala equation, *J. Aerosol Sci.*, *38*, 988–994, doi:10.1016/j.jaerosci.2007.06.009.
- Lelieveld, J., et al. (2008), Atmospheric oxidation capacity sustained by a tropical forest, *Nature*, *452*, 737–740, doi:10.1038/nature06870.
- Lewis, S. L. (2006), Tropical forests and the changing Earth system, *Philos. Trans. R. Soc. London, Ser. B*, *361*, 195–210, doi:10.1098/rstb.2005.1711.
- Lim, H.-J., A. G. Carlton, and B. J. Turpin (2005), Isoprene forms secondary organic aerosol through cloud processing: Model simulations, *Environ. Sci. Technol.*, *39*, 4441–4446, doi:10.1021/es048039h.
- Liu, Y. Q. (2005), Atmospheric response and feedback to radiative forcing from biomass burning in tropical South America, *Agric. For. Meteorol.*, *133*, 40–53, doi:10.1016/j.agrformet.2005.03.011.
- Mace, K. A., P. Artaxo, and R. A. Duce (2003), Water-soluble organic nitrogen in Amazon Basin aerosols during the dry (biomass burning) and wet seasons, *J. Geophys. Res.*, *108*(D16), 4512, doi:10.1029/2003JD003557.
- Maenhaut, W., M. T. Fernandez-Jimenez, I. Rajta, and P. Artaxo (2002), Two-year study of atmospheric aerosols in Alta Floresta, Brazil: Multielemental composition and source apportionment, *Nucl. Instrum. Methods Phys. Res., Sect. B*, *189*, 243–248, doi:10.1016/S0168-583X(01)01050-3.
- Mahowald, N. M., P. Artaxo, A. R. Baker, T. D. Jickells, G. S. Okin, J. T. Randerson, and A. R. Townsend (2005), Impacts of biomass-burning emissions and land use change on Amazonian atmospheric phosphorus cycling and deposition, *Global Biogeochem. Cycles*, *19*, GB4030, doi:10.1029/2005GB002541.
- Malhi, Y., J. T. Roberts, R. A. Betts, T. J. Killeen, W. H. Li, and C. A. Nobre (2008), Climate change, deforestation, and the fate of the Amazon, *Science*, *319*, 169–172, doi:10.1126/science.1146961.
- Marengo, J. A., C. A. Nobre, and A. D. Culf (1997), Climatic impacts of “friagens” in forested and deforested areas of the Amazon Basin, *J. Appl. Meteorol.*, *36*, 1553–1566, doi:10.1175/1520-0450(1997)036<1553:CIOFIF>2.0.CO;2.
- Marengo, J. A., C. A. Nobre, J. Tomasella, M. D. Oyama, G. S. De Oliveira, R. De Oliveira, H. Camargo, L. M. Alves, and I. F. Brown (2008), The drought of Amazonia in 2005, *J. Clim.*, *21*, 495–516, doi:10.1175/2007JCLI1600.1.
- Martins, F. R., and E. B. Pereira (2006), Parameterization of aerosols from burning biomass in the Brazil-SR radiative transfer model, *Sol. Energy*, *80*, 231–239, doi:10.1016/j.solener.2005.03.008.
- Mayol-Bracero, O. L., P. Guyon, B. Graham, G. Roberts, M. O. Andreae, S. Decesari, M. C. Facchini, S. Fuzzi, and P. Artaxo (2002), Water-soluble organic compounds in biomass burning aerosols over Amazonia: 2. Apportionment of the chemical composition and importance of the polyacidic fraction, *J. Geophys. Res.*, *107*(D20), 8091, doi:10.1029/2001JD000522.
- McFiggans, G., et al. (2006), The effect of physical and chemical aerosol properties on warm cloud droplet activation, *Atmos. Chem. Phys.*, *6*, 2593–2649.
- McMurry, P. H., and S. K. Friedlander (1979), New particle formation in the presence of an aerosol, *Atmos. Environ.*, *13*, 1635–1651, doi:10.1016/0004-6981(79)90322-6.
- McMurry, P. H., et al. (2005), A criterion for new particle formation in the sulfur-rich Atlanta atmosphere, *J. Geophys. Res.*, *110*, D22S02, doi:10.1029/2005JD005901.
- Mircea, M., et al. (2005), Importance of the organic aerosol fraction for modeling aerosol hygroscopic growth and activation: A case study in the Amazon Basin, *Atmos. Chem. Phys.*, *5*, 3111–3126.
- Mirme, A., E. Tamm, G. Mordas, M. Vana, J. Uin, S. Mirme, T. Bernotas, L. Laakso, A. Hirsikko, and M. Kulmala (2007), A wide-range multi-channel air ion spectrometer, *Boreal Environ. Res.*, *12*, 247–264.
- Moteki, N., Y. Kondo, Y. Miyazaki, N. Takegawa, Y. Komazaki, G. Kurata, T. Shirai, D. R. Blake, T. Miyakawa, and M. Koike (2007), Evolution of mixing state of black carbon particles: Aircraft measurements over the western Pacific in March 2004, *Geophys. Res. Lett.*, *34*, L11803, doi:10.1029/2006GL028943.
- Nepstad, D. C. (2007), The Amazon's vicious cycles: A report to the World Wide Fund for Nature (WWF), World Wide Fund for Nature, Gland, Switzerland.
- Nepstad, D. C., C. M. Stickler, B. Soares, and F. Merry (2008), Interactions among Amazon land use, forests and climate: Prospects for a near-term forest tipping point, *Philos. Trans.*

- R. Soc. London, Ser. B*, 363, 1737–1746, doi:10.1098/rstb.2007.0036.
- Ng, N. L., J. H. Kroll, M. D. Keywood, R. Bahreini, V. Varutbangkul, R. C. Flagan, J. H. Seinfeld, A. Lee, and A. H. Goldstein (2006), Contribution of first- versus second-generation products to secondary organic aerosols formed in the oxidation of biogenic hydrocarbons, *Environ. Sci. Technol.*, 40, 2283–2297, doi:10.1021/es052269u.
- Nobre, C. A., M. A. Silva Dias, A. D. Culf, J. A. Polcher, J. H. C. Gash, J. A. Marengo, and R. Avissar (2004), The Amazonian climate, in *Vegetation, Water, Humans and the Climate: A New Perspective on an Interactive System*, edited by P. Kabat et al., pp. 79–92, Springer, Berlin.
- Nogues-Paegle, J., and K. C. Mo (1997), Alternating wet and dry conditions over South America during summer, *Mon. Weather Rev.*, 125, 279–291, doi:10.1175/1520-0493(1997)125<0279:AWADCO>2.0.CO;2.
- O'Dowd, C. D., M. C. Facchini, F. Cavalli, D. Ceburnis, M. Mircea, S. Decesari, S. Fuzzi, Y. J. Yoon, and J. P. Putaud (2004), Biogenically driven organic contribution to marine aerosol, *Nature*, 431, 676–680, doi:10.1038/nature02959.
- Oyama, M. D., and C. A. Nobre (2003), A new climate-vegetation equilibrium state for tropical South America, *Geophys. Res. Lett.*, 30(23), 2199, doi:10.1029/2003GL018600.
- Palmer, P. I., D. J. Jacob, A. M. Fiore, R. V. Martin, K. Chance, and T. P. Kurosu (2003), Mapping isoprene emissions over North America using formaldehyde column observations from space, *J. Geophys. Res.*, 108(D6), 4180, doi:10.1029/2002JD002153.
- Palmer, P. I., M. P. Barkley, T. P. Kurosu, A. C. Lewis, J. E. Saxton, K. Chance, and L. V. Gatti (2007), Interpreting satellite column observations of formaldehyde over tropical South America, *Philos. Trans. R. Soc. A*, 365, 1741–1751, doi:10.1098/rsta.2007.2042.
- Pauliquevis, T., L. L. Lara, M. L. Antunes, and P. Artaxo (2007), Aerosol and precipitation chemistry in a remote site in Central Amazonia: The role of biogenic contribution, *Atmos. Chem. Phys. Discuss.*, 7, 11,465–11,509.
- Penner, J. E., S. J. Ghan, and J. J. Walton (1991), The role of biomass burning in the budget and cycle of carbonaceous soot aerosols and their climate impact, in *Global Biomass Burning*, edited by J. Levine, pp. 432–438, MIT Press, Cambridge, Mass.
- Perry, K. D., and P. V. Hobbs (1994), Further evidence for particle nucleation in clear air adjacent to marine cumulus clouds, *J. Geophys. Res.*, 99, 22,803–22,818, doi:10.1029/94JD01926.
- Platnick, S., and S. Twomey (1994), Determining the susceptibility of cloud albedo to changes in droplet concentration with the advanced very high-resolution radiometer, *J. Appl. Meteorol.*, 33, 334–347, doi:10.1175/1520-0450(1994)033<0334:DTSOCA>2.0.CO;2.
- Pope, C. A., and D. W. Dockery (2006), Health effects of fine particulate air pollution: Lines that connect, *J. Air Waste Manage. Assoc.*, 56, 709–742.
- Pöschl, U. (2005), Atmospheric aerosols: Composition, transformation, climate and health effects, *Angew. Chem.*, 44, 7520–7540, doi:10.1002/anie.200501122.
- Posfai, M., R. Simonics, J. Li, P. V. Hobbs, and P. R. Buseck (2003), Individual aerosol particles from biomass burning in southern Africa: 1. Compositions and size distributions of carbonaceous particles, *J. Geophys. Res.*, 108(D13), 8483, doi:10.1029/2002JD002291.
- Posfai, M., A. Gelencser, R. Simonics, K. Arato, J. Li, P. V. Hobbs, and P. R. Buseck (2004), Atmospheric tar balls: Particles from biomass and biofuel burning, *J. Geophys. Res.*, 109, D06213, doi:10.1029/2003JD004169.
- Prather, K. A., T. Nordmeyer, and K. Salt (1994), Real-time characterization of individual aerosol-particles using time-of-flight mass-spectrometry, *Anal. Chem.*, 66, 1403–1407, doi:10.1021/ac00081a007.
- Prenni, A. J., M. D. Petters, S. M. Kreidenweis, and P. J. DeMott (2007), Cloud droplet activation of secondary organic aerosol, *J. Geophys. Res.*, 112, D10223, doi:10.1029/2006JD007963.
- Prenni, A. J., M. D. Petters, S. M. Kreidenweis, C. L. Heald, S. T. Martin, P. Artaxo, R. M. Garland, A. G. Wollny, and U. Poeschl (2009), Relative roles of biogenic emissions and Saharan dust as ice nuclei in the Amazon Basin, *Nat. Geosci.*, 2, 402–405, doi:10.1038/ngeo517.
- Prospero, J. M., R. A. Glaccum, and R. T. Nees (1981), Atmospheric transport of soil dust from Africa to South America, *Nature*, 289, 570–572, doi:10.1038/289570a0.
- Rasmussen, R. A., and M. A. K. Khalil (1988), Isoprene over the Amazon Basin, *J. Geophys. Res.*, 93, 1417–1421, doi:10.1029/JD093iD02p01417.
- Reid, J. S., R. Koppmann, T. F. Eck, and D. P. Eleuterio (2005), A review of biomass-burning emissions part II: Intensive physical properties of biomass-burning particles, *Atmos. Chem. Phys.*, 5, 799–825.
- Reinhardt, T. E., R. D. Ottmar, and C. Castilla (2001), Smoke impacts from agricultural burning in a rural Brazilian town, *J. Air Waste Manage. Assoc.*, 51, 443–450.
- Riipinen, I., et al. (2007), Connections between atmospheric sulphuric acid and new particle formation during QUEST III–IV campaigns in Hyytiälä and Heidelberg, *Atmos. Phys. Chem.*, 7, 1899–1914.
- Rissler, J., E. Swietlicki, J. Zhou, G. Roberts, M. O. Andreae, L. V. Gatti, and P. Artaxo (2004), Physical properties of the sub-micrometer aerosol over the Amazon rain forest during the wet-to-dry season transition—Comparison of modeled and measured CCN concentrations, *Atmos. Chem. Phys.*, 4, 2119–2143.
- Rissler, J., A. Vestin, E. Swietlicki, G. Fisch, J. Zhou, P. Artaxo, and M. O. Andreae (2006), Size distribution and hygroscopic properties of aerosol particles from dry-season biomass burning in Amazonia, *Atmos. Chem. Phys.*, 6, 471–491.
- Roberts, G. C., and A. Nenes (2005), A continuous-flow stream-wise thermal-gradient CCN chamber for atmospheric measurements, *Aerosol Sci. Technol.*, 39, 206–221, doi:10.1080/027868290913988.
- Roberts, G. C., M. O. Andreae, J. C. Zhou, and P. Artaxo (2001), Cloud condensation nuclei in the Amazon Basin: “Marine” conditions over a continent?, *Geophys. Res. Lett.*, 28, 2807–2810, doi:10.1029/2000GL012585.
- Roberts, G. C., P. Artaxo, J. C. Zhou, E. Swietlicki, and M. O. Andreae (2002), Sensitivity of CCN spectra on chemical and physical properties of aerosol: A case study from the Amazon Basin, *J. Geophys. Res.*, 107(D20), 8070, doi:10.1029/2001JD000583.
- Roberts, G., A. Nenes, J. Seinfeld, and M. Andreae (2003), Impact of biomass burning on cloud properties in the Amazon Basin, *J. Geophys. Res.*, 108(D2), 4062, doi:10.1029/2001JD000985.
- Rosenfeld, D., U. Lohmann, G. B. Raga, C. D. O'Dowd, M. Kulmala, S. Fuzzi, A. Reissell, and M. O. Andreae (2008), Flood or drought: How do aerosols affect precipitation?, *Science*, 321, 1309–1313, doi:10.1126/science.1160606.
- Sachse, G. W., R. C. Harriss, J. Fishman, G. F. Hill, and D. R. Cahoon (1988), Carbon-monoxide over the Amazon Basin during the 1985 dry season, *J. Geophys. Res.*, 93, 1422–1430, doi:10.1029/JD093iD02p01422.
- Saleska, S. R., K. Didan, A. R. Huete, and H. R. da Rocha (2007), Amazon forests green-up during 2005 drought, *Science*, 318, 612, doi:10.1126/science.1146663.
- Satyamurty, P., C. A. Nobre, and P. S. Dias (1998), South America, in *Meteorology of the Southern Hemisphere*, edited by D. Karoly and D. G. Vincent, *Meteorol. Monogr.*, 27, 119–139.
- Schafer, J. S., T. F. Eck, B. N. Holben, P. Artaxo, and A. F. Duarte (2008), Characterization of the optical properties of atmospheric aerosols in Amazonia from long-term AERONET monitoring (1993–1995 and 1999–2006), *J. Geophys. Res.*, 113, D04204, doi:10.1029/2007JD009319.

- Schkolnik, G., A. H. Falkovich, Y. Rudich, W. Maenhaut, and P. Artaxo (2005), New analytical method for the determination of levoglucosan, polyhydroxy compounds, and 2-methylerythritol and its application to smoke and rainwater samples, *Environ. Sci. Technol.*, *39*, 2744–2752, doi:10.1021/es048363c.
- Schwartz, J., F. Laden, and A. Zanobetti (2002), The concentration-response relation between PM_{2.5} and daily deaths, *Environ. Health Perspect.*, *110*, 1025–1029.
- Schwarz, J. P., et al. (2006), Single-particle measurements of mid-latitude black carbon and light-scattering aerosols from the boundary layer to the lower stratosphere, *J. Geophys. Res.*, *111*, D16207, doi:10.1029/2006JD007076.
- Shilling, J. E., Q. Chen, S. M. King, T. Rosenoern, J. H. Kroll, D. R. Worsnop, K. A. McKinney, and S. T. Martin (2008), Particle mass yield in secondary organic aerosol formed by the dark ozonolysis of alpha-pinene, *Atmos. Chem. Phys.*, *8*, 2073–2088.
- Shim, C., Y. H. Wang, Y. Choi, P. I. Palmer, D. S. Abbot, and K. Chance (2005), Constraining global isoprene emissions with Global Ozone Monitoring Experiment (GOME) formaldehyde column measurements, *J. Geophys. Res.*, *110*, D24301, doi:10.1029/2004JD005629.
- Silva Dias, M. A. F., et al. (2002), Cloud and rain processes in a biosphere-atmosphere interaction context in the Amazon region, *J. Geophys. Res.*, *107*(D20), 8072, doi:10.1029/2001JD000335.
- Sinha, B. W., et al. (2009), Composition and mixing state of wet season fine mode aerosol collected in the Amazonian tropical rain forest, paper presented at General Assembly, Eur. Geosci. Union, Manaus, Brazil.
- Smith, J. N., M. J. Dunn, T. M. VanReken, K. Iida, M. R. Stolzenburg, P. H. McMurry, and L. G. Huey (2008), Chemical composition of atmospheric nanoparticles formed from nucleation in Tecamac, Mexico: Evidence for an important role for organic species in nanoparticle growth, *Geophys. Res. Lett.*, *35*, L04808, doi:10.1029/2007GL032523.
- Soares-Filho, B. S., D. C. Nepstad, L. M. Curran, G. C. Cerqueira, R. A. Garcia, C. A. Ramos, E. Voll, A. McDonald, P. Lefebvre, and P. Schlesinger (2006), Modelling conservation in the Amazon Basin, *Nature*, *440*, 520–523, doi:10.1038/nature04389.
- Solomon, P. A., and C. Sioutas (2008), Continuous and semicontinuous monitoring techniques for particulate matter mass and chemical components: A synthesis of findings from EPA's particulate matter supersites program and related studies, *J. Air Waste Manage. Assoc.*, *58*, 164–195, doi:10.3155/1047-3289.58.2.164.
- Spracklen, D. V., K. J. Pringle, K. S. Carslaw, M. P. Chipperfield, and G. W. Mann (2005), A global off-line model of size-resolved aerosol microphysics: I. Model development and prediction of aerosol properties, *Atmos. Chem. Phys.*, *5*, 2227–2252.
- Spracklen, D. V., K. S. Carslaw, M. Kulmala, V. M. Kerminen, G. W. Mann, and S. L. Sihto (2006), The contribution of boundary layer nucleation events to total particle concentrations on regional and global scales, *Atmos. Chem. Phys.*, *6*, 5631–5648.
- Stanier, C. O., A. Y. Khlystov, and S. N. Pandis (2004), Nucleation events during the Pittsburgh air quality study: Description and relation to key meteorological, gas phase, and aerosol parameters, *Aerosol Sci. Technol.*, *38*, 253–264, doi:10.1080/02786820390229570.
- Stavrakou, T., J.-F. Müller, I. D. Smedt, M. V. Roozendaal, G. R. van der Werf, L. Giglio, and A. Guenther (2009), Global emissions of non-methane hydrocarbons deduced from SCIAMACHY formaldehyde columns through 2003–2006, *Atmos. Chem. Phys.*, *9*, 3663–3679.
- Stolzenburg, M. R., P. H. McMurry, H. Sakurai, J. N. Smith, R. L. Mauldin, F. L. Eisele, and C. F. Clement (2005), Growth rates of freshly nucleated atmospheric particles in Atlanta, *J. Geophys. Res.*, *110*, D22S05, doi:10.1029/2005JD005935.
- Stratmann, F., et al. (2003), New-particle formation events in a continental boundary layer: First results from the SATURN experiment, *Atmos. Chem. Phys.*, *3*, 1445–1459.
- Streets, D. G. (2007), Dissecting future aerosol emissions: Warming tendencies and mitigation opportunities, *Clim. Change*, *81*, 313–330, doi:10.1007/s10584-006-9112-8.
- Suni, T., et al. (2008), Formation and characteristics of ions and charged aerosol particles in a native Australian Eucalypt forest, *Atmos. Chem. Phys.*, *8*, 129–139.
- Surratt, J. D., et al. (2008), Organosulfate formation in biogenic secondary organic aerosol, *J. Phys. Chem. A*, *112*, 8345–8378, doi:10.1021/jp802310p.
- Svenningsson, B., et al. (2006), Hygroscopic growth and critical supersaturations for mixed aerosol particles of inorganic and organic compounds of atmospheric relevance, *Atmos. Chem. Phys.*, *6*, 1937–1952.
- Swap, R., M. Garstang, S. Greco, R. Talbot, and P. Kållberg (1992), Saharan dust in the Amazon Basin, *Tellus, Ser. B*, *44*, 133–149.
- Swietlicki, E., et al. (2008), Hygroscopic properties of submicrometer atmospheric aerosol particles measured with H-TDMA instruments in various environments—A review, *Tellus, Ser. B*, *60*, 432–469.
- Talbot, R. W., M. O. Andreae, T. W. Andreae, and R. C. Harriss (1988), Regional aerosol chemistry of the Amazon Basin during the dry season, *J. Geophys. Res.*, *93*, 1499–1508, doi:10.1029/JD093iD02p01499.
- Talbot, R. W., M. O. Andreae, H. Berresheim, P. Artaxo, M. Garstang, R. C. Harriss, K. M. Beecher, and S. M. Li (1990), Aerosol chemistry during the wet season in central Amazonia: The influence of long-range transport, *J. Geophys. Res.*, *95*, 16,955–16,969, doi:10.1029/JD095iD10p16955.
- Tegen, I., P. Hollrig, M. Chin, I. Fung, D. Jacob, and J. Penner (1997), Contribution of different aerosol species to the global aerosol extinction optical thickness: Estimates from model results, *J. Geophys. Res.*, *102*, 23,895–23,915, doi:10.1029/97JD01864.
- Textor, C., et al. (2006), Analysis and quantification of the diversities of aerosol life cycles within AeroCom, *Atmos. Chem. Phys.*, *6*, 1777–1813.
- Tie, X. X., S. Madronich, S. Walters, D. P. Edwards, P. Ginoux, N. Mahowald, R. Y. Zhang, C. Lou, and G. Brasseur (2005), Assessment of the global impact of aerosols on tropospheric oxidants, *J. Geophys. Res.*, *110*, D03204, doi:10.1029/2004JD005359.
- Trebs, I., F. X. Meixner, J. Slanina, R. Otjes, P. Jongejan, and M. O. Andreae (2004), Real-time measurements of ammonia, acidic trace gases and water-soluble inorganic aerosol species at a rural site in the Amazon Basin, *Atmos. Chem. Phys.*, *4*, 967–987.
- Trebs, I., et al. (2005), The NH₄⁺NO₃⁻Cl⁻SO₄²⁻-H₂O aerosol system and its gas phase precursors at a pasture site in the Amazon Basin: How relevant are mineral cations and soluble organic acids?, *J. Geophys. Res.*, *110*, D07303, doi:10.1029/2004JD005478.
- Trebs, I., et al. (2008), Aerosol inorganic composition at a tropical site: Discrepancies between filter-based sampling and a semi-continuous method, *Aerosol Sci. Technol.*, *42*, 255–269, doi:10.1080/02786820801992899.
- Tsigaridis, K., and M. Kanakidou (2003), Global modelling of secondary organic aerosol in the troposphere: A sensitivity analysis, *Atmos. Chem. Phys.*, *3*, 1849–1869.
- Twomey, S. (1977), The influence of pollution on the short-wave albedo of clouds, *J. Atmos. Sci.*, *34*, 1149–1152, doi:10.1175/1520-0469(1977)034<1149:TIOPOT>2.0.CO;2.
- Vana, M., M. Kulmala, M. Dal Maso, U. Horrak, and E. Tamm (2004), Comparative study of nucleation mode aerosol particles and intermediate air ions formation events at three sites, *J. Geophys. Res.*, *109*, D17201, doi:10.1029/2003JD004413.
- Vestin, A., J. Rissler, E. Swietlicki, G. P. Frank, and M. O. Andreae (2007), Cloud-nucleating properties of the Amazonian biomass burning aerosol: Cloud condensation nuclei measurements and

- modeling, *J. Geophys. Res.*, *112*, D14201, doi:10.1029/2006JD008104.
- Volkamer, R., F. S. Martini, L. T. Molina, D. Salcedo, J. L. Jimenez, and M. J. Molina (2007), A missing sink for gas-phase glyoxal in Mexico City: Formation of secondary organic aerosol, *Geophys. Res. Lett.*, *34*, L19807, doi:10.1029/2007GL030752.
- Wang, H., and R. Fu (2002), Cross-equatorial flow and seasonal cycle of precipitation over South America, *J. Clim.*, *15*, 1591–1608, doi:10.1175/1520-0442(2002)015<1591:CEFASC>2.0.CO;2.
- Watson, J. G. (2002), Visibility: Science and regulation, *J. Air Waste Manage. Assoc.*, *52*, 628–713.
- Went, F. W. (1960), Blue hazes in the atmosphere, *Nature*, *187*, 641–643, doi:10.1038/187641a0.
- Williams, E., et al. (2002), Contrasting convective regimes over the Amazon: Implications for cloud electrification, *J. Geophys. Res.*, *107*(D20), 8082, doi:10.1029/2001JD000380.
- Worobiec, A., I. Szaloki, J. Osan, W. Maenhaut, E. A. Stefaniak, and R. Van Grieken (2007), Characterisation of Amazon Basin aerosols at the individual particle level by X-ray microanalytical techniques, *Atmos. Environ.*, *41*, 9217–9230, doi:10.1016/j.atmosenv.2007.07.056.
- Yamasoe, M. A., P. Artaxo, A. H. Miguel, and A. G. Allen (2000), Chemical composition of aerosol particles from direct emissions of vegetation fires in the Amazon Basin: Water-soluble species and trace elements, *Atmos. Environ.*, *34*, 1641–1653, doi:10.1016/S1352-2310(99)00329-5.
- Yin, Y., Z. Levin, T. G. Reisin, and S. Tzivion (2000), The effects of giant cloud condensation nuclei on the development of precipitation in convective clouds—A numerical study, *Atmos. Res.*, *53*, 91–116, doi:10.1016/S0169-8095(99)00046-0.
- Yokelson, R. J., T. Karl, P. Artaxo, D. R. Blake, T. J. Christian, D. W. T. Griffith, A. Guenther, and W. M. Hao (2007), The Tropical Forest and Fire Emissions Experiment: Overview and airborne fire emission factor measurements, *Atmos. Chem. Phys.*, *7*, 5175–5196.
- Yokelson, R. J., T. J. Christian, T. G. Karl, and A. Guenther (2008), The tropical forest and fire emissions experiment: Laboratory fire measurements and synthesis of campaign data, *Atmos. Chem. Phys.*, *8*, 3509–3527.
- Yu, F., Z. Wang, G. Luo, and R. Turco (2008), Ion-mediated nucleation as an important global source of tropospheric aerosols, *Atmos. Chem. Phys.*, *8*, 2537–2554.
- Zdrahal, Z., J. Oliveira, R. Vermeylen, M. Claeys, and W. Maenhaut (2002), Improved method for quantifying levoglucosan and related monosaccharide anhydrides in atmospheric aerosols and application to samples from urban and tropical locations, *Environ. Sci. Technol.*, *36*, 747–753, doi:10.1021/es015619v.
- Zhang, Y., R. Fu, H. B. Yu, Y. Qian, R. Dickinson, M. Dias, P. L. D. Dias, and K. Fernandes (2009), Impact of biomass burning aerosol on the monsoon circulation transition over Amazonia, *Geophys. Res. Lett.*, *36*, L10814, doi:10.1029/2009GL037180.
- Zhou, J. C., E. Swietlicki, H. C. Hansson, and P. Artaxo (2002), Submicrometer aerosol particle size distribution and hygroscopic growth measured in the Amazon rain forest during the wet season, *J. Geophys. Res.*, *107*(D20), 8055, doi:10.1029/2000JD000203.
- Zhou, J. Y., and K. M. Lau (1998), Does a monsoon climate exist over South America?, *J. Clim.*, *11*, 1020–1040, doi:10.1175/1520-0442(1998)011<1020:DAMCEO>2.0.CO;2.
- Ziese, M., H. Wex, E. Nilsson, I. Salma, R. Ocskay, T. Hennig, A. Massling, and F. Stratmann (2008), Hygroscopic growth and activation of HULIS particles: Experimental data and a new iterative parameterization scheme for complex aerosol particles, *Atmos. Chem. Phys.*, *8*, 1855–1866.
- Zimmerman, P. R., J. P. Greenberg, and C. E. Westberg (1988), Measurements of atmospheric hydrocarbons and biogenic emission fluxes in the Amazon boundary layer, *J. Geophys. Res.*, *93*, 1407–1416, doi:10.1029/JD093iD02p01407.

M. O. Andreae, U. Pöschl, and I. Trebs, Biogeochemistry Department, Max Planck Institute for Chemistry, D-55128 Mainz, Germany.

P. Artaxo, Institute of Physics, University of São Paulo, 05315-970 São Paulo, Brazil.

D. Baumgardner, Centro de Ciencias de la Atmosfera, Universidad Nacional Autónoma de México, 04510 México City, México.

Q. Chen and S. T. Martin, School of Engineering and Applied Sciences, Harvard University, Cambridge, MA 02138, USA. (scot_martin@harvard.edu)

A. H. Goldstein, Department of Environmental Science, Policy, and Management, University of California, Berkeley, CA 94720, USA.

A. Guenther, Biosphere-Atmosphere Interactions Group, Atmospheric Chemistry Division, National Center for Atmospheric Research, 1850 Table Mesa Dr., Boulder, CO 80307, USA.

C. L. Heald, Department of Atmospheric Science, Colorado State University, Fort Collins, CO 80523, USA.

O. L. Mayol-Bracero, Institute for Tropical Ecosystem Studies, University of Puerto Rico, San Juan, PR 00936, Puerto Rico.

P. H. McMurry, Department of Mechanical Engineering, University of Minnesota, Rm. 1101 MechE, Minneapolis, MN 55455, USA.

T. Pauliquevis, Climate and Environmental Modelling Group, Instituto Nacional de Pesquisas da Amazônia, 69060-001, Manaus, Brazil.

K. A. Prather, Department of Chemistry and Biochemistry, University of California, San Diego, La Jolla, CA 92093, USA.

G. C. Roberts, Center for Atmospheric Sciences, Scripps Institution of Oceanography, University of California, San Diego, La Jolla, CA 92037, USA.

S. R. Saleska, Department of Ecology and Evolutionary Biology, University of Arizona, Tucson, AZ 85721, USA.

M. A. Silva Dias, Center for Weather Forecasting and Climate Studies, National Institute for Space Research, University of São Paulo, 12630-000 São Paulo, Brazil.

D. V. Spracklen, Institute for Climate and Atmospheric Science, School of Earth and Environment, University of Leeds, Leeds LS2 9JT, UK.

E. Swietlicki, Department of Physics, Lund University, SE-221 00 Lund, Sweden.